

1 **Integrating Saharan dust forecasts into a regional chemical transport model:**
2 **a case study over Northern Italy**

3

4 **C. Carnevale^{1*}, G. Finzi¹, E. Pisoni¹, M. Volta¹, P. Kishcha², P. Alpert²**

5

6 ¹*Dept. of Information Engineering, Faculty of Engineering, University of Brescia, Italy.*

7

**Corresponding author: carneval@ing.unibs.it*

8

²*Department of Geophysics and Planetary Sciences, Tel-Aviv University,*

9

69978 Tel-Aviv, Israel

10

11 **Abstract**

12 The Po Valley in Northern Italy is frequently affected by high PM10 concentrations,
13 where both natural and anthropogenic sources play a significant role. To improve air
14 pollution modeling, 3D dust fields, produced by means of the DREAM dust forecasts,
15 were integrated as boundary conditions into the mesoscale 3D deterministic Transport
16 Chemical Aerosol Model (TCAM). A case study of the TCAM and DREAM integration
17 was implemented over Northern Italy for the period May 15 – June 30, 2007. First, the
18 Saharan dust impact on PM10 concentration was analyzed for eleven remote PM10 sites
19 with the lowest level of air pollution. These remote sites are the most sensitive to Saharan
20 dust intrusions into Northern Italy, because of the absence of intensive industrial
21 pollution. At these remote sites, the observed maxima in PM10 concentration during dust
22 events is evidence of dust aerosol near the surface in Northern Italy. Comparisons

23 between modeled PM10 concentrations and measurements at 230 PM10 sites in Northern
24 Italy, showed that the integrated TCAM-DREAM model more accurately reproduced
25 PM10 concentration than the base TCAM model, both in terms of correlation and mean
26 error. Specifically, the correlation median increased from 0.40 to 0.65, while the
27 normalized mean absolute error median dropped from 0.5 to 0.4.

28

29 Keywords: Aerosol, Saharan dust, Dust forecast, Multiphase chemical transport model.

30

31 **1. Introduction**

32 According to the EU Directive 2008/50, where exceedance of the pollution limit can be
33 attributed to natural sources rather than to anthropogenic sources, the exceedance can be
34 discounted after scientific validation, by European Union State Members,. In connection
35 with this, in recent years, a number of experimental studies have been performed to
36 investigate the contribution of Saharan dust to PM10 concentrations in the Euro-
37 Mediterranean basin (Gobbi et al., 2007, Escudero et al., 2007, Masson et al., 2010,
38 Perrino et al., 2008, Pederzoli et al., 2010, Rodriguez et al., 2001).

39

40 Saharan dust is frequently transported over the Italian Peninsula, especially during the
41 summer season (Pederzoli et al., 2010). Previous studies, based on lidar measurements of
42 dust vertical distribution over Rome, showed that, on average, the dust layer is far from
43 the surface and is located within a wide range of altitude from 0.5 km to 8 km (Kishcha et
44 al., 2005). However, in accordance with recent experimental studies, carried out
45 independently by Gobbi et al. (2007), Perrino et al. (2008), and Pederzoli et al. (2010),
46 some mixing of dust with local aerosols below the dust layer bottom boundary is often
47 highly probable. In particular, on approximately 30% of days in 2001, dust contribution
48 to daily PM10 levels in Rome was estimated at about $20 \mu\text{g m}^{-3}$ (Gobbi et al., 2007). In
49 accordance with chemical analysis of PM10 measurements in Central Italy, Saharan dust
50 contributes as much as 32 % to the total PM10 mass during dust transport events (Perrino
51 et al., 2008). The contribution of Saharan dust to PM10 monthly concentrations, as much

52 as 8 - 10 $\mu\text{g m}^{-3}$, was estimated at seven Italian locations, from the Ispra site in North Italy
53 to the Lampedusa site in South Italy (Pederzoli et al., 2010). Therefore, it is essential to
54 give correct model estimates of the contribution of Saharan dust to air quality in Italy.

55

56 Regional dust models are considered to be a useful tool for simulating dust contribution
57 to surface aerosol concentration, where dust transport is significant (Kishcha et al., 2008,
58 Papanastasiou et al., 2010, Pederzoli et al., 2010). However, despite the large number of
59 available chemical weather forecasting systems for simulating aerosol contribution to air
60 quality in Europe (<http://chemicalweather.eu/Links>), few of these models include Saharan
61 dust. On the other hand, multiphase models, simulating the physical-chemical processes
62 involving secondary pollutants in the troposphere, are key tools for evaluating the
63 effectiveness of emission control strategies and thereby for the improvement in air
64 quality. Due to the complexity of phenomena involved in the formation and accumulation
65 of aerosols, these models require detailed input in terms of meteorological fields,
66 emission sources and boundary conditions. If we improve boundary condition fields on
67 Saharan dust transport, we could give information to our mesoscale model about external
68 pollutant fluxes entering the model domain.

69

70 This study was aimed at giving better model estimates of air pollution in Northern Italy.
71 Saharan dust forecasts over the Mediterranean region, daily computed at Tel-Aviv
72 University (Kishcha et al., 2008) were used to improve boundary conditions of the
73 mesoscale multiphase Transport Chemical Aerosol Model (TCAM) model, developed at

74 the University of Brescia (Carnevale et al., 2008). Integrating Saharan dust forecasts into
75 TCAM allowed us to obtain better model estimates of PM10 concentration.

76

77 **2. Modelling System Configuration**

78 Two different configurations of the Gas Aerosol Modeling Evaluation System (GAMES)
79 (Volta and Finzi, 2006) are considered in this study. In the base configuration (Fig. 1), the
80 modeling system includes: the TCAM chemical transport model (Carnevale et al., 2008);
81 the meteorological pre-processor PROMETEO; the emission processor POEM-PM
82 (Carnevale et al., 2006); and a pre-processor computing the boundary conditions on the
83 basis of CHIMERE continental scale simulations (Schmidt et al. 2001). In the second
84 model configuration, Saharan dust forecasts were integrated into the base model system,
85 in order to improve boundary conditions on large-scale transport of dust over the model
86 domain (Fig. 2). In this second model configuration, dust was a passive component: no
87 interaction between dust and other aerosol species was included.

