

תורת האקלים

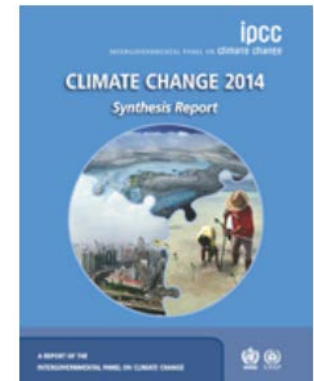
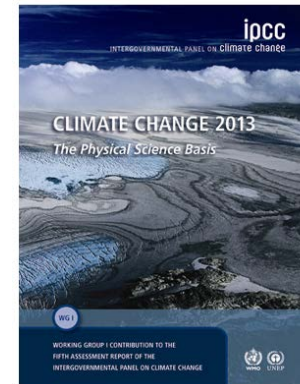
יום ה, 9-12, קפלון 205

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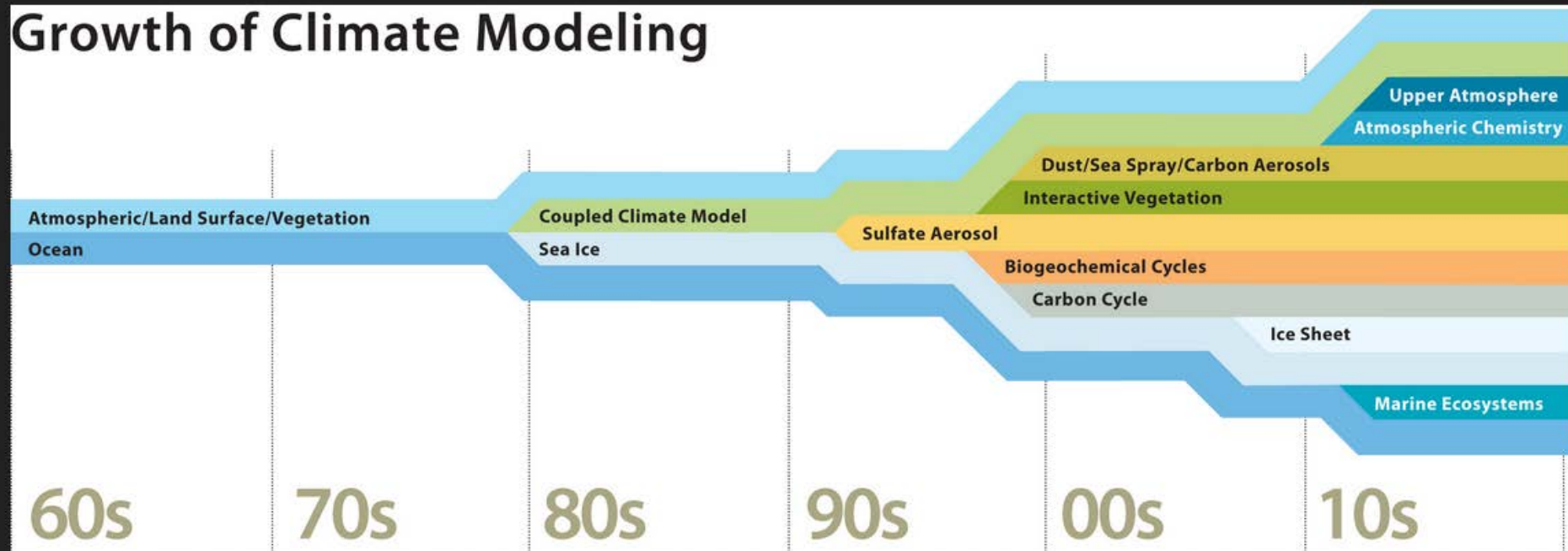
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# A brief history of Climate Science

Modern climate science is relatively new since it heavily relies on computer models which were only developed in the second half of the 20<sup>th</sup> century



## Growth of Climate Modeling



## The earliest thoughts and studies on Climate:

- Climatology- compilation of weather statistics for different geographical locations

“Climate is what you expect, weather is what you get”

- Ice ages – how could the climate have changed so much?
- The Greenhouse Effect (first suggested by Fourier in 1820)

