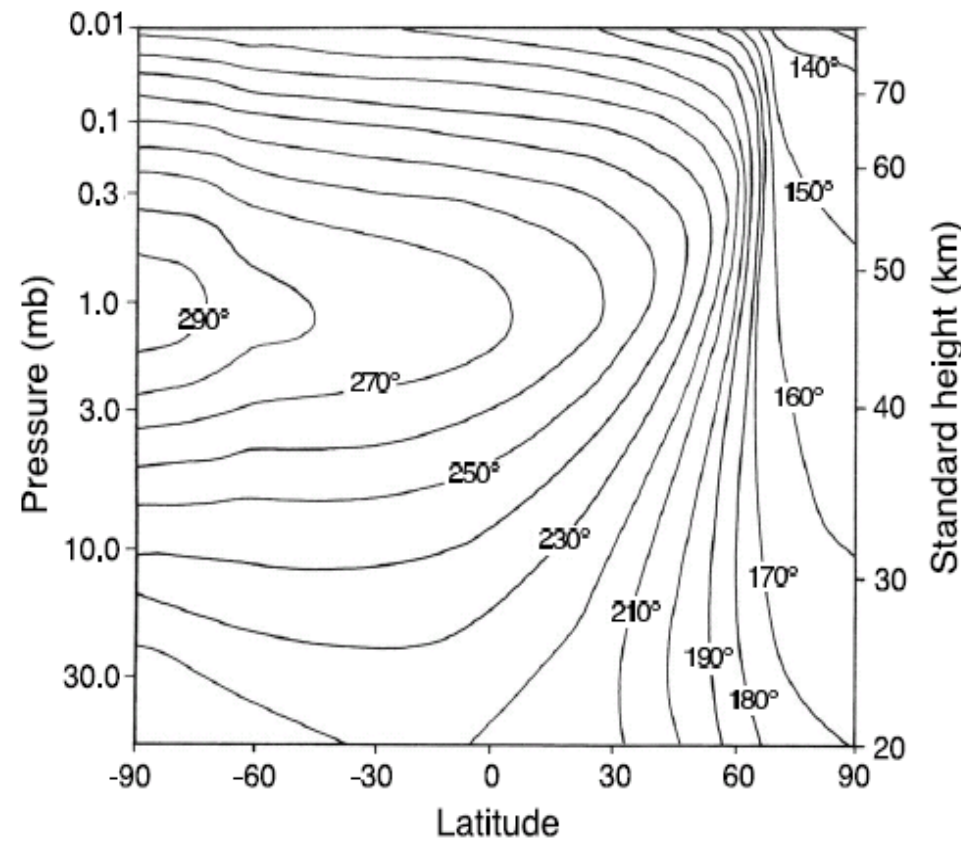
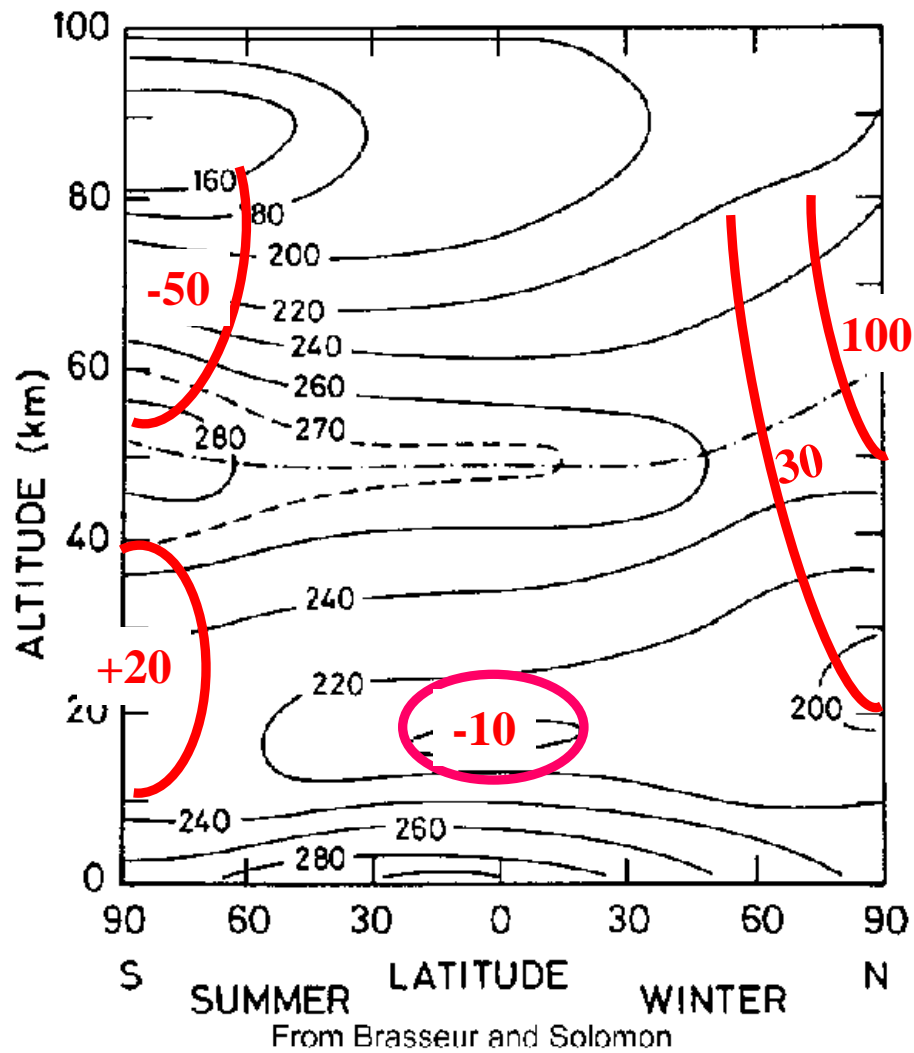


Fig. 13.12 The zonally averaged temperature and zonal wind in January. The temperature contour interval is 10 K, and values less than 220 K are shaded. Zonal wind contours are 10 m s^{-1} and negative (westward) values are shaded.⁷



From Vallis (2006) Adapted from Fels (1985), with the help of K. Hamilton.