88

89 **2.1 TCAM model**

90 TCAM is a 3-D multiphase Eulerian model which solves a system of partial differential
91 equations for each time step. TCAM describes horizontal/vertical transport; multiphase
92 chemical reactions; and gas-to-particle conversion phenomena using a splitting operator
93 technique (Marchuk, 1975). Horizontal transport is solved using a chapeau function
94 approximation (Pepper et al., 1979) and a nonlinear Forester filter (Forester, 1977), while
95 the vertical transport PDE system is solved using a hybrid implicit/explicit scheme. Gas

96 chemistry is simulated using a modified version of the SAPRC97 scheme (Carnevale et
97 al., 2010). The ODE chemical kinetic system is solved by means of the Implicit-Explicit
98 Hybrid (IEH) solver (Chock et al., 1994), which splits the species into fast and slow ones,
99 according to their reaction rates. The system of fast species is solved by means of the
100 implicit Livermore Solver for Ordinary Differential Equations (LSODE) (Hindmarsh,
101 1975) implementing the Adams predictor/corrector method in a non-stiff case and the
102 Backward Differentiation Formula method in a stiff case (Carnevale et al., 2008). The
103 slow species system is solved using the Adams-Bashfort method (Carnevale et al., 2008).
104 The aerosol module simulates the particles by means of a fixed-moving approach. A
105 generic particle is represented by an internal core containing non-volatile material, such
106 as elemental carbon and crustal aerosol. The core dimension of each size class is
107 established at the beginning of the simulation and is held constant during all the
108 computations. The volatile material is assumed to reside in the outer shell of the particle
109 whose dimension is evaluated by the module at each time step, on the basis of the total
110 mass and of the total number of suspended particles. The aerosol module describes the
111 dynamics of 21 chemical compounds: twelve inorganic species (H_2O , SO_4^{2-} , NH_4^+ , Cl^- ,
112 NO_3^- , Na^+ , H^+ , SO_2 (aq), H_2O_2 (aq), O_3 (aq), elemental carbon and others), and 9
113 organic species consisting of one generic primary organic species and 8 classes of
114 secondary organic species. Each chemical species is split into n (namely n=10) size bins.
115 The inorganic species thermodynamic equilibrium is solved using ISORROPIA-II (Nenes
116 et al., 1998), (Fountoukis and Nenes, 2007), (<http://nenes.eas.gatech.edu/ISORROPIA>).

117 The TCAM model has been widely used and validated in the frame of a number of
118 national and international projects (Cuvelier et al., 2007) (Carnevale et al., 2008) (Di
119 Nicolantonio et al. 2009).

120 **2.2 DREAM desert dust forecasts**

121 DREAM is a model designed to predict the atmospheric cycle of mineral dust aerosol
122 (Nickovic et al., 2001). It solves the Euler-type partial differential nonlinear equation for
123 dust mass continuity. The NCEP/Eta regional atmospheric model drives the aerosol.
124 During the model integration, calculations of the surface dust injection fluxes are made
125 over the model cells declared as deserts. DREAM includes dust sources located in the
126 Western, Central, and Eastern Sahara, as well as in the Arabian Peninsula. Wind erosion
127 of the soil in the DREAM parameterization scheme is controlled by several factors, such
128 as the type of soil (soil texture), type of vegetation cover, soil moisture content, and
129 surface atmospheric turbulence. Once injected into the air, dust aerosol is driven by
130 turbulence in the early stage of the process, when dust is lifted from the ground to the
131 upper levels, and subsequently by winds in the later phases of the process, when dust
132 travels away from the sources. DREAM includes processes of wet (Giorgi, 1986) and dry
133 deposition. Eight size bins covering the particle effective sizes from 0.1 μm to 8.0 μm are
134 used in the model.

135

136 DREAM has been used for several years for producing operational dust forecasts at Tel-
137 Aviv University, Israel. Several publications on the DREAM model validation against
138 lidar measurements, AERONET data, and PM10 measurements are available (Amiridis et

139 al., 2009, Kishcha et al., 2007, Papanastasiou et al., 2010, Papayannis et al., 2007, Perez
140 et al., 2006). At Tel-Aviv University, since the year 2006, DREAM has been producing
141 daily forecasts of 3-D distribution of dust concentrations over the Mediterranean region,
142 the Middle East, Europe, and the Atlantic Ocean (<http://wind.tau.ac.il/dust8/dust.html>). In
143 the current study, dust forecasts over the Mediterranean domain (20°W – 45°E, 15°N –
144 50°N) were used, with horizontal resolution of 0.3° and 24 vertical levels between 86 m
145 and approximately 15 km above the sea level. DREAM is initialized with the NCEP
146 analysis and the lateral boundary data are updated every six hours by the NCEP GFS
147 model. The runs start at 1200 UTC and forecasts are performed for 3-hour periods up to
148 72 hours ahead. With respect to dust, DREAM is initialized with 3-D dust distributions
149 from previous days.

150 **3. The case study**

151 **3.1 Dust event characterization**

152 During the period under consideration in the current study from May 15 to June 30, 2007,
153 four dust events were observed over Northern Italy: May 20 – 29, June 3 – 15, June 18 –
154 22, June 24 – 26. Figs. 3 and 4 illustrate Saharan dust transport by showing modeled dust
155 distribution; wind vectors, and satellite images with dust. The 3000 m altitude (~ 700
156 hPa) was chosen for displaying wind vectors because the average transportation of the
157 dust takes place above the humid air of the PBL, between 600–800 hPa (Carlson and
158 Prospero, 1972, Hamonou et al., 1999, Barkan et al., 2005, Alpert et al., 2004). Three of
159 the dust events showed a similar situation with dust close to the surface. However, it
160 should be noted that the dust event on June 3 – 15, 2007, was mainly characterized by

161 dust distant from the surface, which made an insignificant contribution to PM10
162 concentration (Fig. 3).

163

164 The synoptic situation during the aforementioned dust events was characterized as
165 follows:

166 May 20 – 29, 2007 - in this phenomenon, Saharan dust was transported to Italy
167 simultaneously along two different routes: (a) from the Eastern Sahara (Fig. 3a), and (b)
168 from the Western Sahara (Fig. 3 b). The synoptic situation was characterized by a specific
169 structure consisting of two intensive low-pressure systems, one over the Eastern
170 Mediterranean and the other over the Eastern Atlantic. A strong anticlockwise airflow,
171 associated with the low over the Eastern Mediterranean, transported dust in an
172 anticlockwise movement from the Eastern Sahara, around the Eastern Mediterranean, to
173 Northern Italy (Fig. 3b). Simultaneously, the other low over the Eastern Atlantic
174 produced favorable conditions for the development of a heavy dust storm over the
175 Western Mediterranean, accompanied by dust intrusions into Northern Italy (Fig. 3b).
176 Satellite images with dust over sea areas adjacent to Northern Italy supported model
177 simulations (Fig. 3c).

178

179 June 3 – 15, 2007 – the long dust presence over Northern Italy was due to a sequence of
180 dust events. From June 3 – 4, 2007, an intensive low-pressure system over the Central
181 Mediterranean created favorable conditions for dust transport from dust sources in the

182 Libyan Desert to Italy (Fig. 3d). As this low-pressure system shifted eastward during June
183 5 – 9, 2007, dust from the eastern part of the Sahara was transported to Northern Italy.