# The Greenhouse Effect:

Joseph Fourier



Jean-Baptiste Joseph Fourier

**Fourier:** 1820s- the effective temperature of the Earth is below freezing. Suggested the atmosphere might trap heat

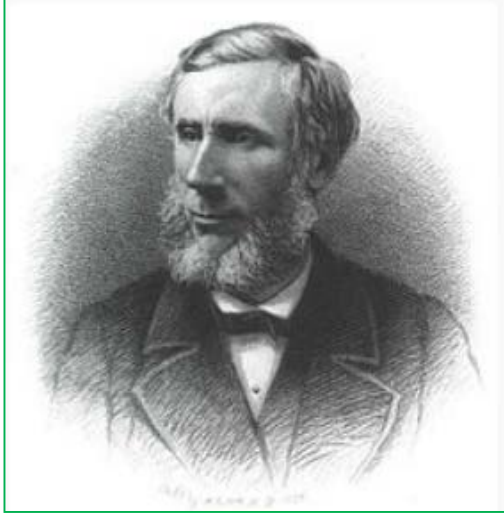
Claude Pouillet



Claude Servais Mathias Pouillet

**Pouillet:** 1837- first measurement of the solar constant

John Tyndall



**Tyndall:** 1859- Experimentally proved the possibility of the Greenhouse Effect by showing that atmospheric gasses absorb IR. Noted the dominance of water vapor

# Ice Ages:

18<sup>th</sup> -19<sup>th</sup> century: attribution of erratic boulders to the dispersion by glaciers (various scientists, including Geothe)



Yeager Rock  
(Washington state)

Karl Friedrich Schimper



German naturalist and poet

**Schimper:** 1830s- there existed times when the earth was much colder and most of Europe covered by ice - Called these times “Eiszeit” (Ice Ages)

Louis Agassiz



**Agassiz:** 1837- extended and presented the ice age idea. Community rejected this idea because contradicted established idea that Earth has been gradually cooling since its creation.

Lived on a glacier to study its motion.

1840- published book “Study on Glaciers”

# Ice Ages:

James Croll  
FRS



**Croll:** 1875 – “Climate and time in their geological relations” - tied changes in climate and glaciation to orbital variations.

His specific mechanism was wrong but the basic idea of orbital forcing effect was correct.

## Ice Ages and the Greenhouse Effect:

Svante Arrhenius



**Arrhenius:** 1896 – calculated how changes in CO<sub>2</sub> could affect Earth’s temperature, to explain ice ages. His formula relating radiative forcing change to CO<sub>2</sub> change still holds today:  $\Delta F = \alpha \ln(C/C_0)$

Predicted that fossil fuel burning would result in global warming, and included a water vapor feedback. Treated CO<sub>2</sub> as forcing water vapor changes because it is long lived.

Estimated doubling of CO<sub>2</sub> in 1000 years.



# Milutin Milankovich (1879-1958)

Milutin Milanković



Milankovitch as a student with a pocket watch.

A Serbian mathematician and civil engineer, who became interested in climate when read about ice ages and realized that meteorology was mostly empirical, a few physical theories used, and hardly any math.

Correctly calculated the intensity, temporal variations and latitudinal distribution of solar insolation