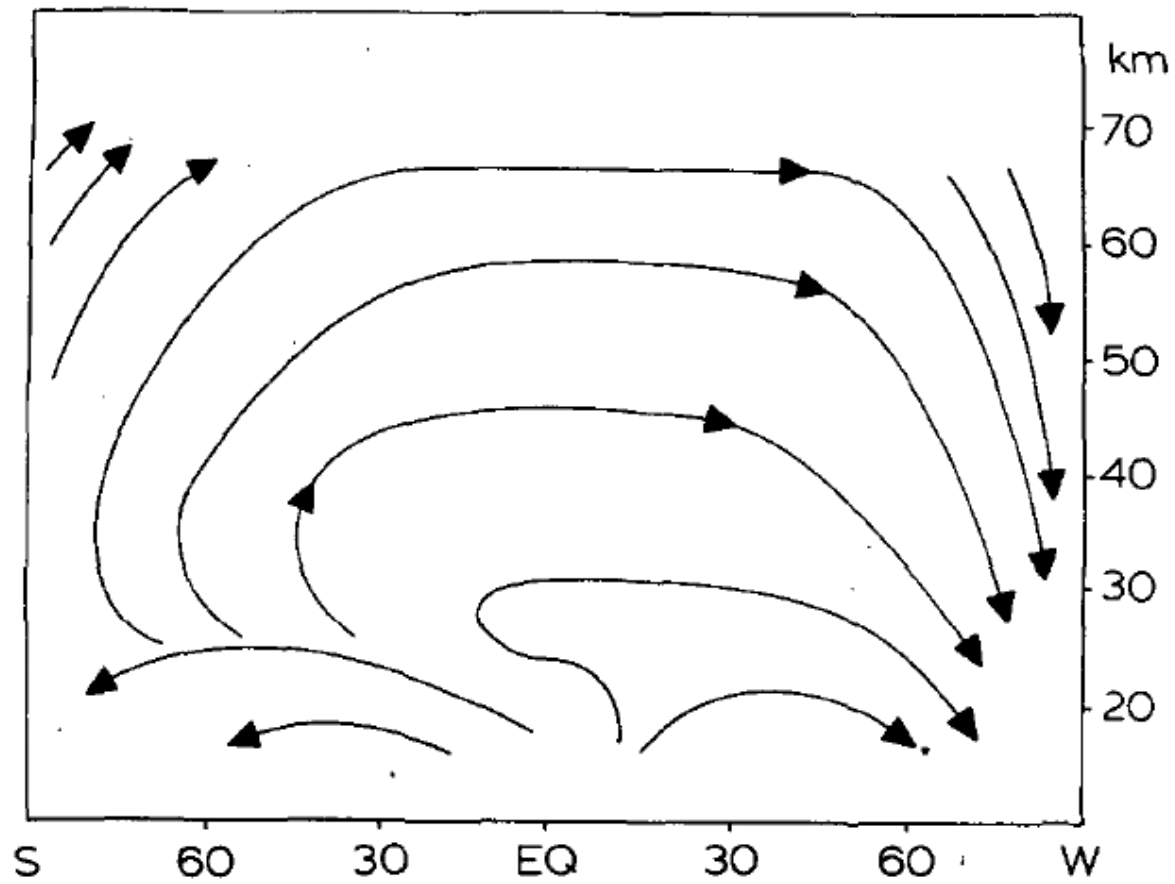
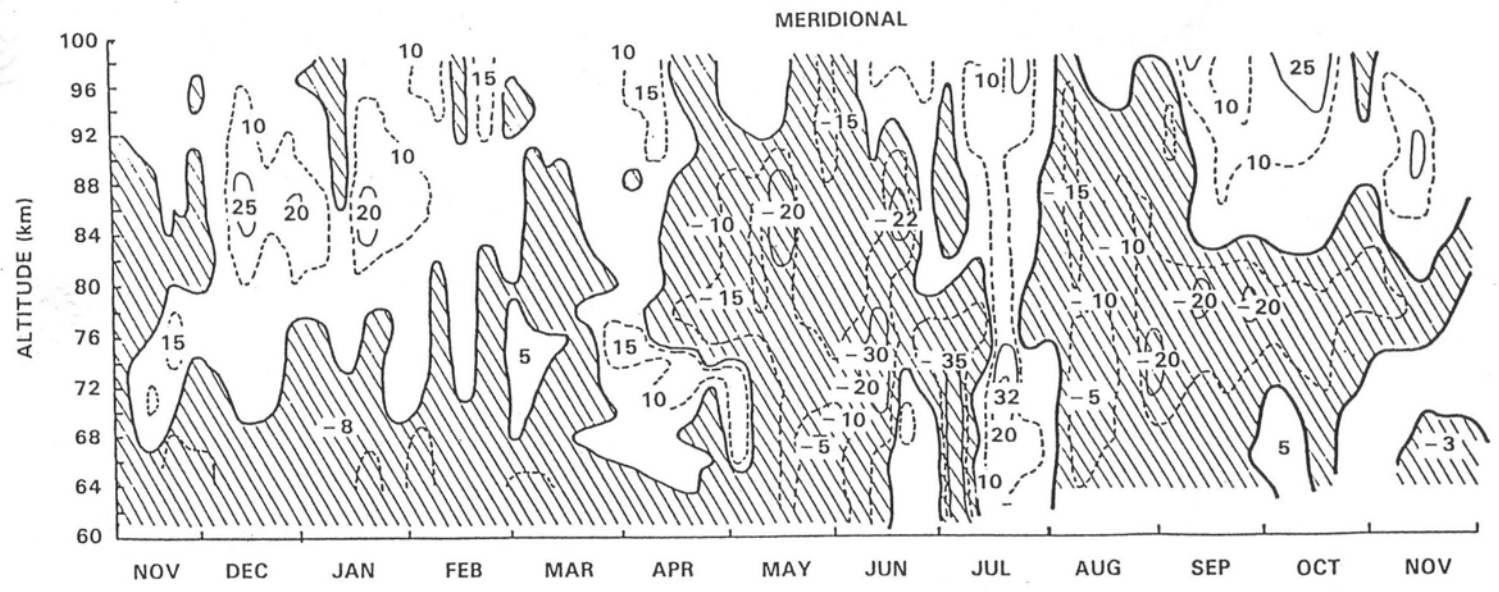
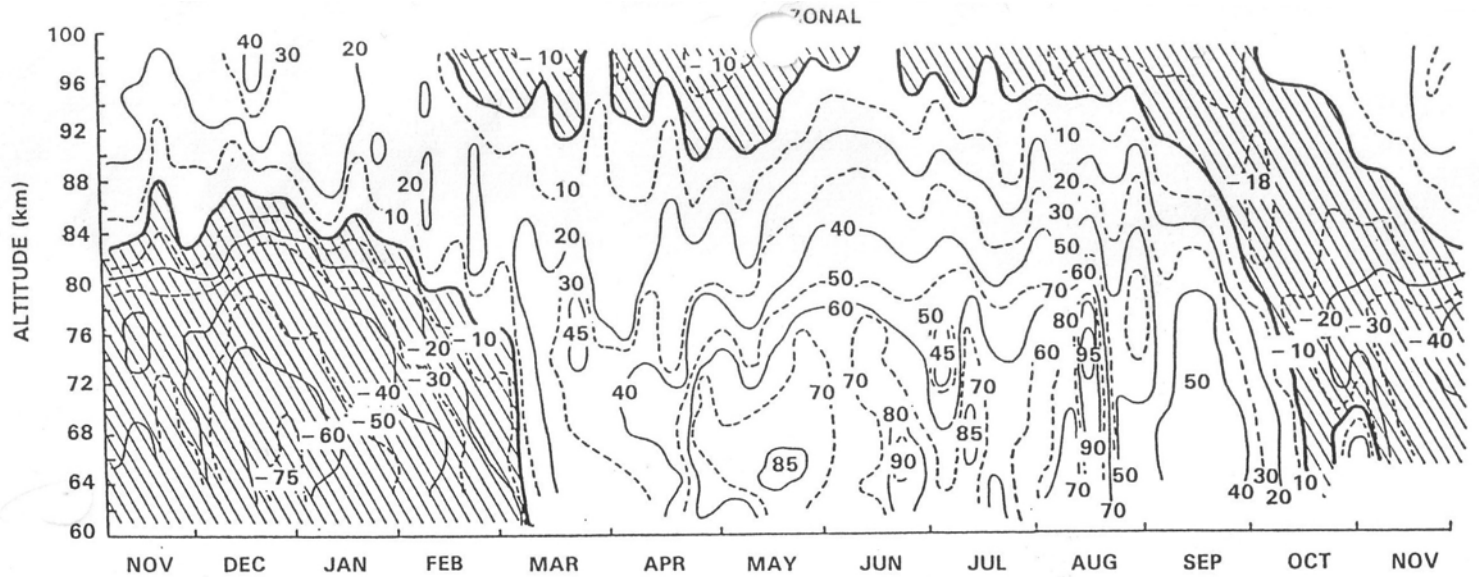