184 A low-pressure system over the tropical east Atlantic, generating south-west winds in its
185 eastern flank, could cause dust transport to Northern Italy from June 10 – 15, 2007. As
186 mentioned above, in accordance with DREAM 3-D dust distributions over Northern Italy,
187 this dust event was mainly characterized by dust distant from the surface, which made an
188 insignificant contribution to PM10 concentration. This is illustrated in Fig. 3e presenting
189 dust distribution within the vertical cross-section along 10.2° longitude. In the second
190 dust event, dust was not seen from space because of cloud cover over Northern Italy and
191 adjacent sea areas (Fig. 3f). Nevertheless, for the sake of consistency, we presented the
192 satellite image for the second dust event, as we did for all other events.

193

194 June 18 – 22, 2007 - an intensive Atlantic cyclone was the main causal factor for a severe
195 dust storm over the Western Mediterranean, accompanied by a direct dust transport along
196 the route from North Africa, through Italy, into the Po Valley (Fig. 4a). The modeled dust
197 transport was supported by satellite images with dust (Fig. 4b).

198

199 June 24 – 25, 2007 - as a consequence of the aforementioned Atlantic low, an extensive
200 high-pressure system arose over the whole Mediterranean Sea. As a result, a dust storm
201 took place over the Mediterranean, accompanied by dust intrusion into Northern Italy
202 (Fig. 4 c and d).

203 **3.2 Impact on TCAM performances**

204 As mentioned, the impact of DREAM dust boundary conditions on TCAM performance
205 over Northern Italy was assessed by comparing two model configurations, one with and
206 one without DREAM boundary conditions. The TCAM configuration without DREAM
207 boundary conditions is called hereafter the base TCAM model. Both TCAM model
208 configurations were run over the model domain (Fig. 2) with 10 km × 10 km horizontal
209 resolution and 11 vertical levels, from 10m to 4000 m above ground level. Following
210 that, the DREAM dust boundary condition impact on TCAM performance was assessed
211 in terms of correlation and mean normalized model absolute error, by comparing modeled
212 PM10 concentrations with PM10 measurements for the period under investigation.

213

214 230 PM10 monitoring sites, scattered all around the model domain (Fig. 2, crosses), were
215 used in order to comprehensively capture the distribution of particulate emissions
216 affecting the domain. At different PM10 sites, PM10 concentrations, averaged over the
217 period under consideration, varied between 10 and 50 $\mu\text{g}/\text{m}^3$. The sites with the lowest
218 average PM10 concentration are remote from both industrial centers and highly-
219 populated areas. These remote PM10 sites show the minimal level of air pollution in
220 Northern Italy.

221

222 **3.2.1. TCAM-DREAM performance over remote PM10 sites**

223 It is reasonable to suggest that these remote PM10 sites are most sensitive to Saharan dust
224 intrusions into Northern Italy, because of the absence of intensive industrial pollution.

225 The dust impact on PM10 concentration at these sites should be noticeable on the
226 background of low PM10 concentrations. Consequently, first we analyzed the Saharan
227 dust impact on TCAM performance for the eleven PM10 sites, which were characterized
228 by the lowest average PM10 concentrations ($PM_{10} \leq 14 \mu\text{g}/\text{m}^3$) over the period under
229 consideration

230

231 Fig. 5 represents time series of measured and modeled PM10 concentration averaged over
232 the aforementioned remote sites. It is seen that the observed PM10 concentrations (stars)
233 show two pronounced maxima during the two dust events May 20 – 29 and June 18 – 22,
234 2007. The PM10 concentrations in these maxima are approximately two times higher
235 than the background PM10 level. At these remote sites, such a strong deviation from the
236 background level can not be attributed to anthropogenic aerosol emissions, because of
237 their distance from anthropogenic sources. Moreover, the base TCAM model was not
238 capable of reproducing the two observed PM10 maxima (Fig. 5), because the dust effect
239 was not included. The two aforementioned facts provided us with evidence of dust
240 aerosol near the surface in Northern Italy. Due to the use of improved boundary
241 conditions, the integrated TCAM-DREAM model more accurately reproduced PM10
242 measurements than the base TCAM model. For the two other dust events on June 3 – 15
243 and June 24 – 26, 2007, modeled PM10 concentrations were close to the background
244 PM10 concentrations, in accordance with PM10 measurements.

245

246 **3.2.2. TCAM-DREAM performance over 230 PM10 sites in Northern Italy**

247 The time series of measured PM10 concentration, averaged over all available 230 PM10
248 monitoring sites, were obtained for each day from May 15 to June 30, 2007. For the 230
249 PM10 sites distributed over the model domain, model-vs.-measurement comparisons for
250 both model configurations showed that the use of DREAM boundary conditions led
251 TCAM to better performance (Fig. 6). In particular, the integrated TCAM-DREAM
252 model was able to capture high peaks in measured PM10 concentration during the two
253 dust events, May 20 – 29, and June 18 – 22, 2007. Note that, although the integrated
254 TCAM-DREAM model more accurately reproduced PM10 measurements than the base
255 TCAM model, sometimes there was some discrepancy between modeled concentrations
256 and measurements. Specifically, during the period under consideration, TCAM-DREAM
257 modeled concentrations were somewhat lower than PM10 measurements. In addition,
258 there was a three-day delay between the measured and modeled PM10 maxima during the
259 first dust event on May 21 – 29, 2007 (Fig. 6). A possible reason for the delay could be
260 errors in the meteorological parameters.

261

262 The better capability of the integrated TCAM-DREAM model over the base TCAM
263 model is illustrated by the better correspondence of 95th percentiles between observed and
264 modeled PM10 concentration (Fig. 7). To further support this point, the DREAM dust
265 boundary condition impact on TCAM performance was assessed in terms of correlation
266 (r) and normalized mean absolute error (nmae). Fig. 8 represents box-plots of the
267 correlation coefficient (Fig. 8a) and the normalized mean absolute error (Fig. 8b),
268 computed on the basis of model-vs.-measurement comparisons at 230 PM10 monitoring

269 sites. The boxes show a significant increase in the r median from 0.40 to 0.65 achieved by
270 using DREAM dust boundary conditions (Fig. 8a). The $nmae$ median dropped from 0.5
271 for the base simulation to 0.4 for the simulation with DREAM boundary conditions (Fig.
272 8b). This result indicates the improvement in model performance due to the use of
273 DREAM boundary conditions.

274

275 In order to demonstrate how the improvement in model performance covers the model
276 domain, the difference in correlation between measured and modeled PM10
277 concentrations, based on the two model configurations, was analyzed (Fig. 9a). The
278 improvement was characterized by a positive correlation difference higher than 0.1. The
279 map of the correlation differences highlights the fact that the noticeable improvement in
280 TCAM model performance was observed almost over the entire model domain. We used
281 a map of $nmae$ differences in order to demonstrate how the decrease in normalized mean
282 absolute error covers the model domain. The improvement was characterized by a
283 negative $nmae$ difference less than -0.1. Shown in Fig. 9b, the map of $nmae$ differences
284 demonstrates a noticeable improvement in the integrated TCAM-DREAM model
285 performance almost over the entire model domain, achieved by the use of DREAM
286 boundary conditions.

287

288 **Conclusions**

289 To improve air pollution modeling, 3D dust fields, produced by means of the DREAM
290 dust forecasts, were integrated as boundary conditions into the TCAM chemical transport

291 model. This allowed us to reproduce Saharan dust contribution to PM10 concentration
292 over Northern Italy for the period May 15 – June 30, 2007, when four significant dust
293 events were observed.

294

295 First, Saharan dust impact on TCAM performance was analyzed at eleven remote PM10
296 sites which had the lowest level of air pollution ($PM_{10} \leq 14 \mu\text{g}/\text{m}^3$) over the period under
297 consideration. For those remote sites, the observed high PM10 concentrations during dust
298 events stood prominently on the background of low PM10 concentrations. At the remote
299 sites, such a strong deviation from the background level can not be attributed to
300 anthropogenic aerosol emissions because of their distance from anthropogenic sources.
301 The observed maxima in PM10 concentration during dust events is evidence of dust
302 aerosol near the surface in Northern Italy. During all dust events under consideration, the
303 integrated TCAM-DREAM model produced more accurate PM10 concentrations than the
304 base TCAM model.

305

306 A comparison between modeled PM10 concentrations and PM10 measurements during
307 the period under study May 15 – June 30, 2007, was carried out at 230 PM10 monitoring
308 sites, distributed within the model domain. This model-vs.-measurement comparison
309 showed that the integrated TCAM –DREAM model more accurately reproduced PM10
310 concentrations than the base TCAM model, both in term of correlation and mean error.
311 Specifically, as estimated for 230 PM10 monitoring sites, the correlation median
312 increased from 0.40 to 0.65, while the normalized mean absolute error median dropped

313 from 0.5 to 0.4. Noticeable improvement in TCAM model performance was obtained
314 almost over the entire model domain.

315

316 The obtained results indicate that dust can contribute significantly to PM10
317 concentrations in Northern Italy. This dust contribution should be taken into account
318 when estimating the exceedance of pollution limits. Our results are of importance to
319 countries which have to pay a penalty for exceeding the pollution limit. By subtracting
320 dust contribution from PM10 measurements, these countries could show lower rates of
321 man-made pollution.

322

323 **Acknowledgments**

324 The work has been partially supported by COST ES0602 action and MIUR (Ministero
325 Italiano Università e Ricerca). The authors acknowledge Dr. G. Curci (CETEMPS, Italy)
326 for the Chimere boundary conditions.

327 **References**

328 Alpert P., Kishcha P, Shtivelman A, Krichak S, Joseph J. Vertical distribution
329 of Saharan dust based on 2.5-year model predictions. Atmospheric
330 Research 2004;70: 109–130.
331 Amiridis V, Kafatos M, Perez C, Kazadzis S, Gerasopoulos E, Mamouri RE,
332 et al. The potential of the synergistic use of passive and active remote
333 sensing measurements for the validation of a regional dust model. Annales
334 Geophysicae 2009;27:3155-3164.

335 Barkan J, Alpert P, Kutiel H, Kishcha P. Synoptics of dust transportation days
336 from Africa toward Italy and central Europe, *J. Geophys. Res.* 2005;110,
337 D07208, doi:10.1029/2004JD005222.

338 Carlson TN, Prospero JM. The large-scale movement of Saharan air
339 outbreaks over the northern equatorial Atlantic, *Journal of Applied*
340 *Meteorology* 1972;11: 283–297.

341 Carnevale C, Gabusi V, Volta M. POEM-PM: an emission model for
342 secondary pollution control scenarios, *Environmental Modelling and*
343 *Software*, 2006;21:320-329.

344 Carnevale C, Decanini E, Volta M. Design and validation of a multiphase 3D
345 model to simulate tropospheric pollution, *Science of the Total*
346 *Environment*, 2008;390:166-176.

347 C. Carnevale, E. Pisoni and M. Volta. A non-linear analysis to detect the
348 origin of PM10 concentrations in Northern Italy, *Science of the Total*
349 *Environment*, 2010;409:182-191.

350 Chock DP, Winkler SL, Sun P. Comparison of Stiff Chemistry Solvers for
351 Air Quality Modeling. *Environ. Sci. Technol.*, 1994;28:1882–1892.

352 Cuvelier C, Thunis P, Vautard R, Amann M, Bessagnet B, Bedogni M, et al.
353 CityDelta: A model intercomparison study to explore the impact of
354 emission reductions in European cities in 2010. *Atmospheric Environment*,
355 2007;41:189-207.

356 Di Nicolantonio W, Cacciari A, Petritoli A, Carnevale C, Pisoni E, Volta M
357 et al. MODIS and OMI satellite observations supporting air quality
358 monitoring. *Radiation Protection Dosimetry*, 2009;137:280-287.

359 Ecudero M, Querol X, Avila A, Cuevas E. Origin of the exceedances of the
360 European daily PM limit value in regional background areas of Spain.
361 *Atmospheric Environment*, 2007;41:730-744.

362 Forester C. Higher order monotonic convection difference schemes. Journal
363 of Computational Physics, 1977;23:1–22.

364 Fountoukis C, Nenes A. ISORROPIA II: A Computationally Efficient
365 Aerosol Thermodynamic Equilibrium Model for K^+ , Ca^{2+} , Mg^{2+} , NH_4^+ ,
366 Na^+ , SO_4^{2-} , NO_3^- , Cl^- , H_2O Aerosols, Atmospheric Chemistry and
367 Physic, 2007;7:4639–4659.

368 Giorgi, F. A particle dry-deposition parameterization scheme for use in tracer
369 transport models. J. Geophys. Res., 1986;91:9794–9806.

370 Gobbi GP, Barnaba F, Ammannato L. Estimating the impact of Saharan dust
371 on the year 2001 PM10 record of Rome, Italy. Atmospheric Environment,
372 2007;41:261-275.

373 Hamonou E, Chazette P, Bali D, Dulac F, Schneider X, Galani F, Ancellet G,
374 Papayanis A. Characterization of the vertical structure of the Saharan dust
375 transport to the Mediterranean basin. J. Geophys. Res. 1999;104:256– 270.

376 Hindmarsh A. LSODE and LSODEI, two new initial value ordinary
377 differential equation solvers, ACM-SIGNAL Newsletter, 1980;15:10–11.