FIG. 4. Streamlines associated with Lagrangian-mean velocities.

Dunkerton (1978)



1984

Figure 8-3. Zonal and meridional winds (ms^{-1}) in the mesosphere measured with a PR radar at Adelaide (35°S , 128°E). Shaded areas indicate regions of westward and southward flow.

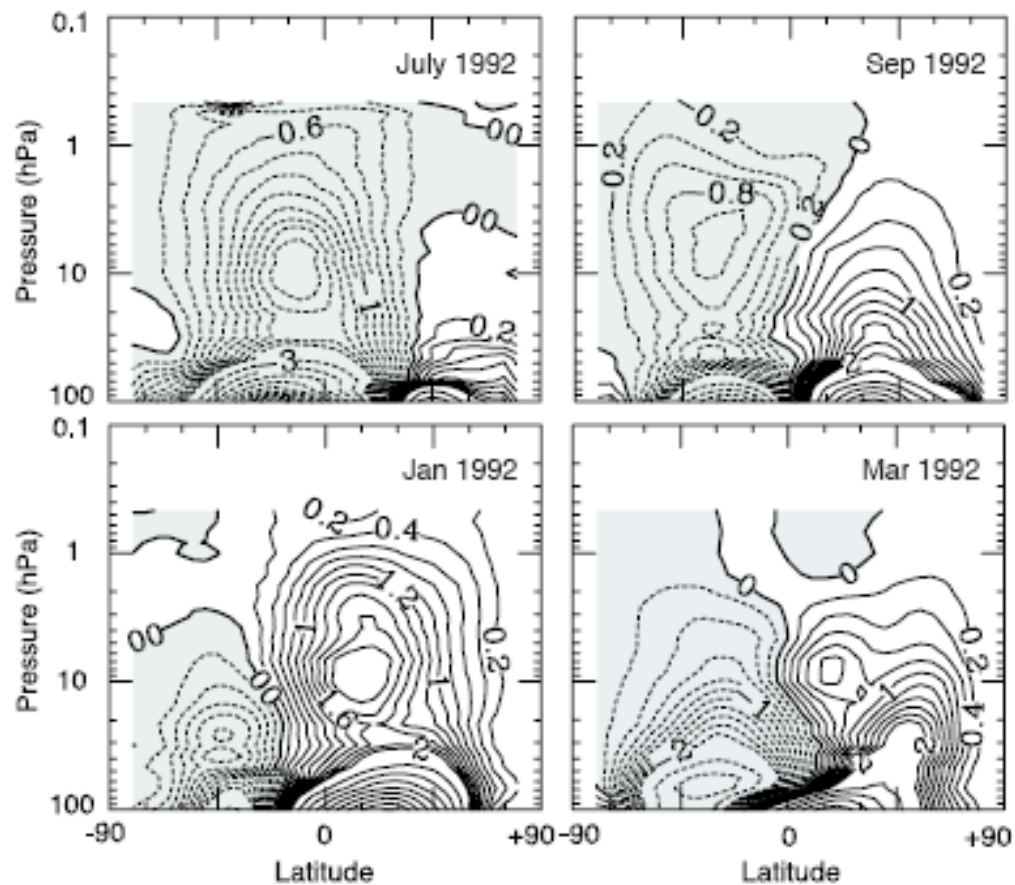
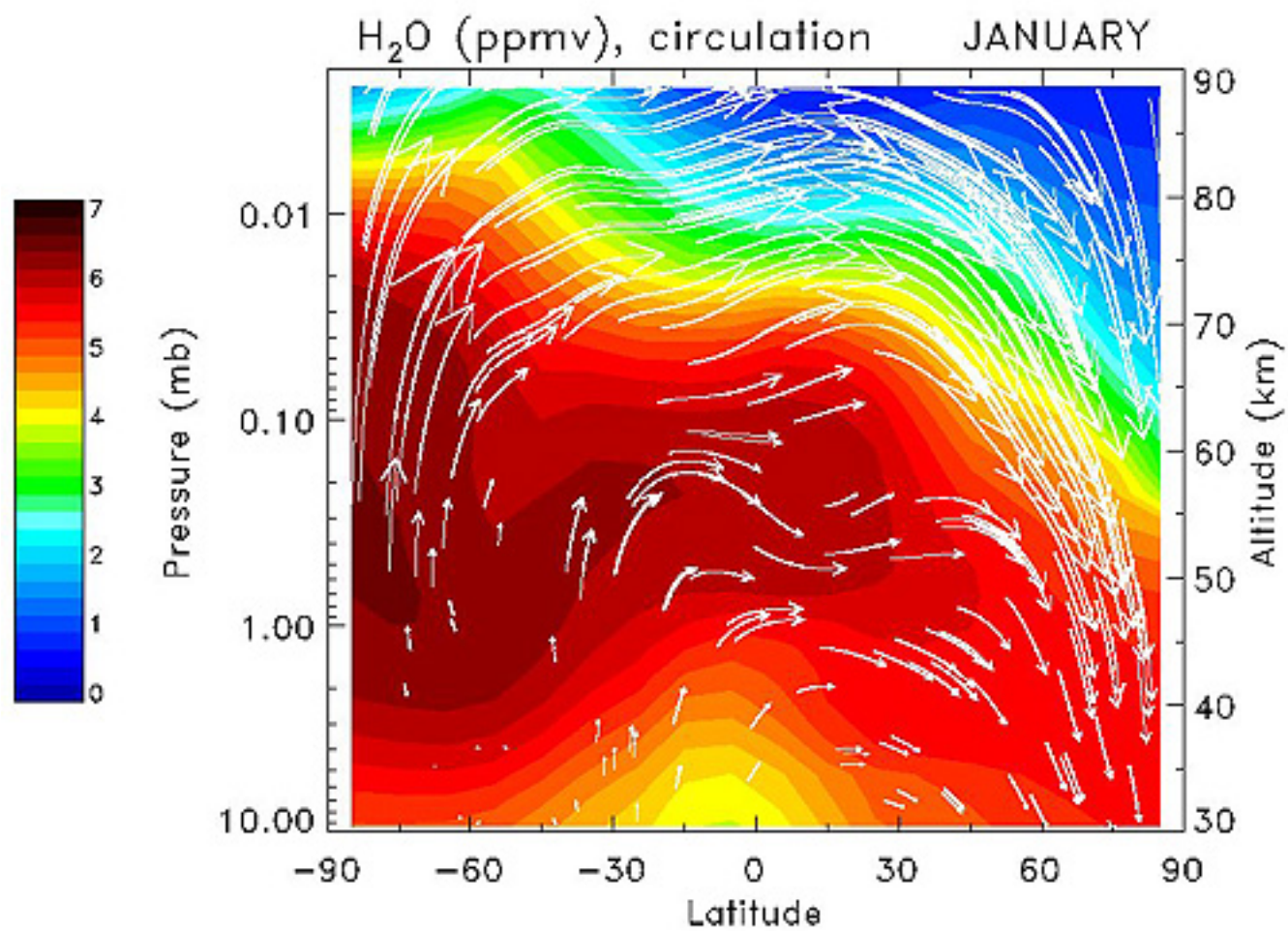
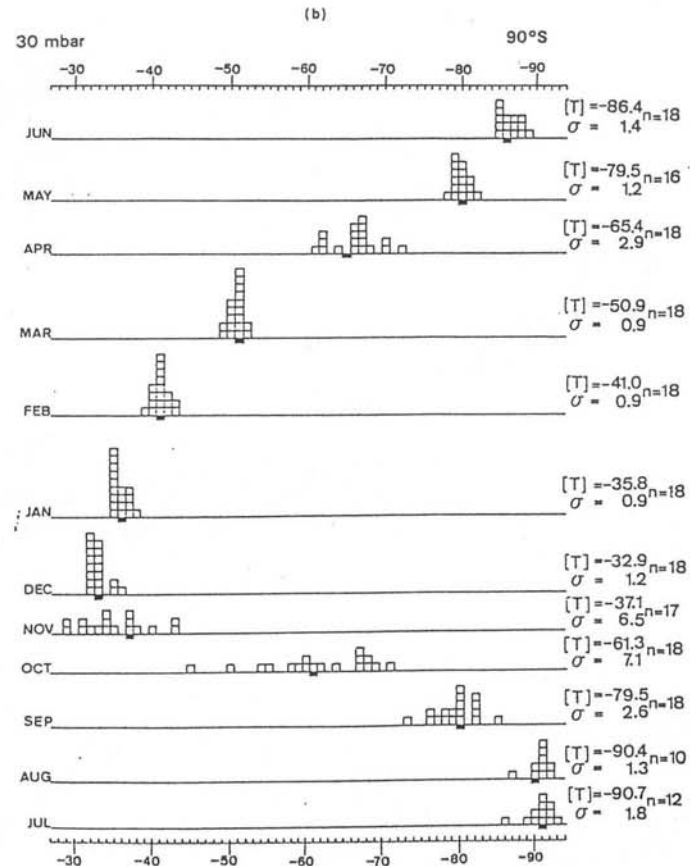
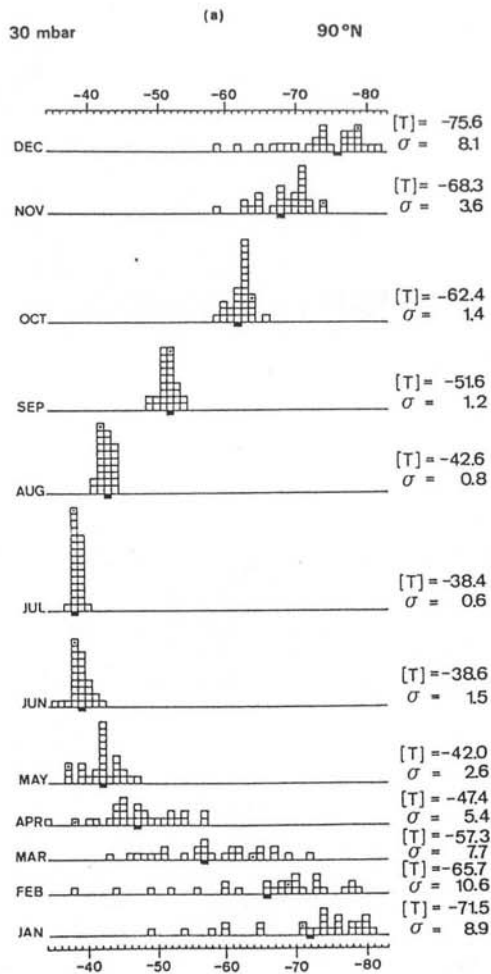


Fig. 13.13 The observed thickness-weighted (residual) streamfunction in the stratosphere, in Sverdrups (10^9 kg s^{-1}). The circulation is clockwise where the contours are solid. Note the stronger circulation in the winter hemispheres, whereas the equinoctial circulations (September, March) are more inter-hemispherically symmetric.⁸





1.7(a)
Figure 6-25. Frequency distribution of the monthly mean 30 mb temperatures (°C). The interval is 1 °C. The long-term average [T] is given on the right hand side of the picture together with the standard deviation σ , and [T] is also marked as a black box in the frequency distribution. (a) is for the North Pole using radiosonde data for the period July 1955-Dec. 1982,

Figure 6-25. (b) is for the South Pole for the period 1961-1978 using, for each month, data for the number of years specified by n. [After Naujokat, 1981, and Labitzke and Naujokat, 1983].