378 Kishcha P, Barnaba F, Gobbi GP, Alpert P, Shtivelman A, Krichak S et al.
379 Vertical distribution of Saharan dust over Rome (Italy): comparison
380 between 3-year model predictions and lidar soundings. J. Geophys. Res.,
381 2005;110.

382 Kishcha P, Alpert P, Shtivelman A, Krichak S, Joseph J, Kallos G et al.
383 Forecast errors in dust vertical distributions over Rome (Italy): Multiple
384 particle size representation and cloud contributions. J. Geophys. Res.,
385 2007;112.

386 Kishcha P, Nickovic S, Ganor E, Kordova L, Alpert P. Saharan dust over the
387 Eastern Mediterranean: Model sensitivity. Air pollution modelling and its
388 applications XIX. 2008;4:358-366.

389 Marchuk G. *Methods of Numerical Mathematics*, Springer-Verlag, New
390 York; 1975.

391 Masson O, Piga D, Guarriaran R, D'Amico D. Impact of exceptional Saharan
392 dust outbreak in France: PM10 and artificial radionuclides concentrations
393 in air in dust deposit. *Atmospheric Environmnet*, 2010;44: 2478-2486.

394 Nenes A, Pandis S, Pilinis C. ISORROPIA: A new thermodynamic
395 equilibrium model for multiphase multicomponent inorganic aerosols.
396 *Aquatic Geochemistry*, 1998;4:123-152.

397 Nickovic S, Kallos G, Papadopoulos A, Kakaliagou O. A model for
398 prediction of desert dust cycle in the atmosphere. *J. Geophys. Res.*,
399 2001;106:18113-18129.

400 Papanastasiou DK, Poupkou A, Katragkou E, Amiridis V, Melas D,
401 Mihalopoulos N et al. An assessment of the efficiency of dust regional
402 modelling to predict Saharan dust transport episodes. *Advances in*
403 *Meteorology*, 2010;doi:10.1155/2010/154368.

404 Papayannis A, Zhang H, Amiridis V, Ju H, Chourdakis G, Georgoussis G et
405 al. Extraordinary dust event over Beijing, China, during April 2006: Lidar,
406 Sun photometric, satellite observations and model validation. *Geophys.*
407 *Res. Let.*, 2007;34:doi:10.1029/2006GL029125.

408 Pederzoli A, Mircea M, Finardi S, Di Sarra A, Zanini G. Quantification of
409 Saharan dust contribution to PM10 concentrations over Italy during 2003 –
410 2005. *Atmospheric Environment*, 2010;44:4181-4190.

411 Pepper D, Kern C, Long P. Modelling the dispersion of atmospheric pollution
412 using cubic splines and chapeau functions, *Atmospheric Environment*,
413 1979;13:223–237.

414 Pérez C, Nickovic S, Pejanovic G, Baldasano JM, Özsoy E. A long Saharan
415 dust event over the western Mediterranean: Lidar, sun photometer

416 observations and regional dust modeling. *J. Geophys. Res.*, 2006;111:
417 D15214, doi:10.1029/2005JD006579.

418 Perrino C, Canepari S, Cardarelli E, Catrambone M, Sagolini T. Inorganic
419 constituents of urban air pollution in the Lazio region (Central Italy).
420 *Environmental Monitoring and Assessment*, 2008;136:69-86.

421 Rodriguez S, Querol X, Alastuey A, Kallos, G, Kakaliagou O. Saharan dust
422 contributions to PM10 and TSP levels in Southern and Eastern Spain,
423 *Atmospheric Environment*, 2001;35:2433-2447.

424 Schmidt H, Derognat C, Vautard R, Beekman M. A comparison of simulated
425 and observed ozone mixing ratios for the summer of 1998 in Western
426 Europe. *Atmospheric Environment*, 2001;35:6277–6297.

427 Volta M, Finzi G. GAMES, a comprehensive gas aerosol modelling
428 evaluation system, *Environmental Modelling and Software*, 2006;21;587-
429 594.

430 **Figure captions**

431

432 Fig. 1. The TCAM model configuration with embedded DREAM dust forecasts.

433

434 Fig. 2. The Northern Italy model domain with PM10 monitoring sites (crosses). The red
435 rings designate the remote PM10 sites with the lowest average PM10 concentration.

436

437 Fig. 3. DREAM dust forecasts and SeaWiFS satellite images of Saharan dust transport to
438 Northern Italy in the first (top) and second (bottom) dust events. In Fig. 3 a, b, and d the
439 colors designate dust loading in g/m^2 , while in Fig. 3e the colors designate dust
440 concentration in $\mu g/m^3$ within the vertical cross-section along the 10.2o longitude. Wind
441 vectors designate wind at 3000 m altitude.

442

443 Fig. 4. DREAM dust forecasts (left panel) and SeaWiFS satellite images (right panel) of
444 Saharan dust transport to Northern Italy: (a and b) June 19, 2007, and (c and d) June 24,
445 2007. In the left panel, the colors designate dust loading in g/m^2 , while wind vectors
446 designate wind at 3000 m altitude.

447

448 Fig. 5. Time series of measured and modeled PM10 concentrations averaged over 11
449 remote PM10 monitoring sites. The blue underlines designate periods with the dust
450 events.

451

452 Fig. 6. Time series of measured PM10 concentration (stars) and modeled PM10
453 concentration simulated by TCAM with DREAM boundary conditions (solid line) and
454 without DREAM (dashed line). The data were averaged over 230 PM10 monitoring sites.
455 The blue underlines designate periods with the dust events.

456

457 Fig. 7. Comparison of 95th percentiles between measured and modeled PM10
458 concentration during the dust events, simulated by TCAM with DREAM boundary
459 conditions (TCAM+DREAM) and without DREAM (TCAM).

460

461 Fig. 8. Box plots of two indices of model performance: (a) the median of correlation, and
462 (b) the median of normalized mean absolute error, computed on the basis of model-vs.-
463 measurement comparisons at 230 PM10 monitoring sites. The performance of two model
464 configurations is compared: TCAM with DREAM boundary conditions
465 (TCAM+DREAM) and the base TCAM model (TCAM).

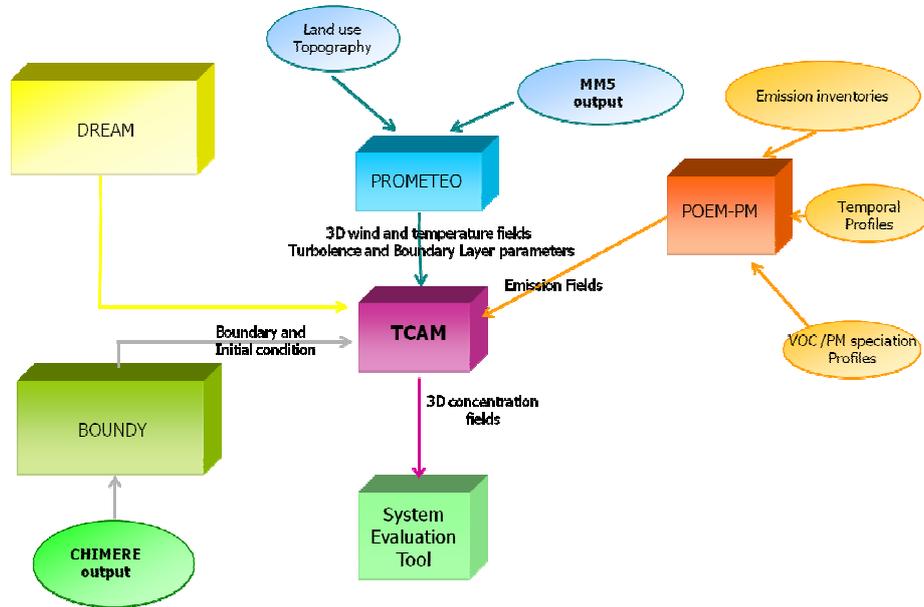
466

467 Fig. 9. Maps of differences in (a) correlation and in (b) normalized mean absolute error
468 based on two model simulations with and without the DREAM boundary conditions.
469 (blue: noticeably decreasing (< -0.1), red: noticeably increasing (> 0.1), green: almost
470 unchanged).

471

472

473



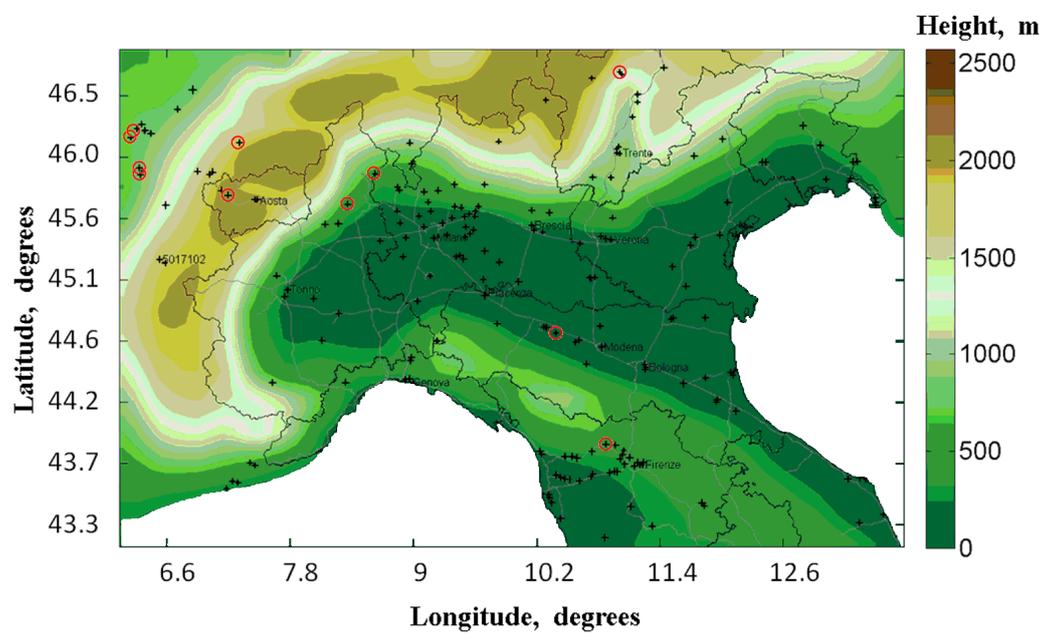
474

475

Fig. 1. The TCAM model configuration with embedded DREAM dust forecasts.