1.7(b)

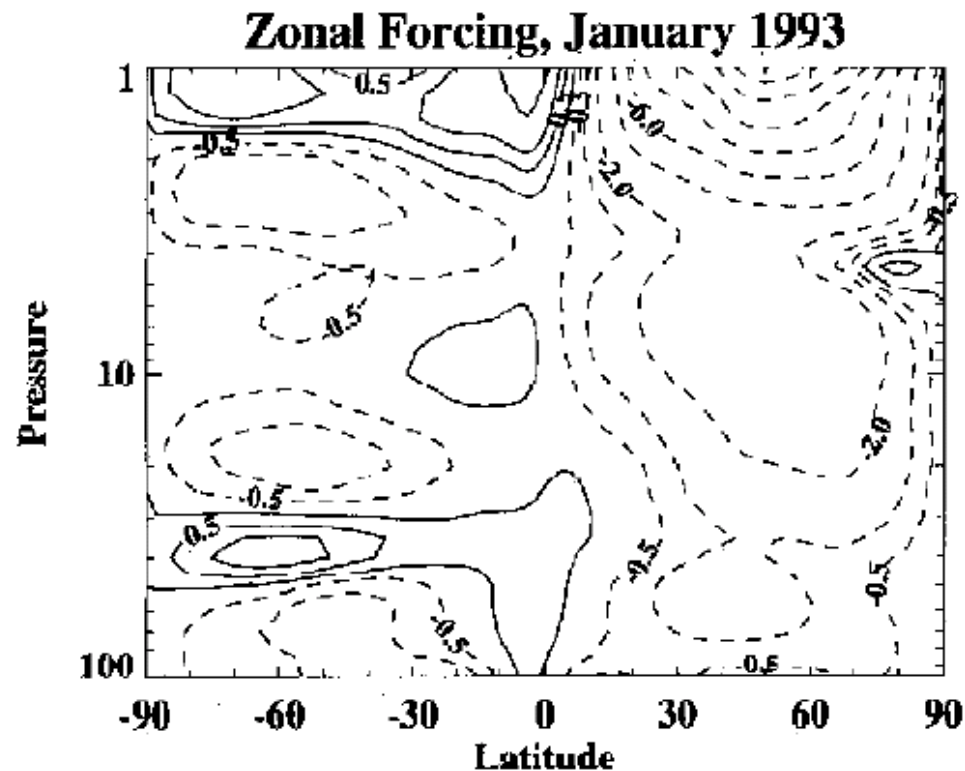
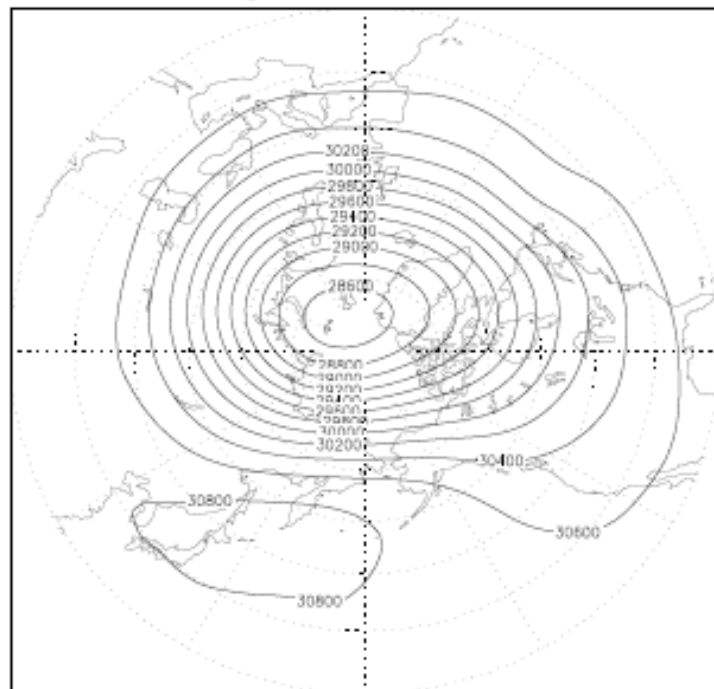
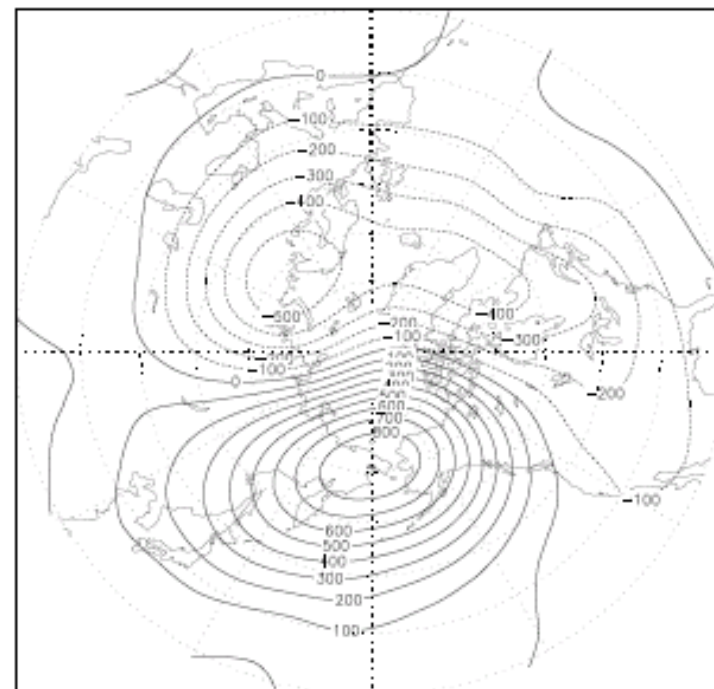


Figure 5. Wave-induced zonal force per unit mass, \bar{G} , in units of meters per second per day, for January 1993, deduced indirectly from the momentum balance (equation (1)) using a radiatively derived meridional circulation and the observed zonal mean wind. Contours are at $2 \text{ m s}^{-1} \text{ d}^{-1}$ intervals, with additional contours of ± 1 and $\pm 0.5 \text{ m s}^{-1} \text{ d}^{-1}$. The eastward force centered near 60°S and 40 hPa may be an artifact caused by the ill-conditionedness in calculating radiative heating rates from observed temperatures, especially in the summertime lower stratosphere. (After *Rosenlof* [1995]).