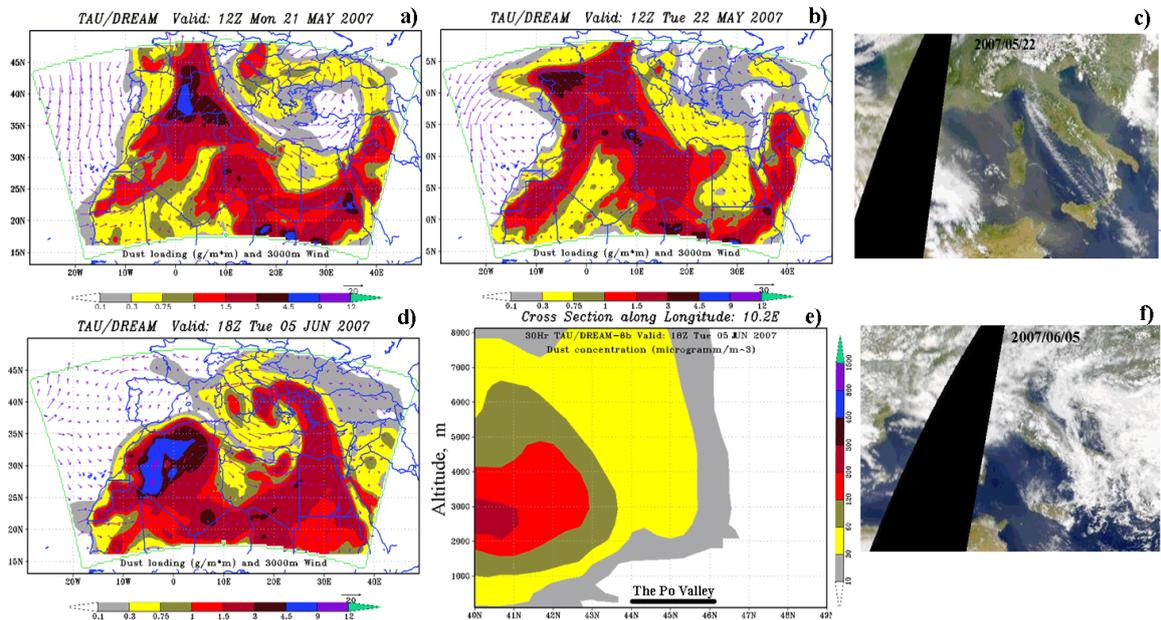
476

477



478

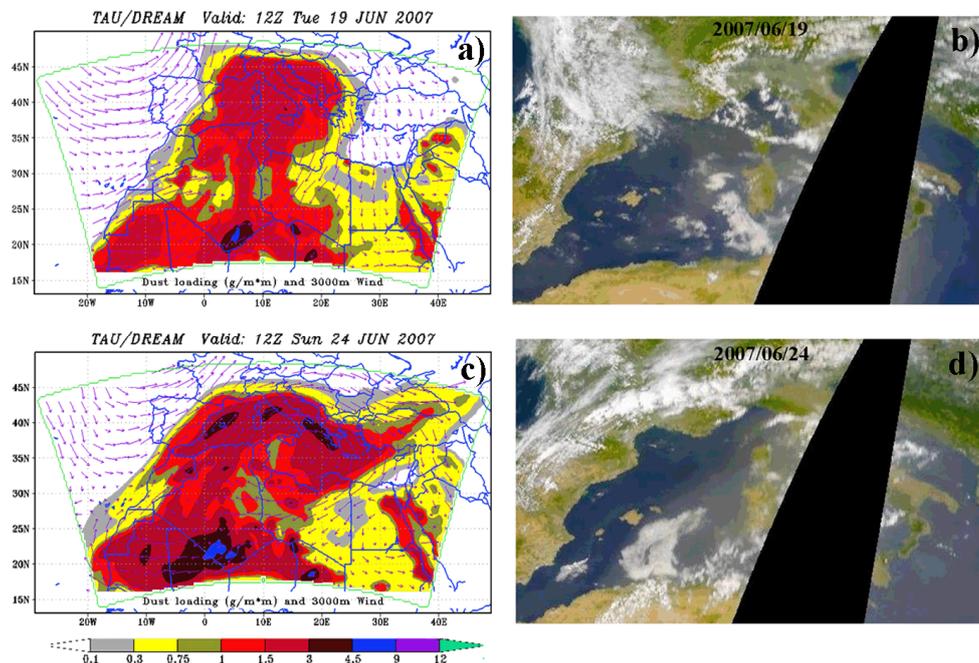
479 Fig. 2. The Northern Italy model domain with PM10 monitoring sites (crosses). The red
480 rings designate the remote PM10 sites with the lowest average PM10 concentration.



482

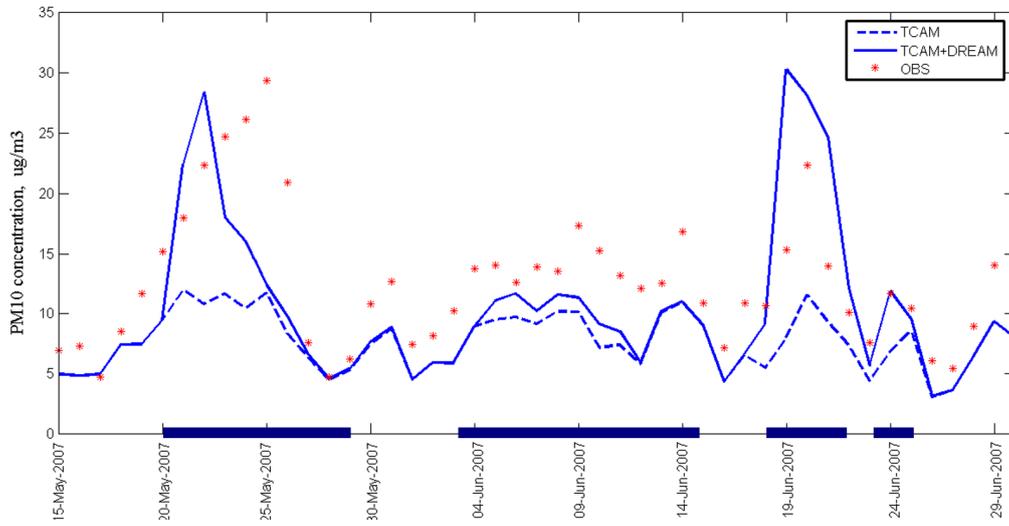
483 Fig. 3. DREAM dust forecasts and SeaWiFS satellite images of Saharan dust transport to
 484 Northern Italy in the first (top) and second (bottom) dust events. In Fig. 3 a, b, and d the
 485 colors designate dust loading in g/m^2 , while in Fig. 3e the colors designate dust
 486 concentration in $\mu\text{g}/\text{m}^3$ within the vertical cross-section along the 10.2° longitude. Wind
 487 vectors designate wind at 3000 m altitude.

488



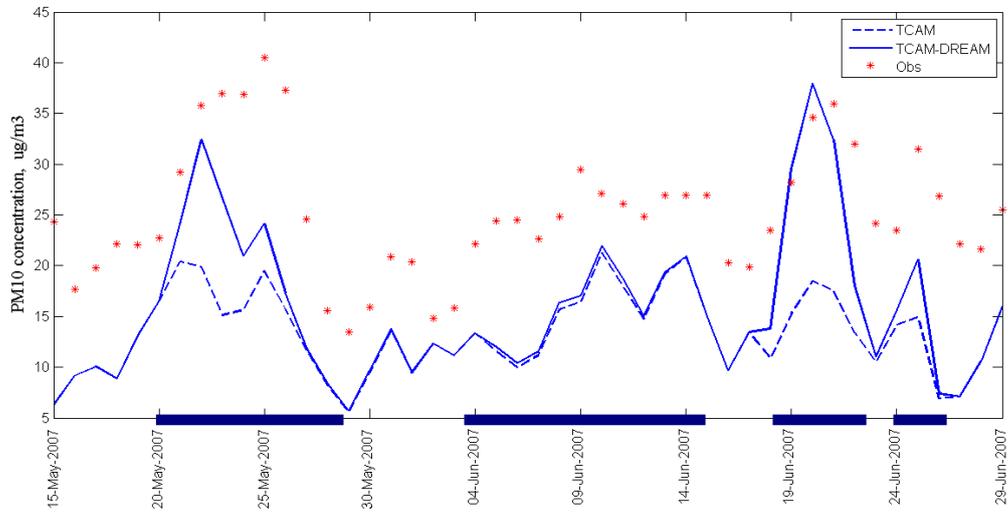
490 Fig. 4. DREAM dust forecasts (left panel) and SeaWIFS satellite images of Saharan dust
 491 transport to Northern Italy: (a and b) June 19, 2007, and (c and d) June 24, 2007. In the
 492 left panel, the colors designate dust loading in g/m^2 , while wind vectors designate wind at
 493 3000 m altitude.

494



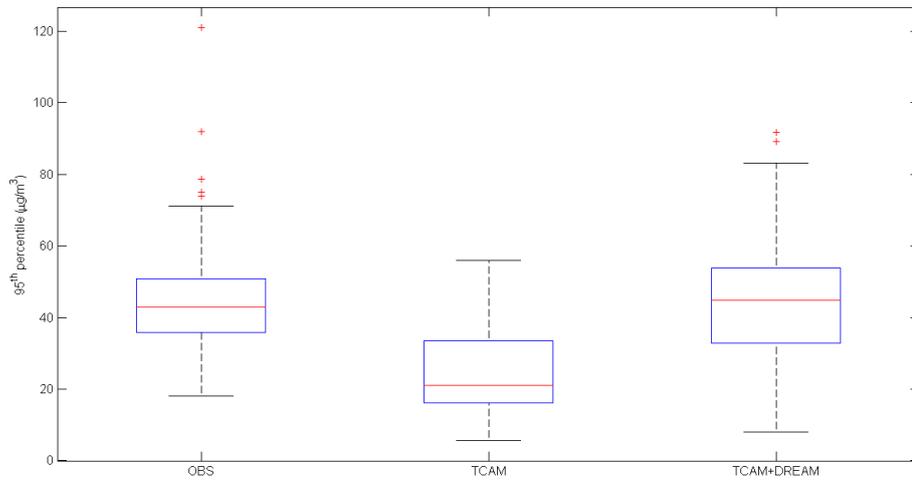
495

496 Fig. 5. Time series of measured and modeled PM10 concentrations averaged over 11
497 remote PM10 monitoring sites. The blue underlines designate periods with the dust
498 events.



500

501 Fig. 6. Time series of measured PM10 concentration (stars) and modeled PM10
 502 concentration simulated by TCAM with DREAM boundary conditions (solid line) and
 503 without DREAM (dashed line). The data were averaged over 230 PM10 monitoring sites.
 504 The blue underlines designate periods with the dust events.

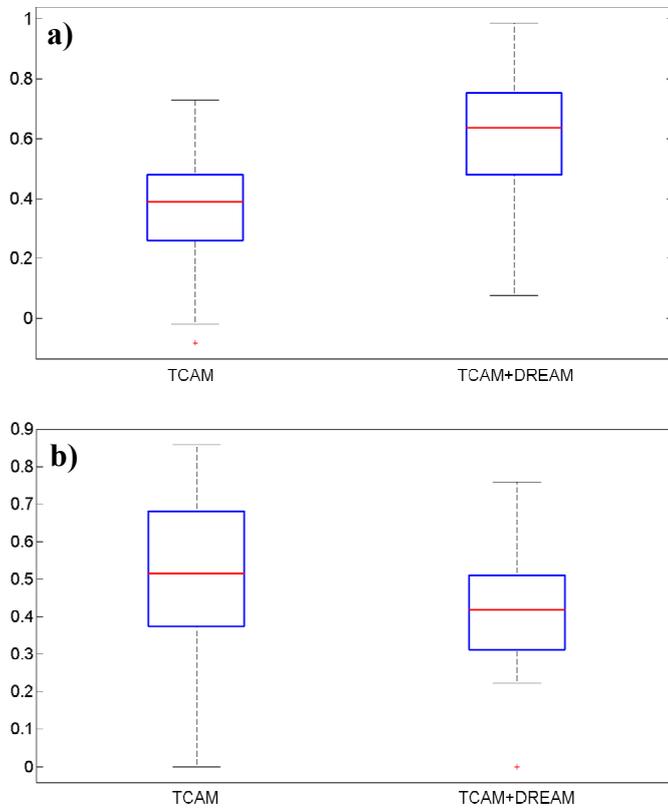


505

506 Fig. 7. Comparison of 95th percentiles between measured and modeled PM10
 507 concentration during the dust events, simulated by TCAM with DREAM boundary
 508 conditions (TCAM+DREAM) and without DREAM (TCAM).

509

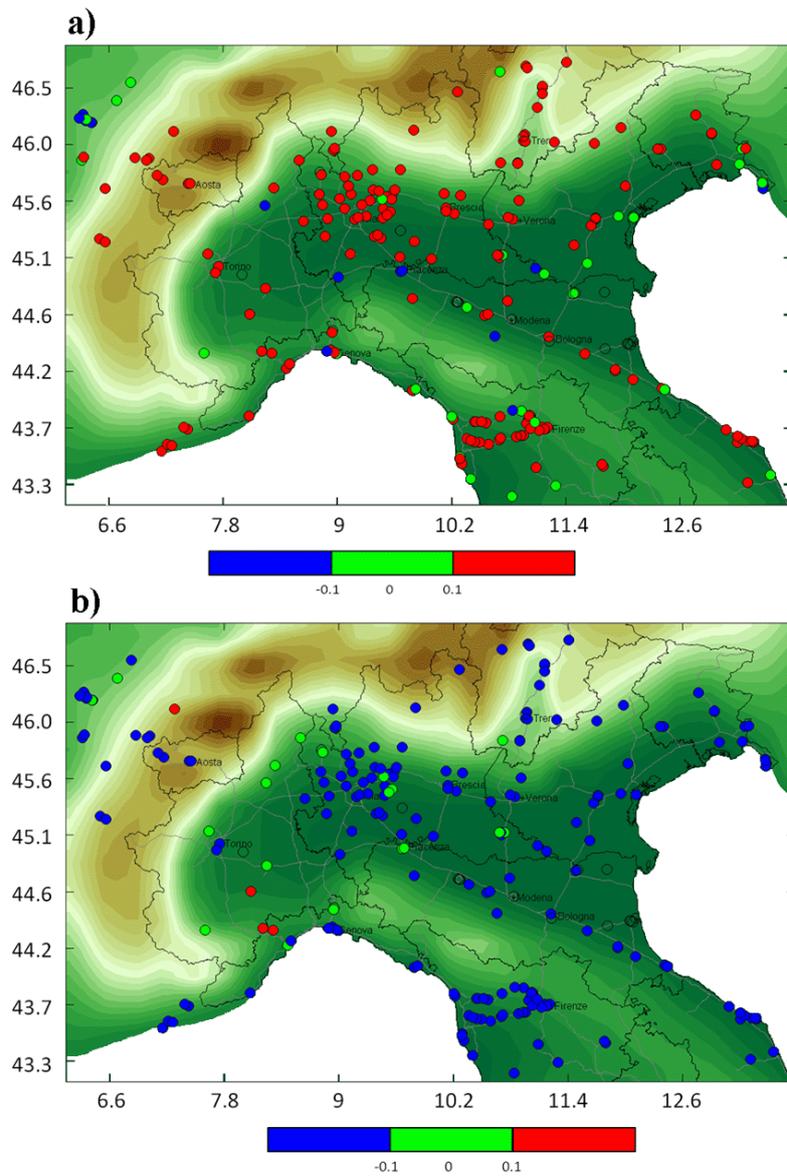
510



511

512 Fig. 8. Box plots of two indices of model performance: (a) the median of correlation, and
513 (b) the median of normalized mean absolute error, computed on the basis of model-vs.-
514 measurement comparisons at 230 PM10 monitoring sites. The performance of two model
515 configurations is compared: TCAM with DREAM boundary conditions
516 (TCAM+DREAM) and the base TCAM model (TCAM).

517



519 Fig. 9. Maps of differences in (a) correlation and in (b) normalized mean absolute error
 520 based on two model simulations with and without the DREAM boundary conditions.
 521 (blue: noticeably decreasing (< -0.1), red: noticeably increasing (> 0.1), green: almost
 522 unchanged).