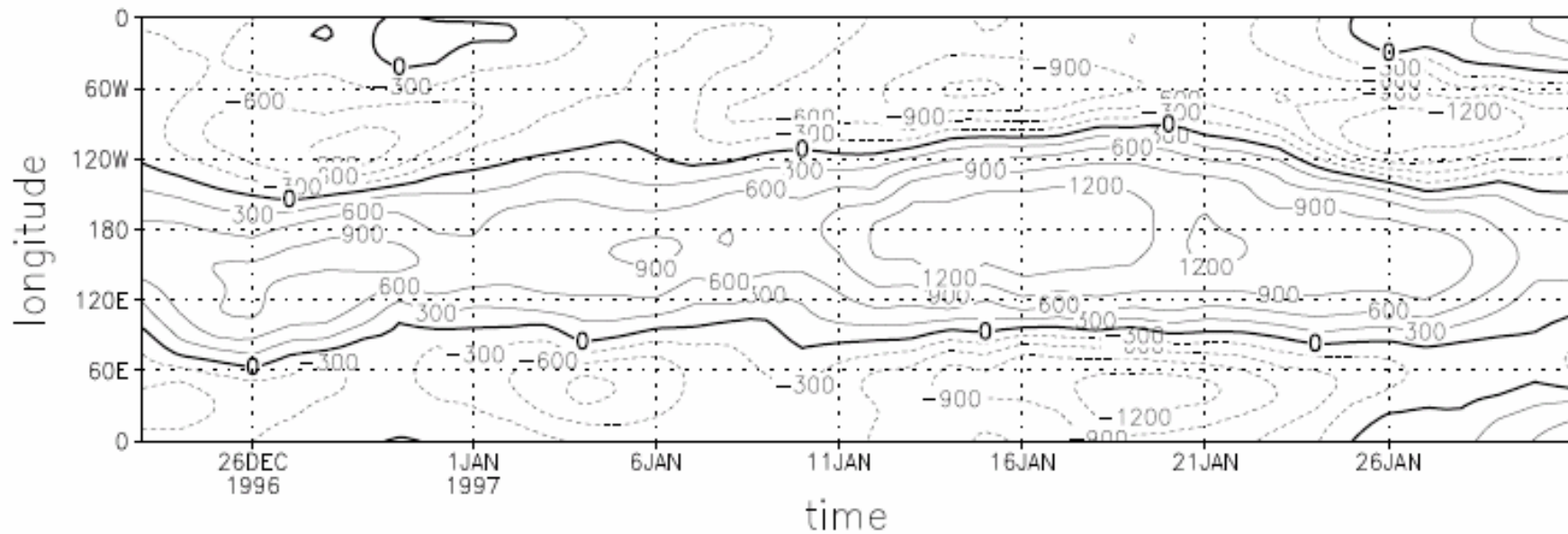
Geopotential Height 10 hPa, Dec 96 – Feb 97



Z-<Z>, 10 hPa Dec 96 – Feb 97

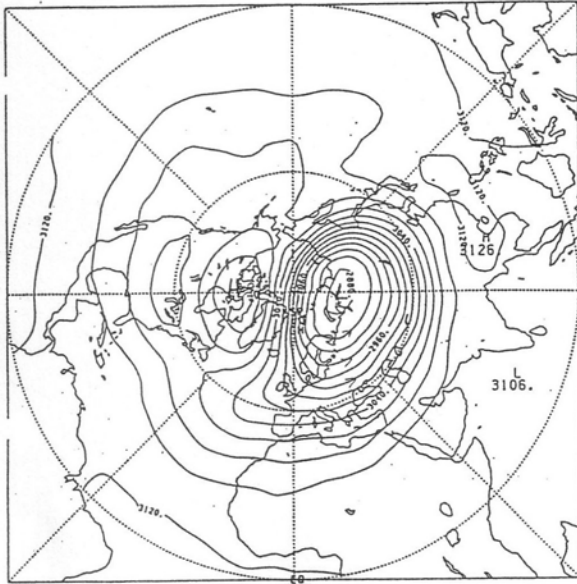


Z - <Z>, 60N 10hPa



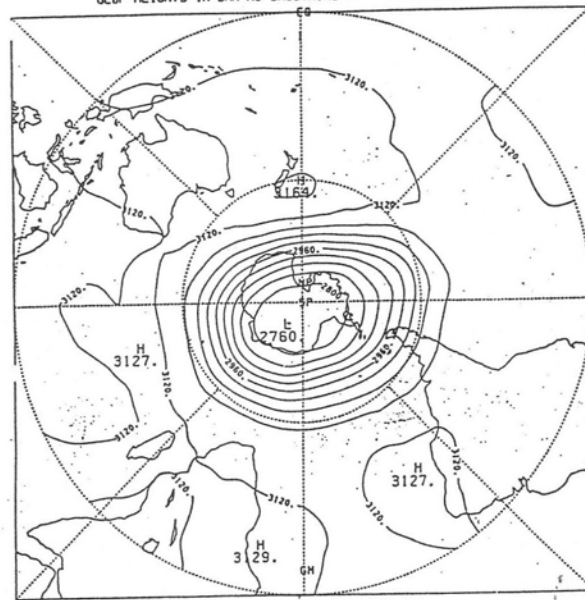
GEOP. HEIGHTS IN DAM NO SMOOTHING AT 10MB ON 27/11/79

NH



1.8(a)

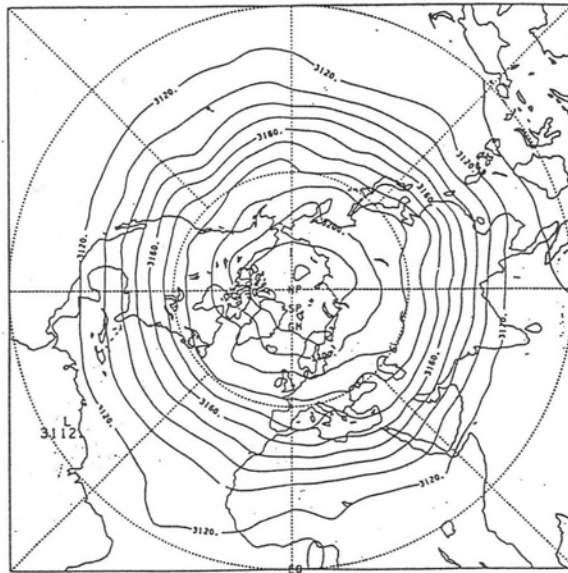
GEOP. HEIGHTS IN DAM NO SMOOTHING AT 10MB ON 14/ 9/79



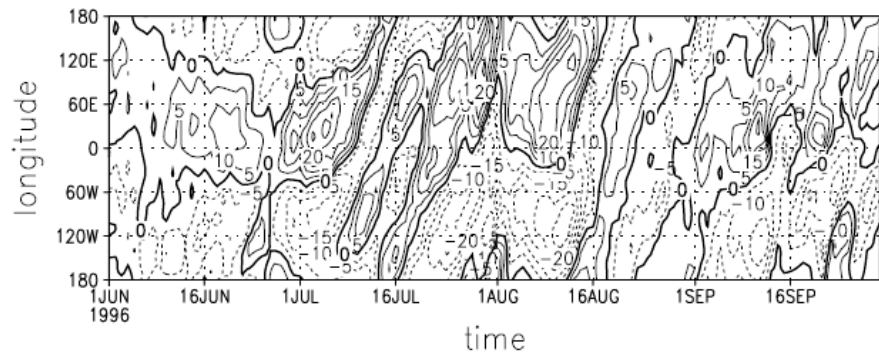
1.8(b)

summer NH

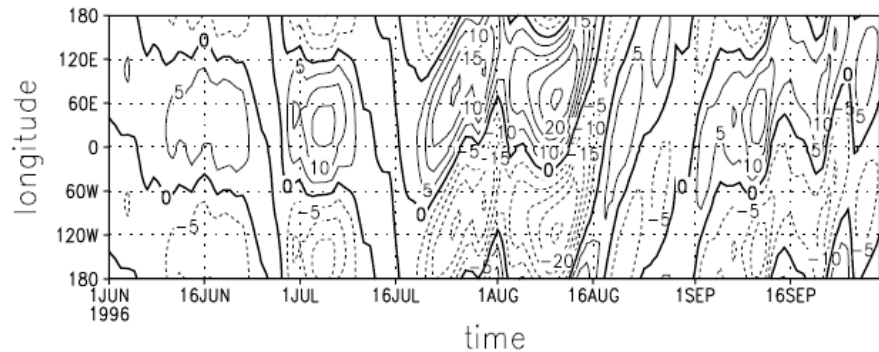
GEOP. HEIGHTS IN DAM NO SMOOTHING AT 10MB ON 26/ 7/79



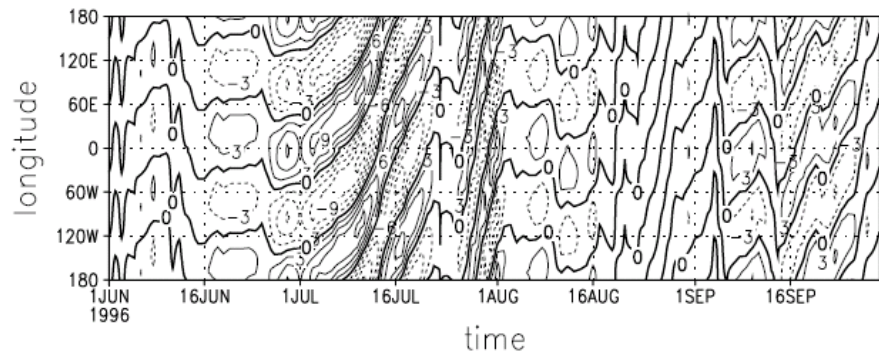
T - <T> at 58S, 10 mb



wave 1 T

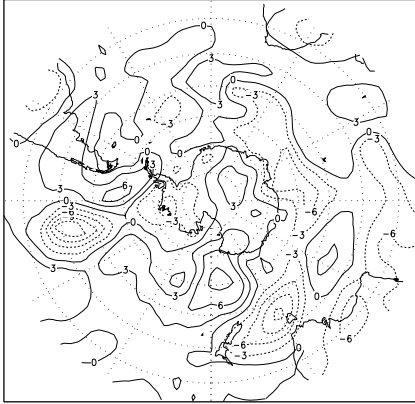


wave 2 T

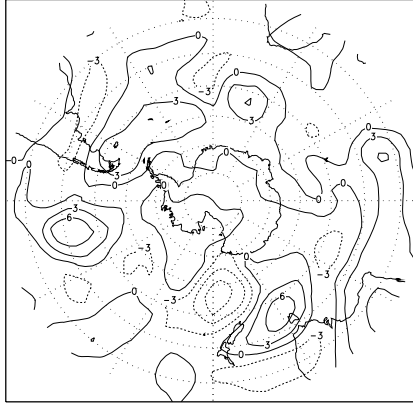


Jan 28 1996, SH

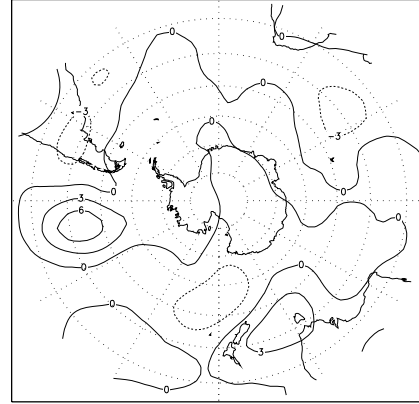
500 mb



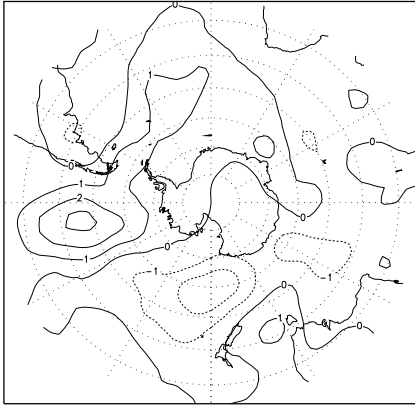
150 mb



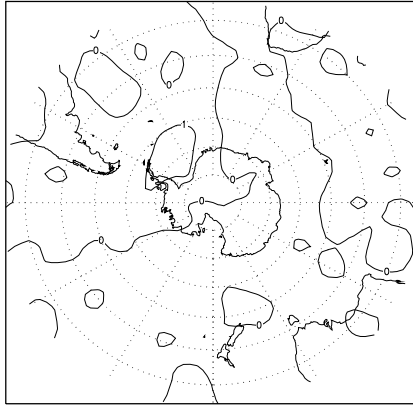
70 mb



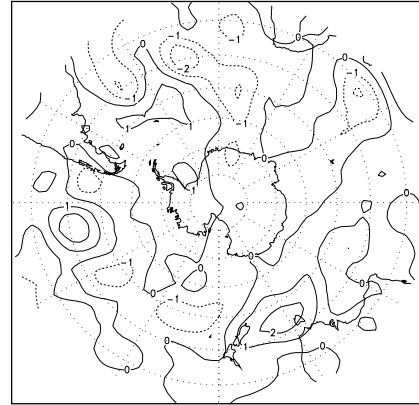
30 mb



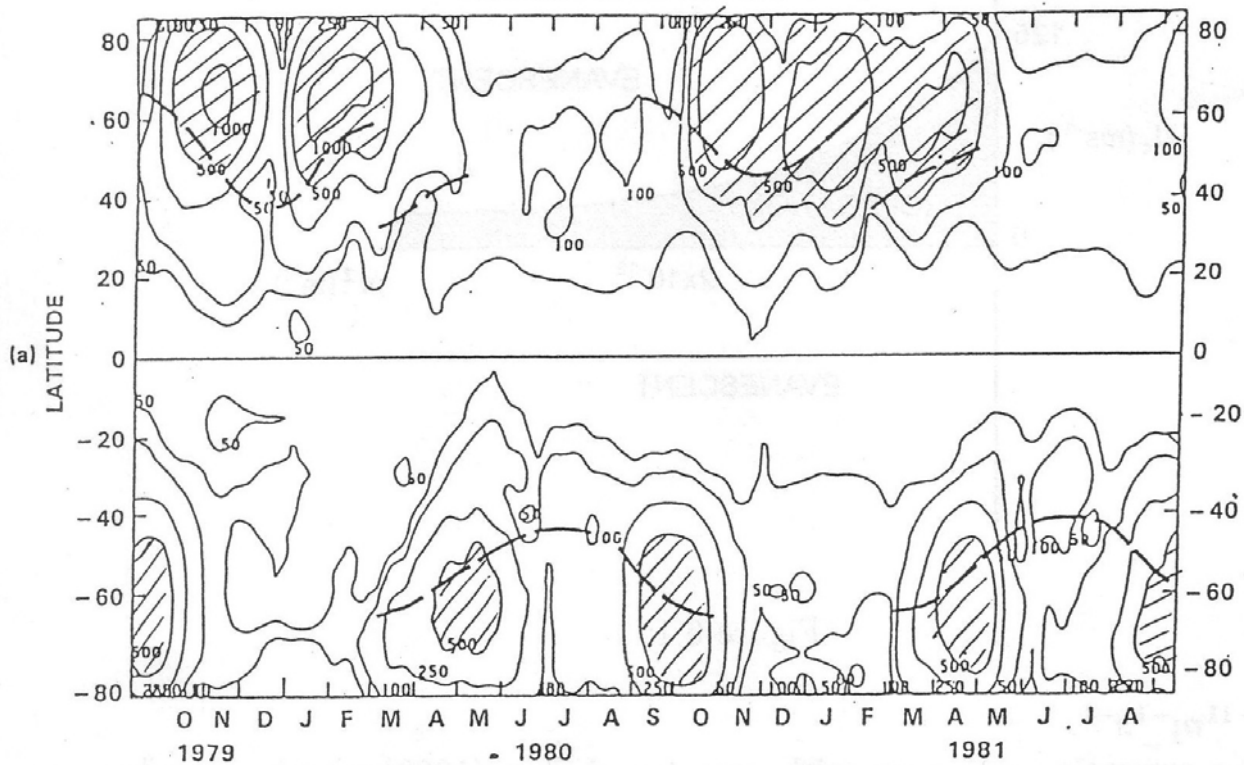
5 mb



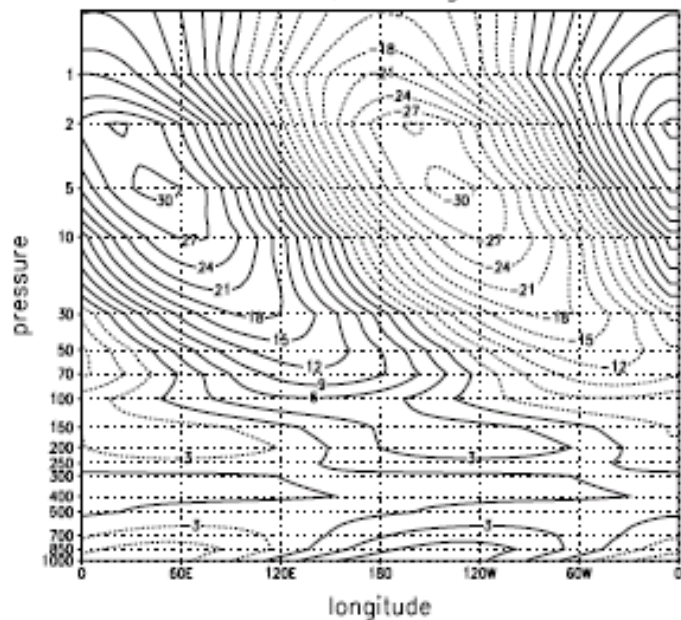
0.4 mb



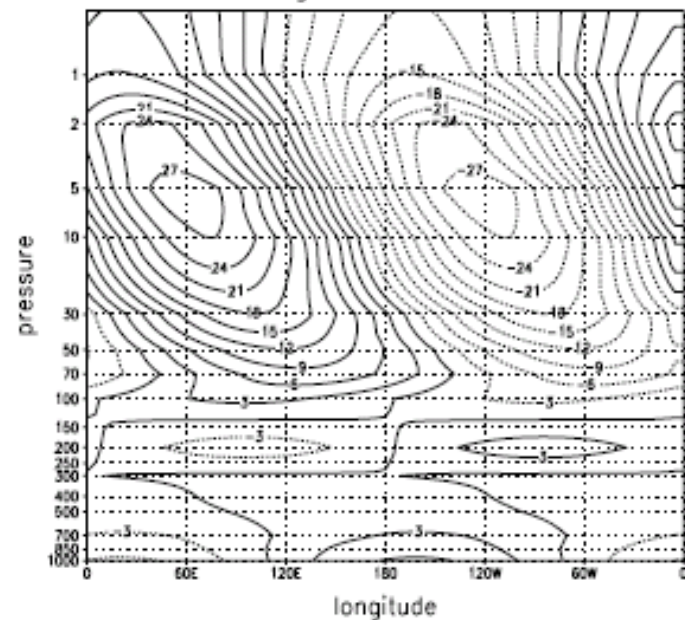
1 mb AMP. OF ST. WAVES (WN = 1)



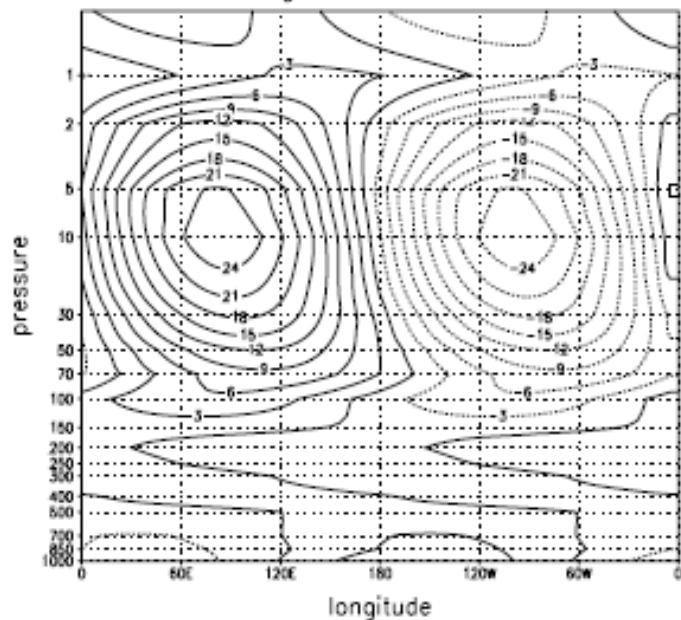
wave 1 T at 60S, on August 10 1996



August 11 1996



August 12 1996



August 13 1996