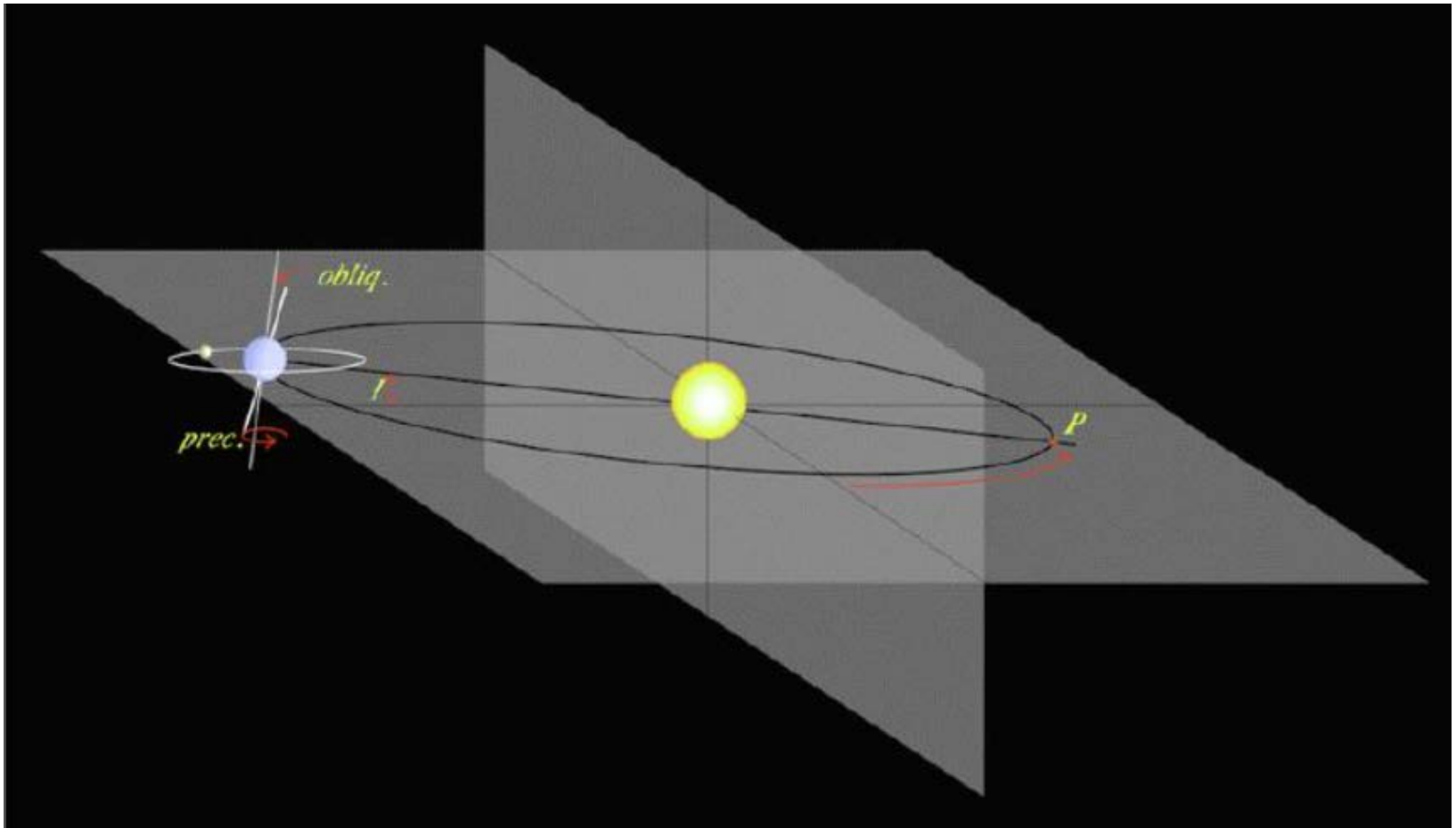
Applied this theory to study climate change and the climates of other planets.

1920- published a comprehensive book:  
*Mathematical Theory of Heat Phenomena Produced by Solar Radiation.*

Milankovic's theory:

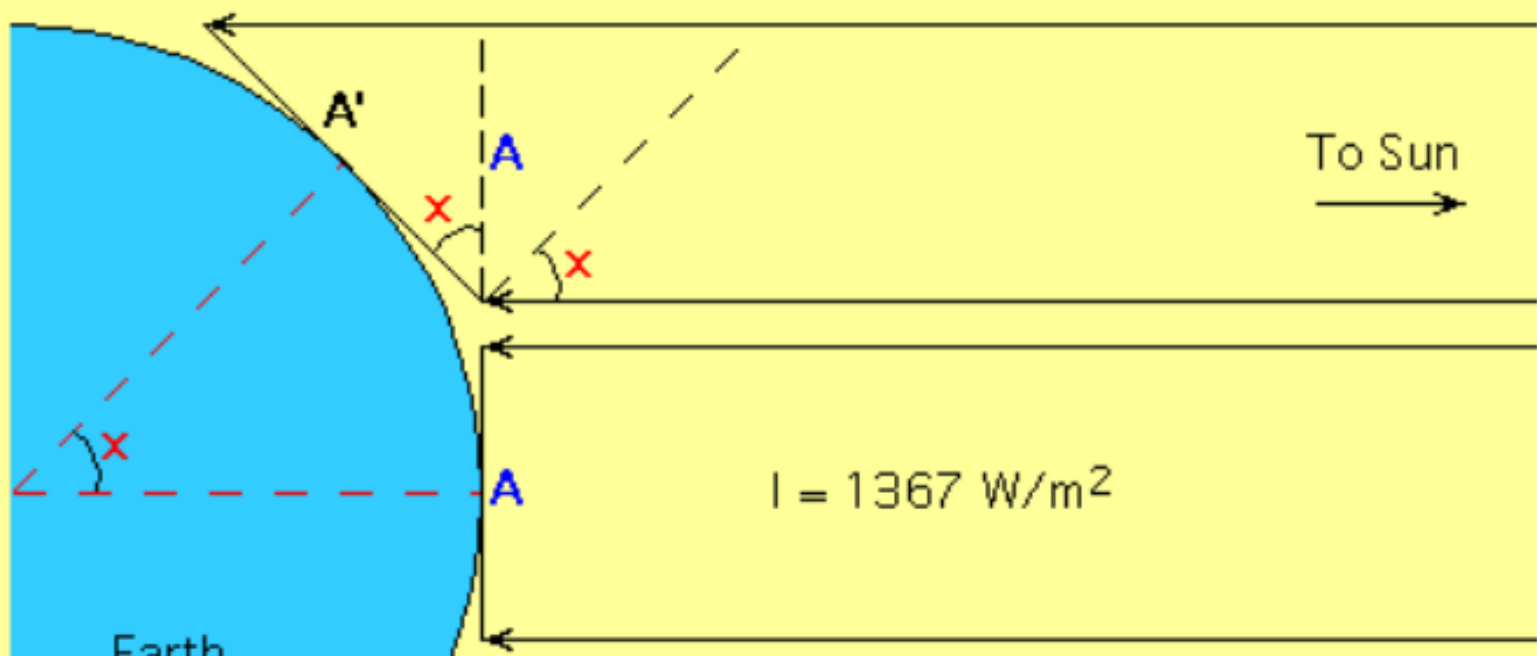
The planets and moon cause variations in the Earth's orbit around the sun, affecting insolation.

Main variations: eccentricity, obliquity and precession





As the angle of incidence (x) increases, insolation decreases.

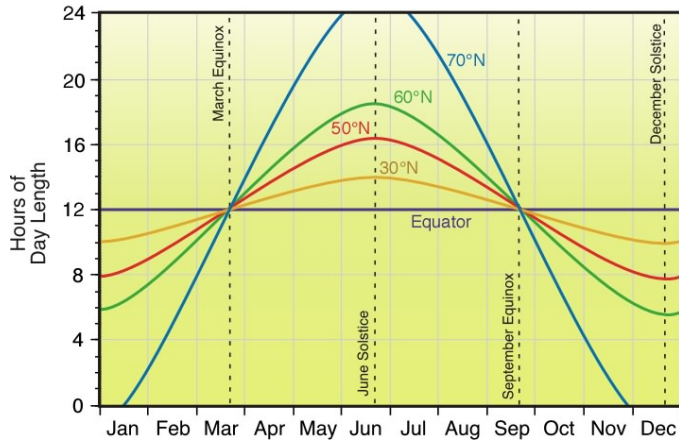
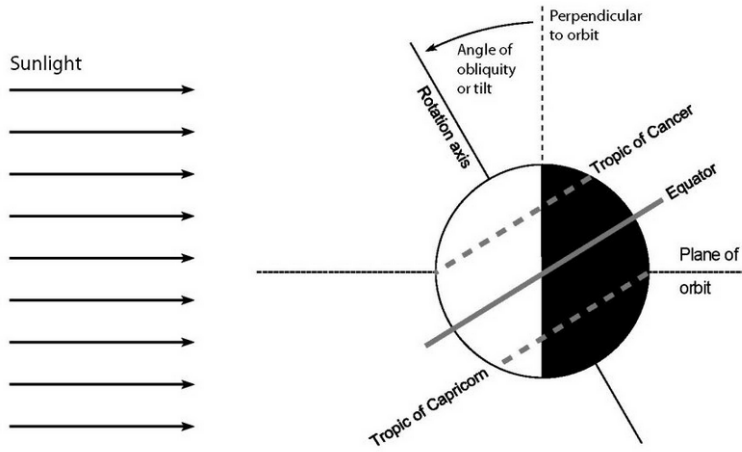


$x$  = degrees latitude away from subsolar point.

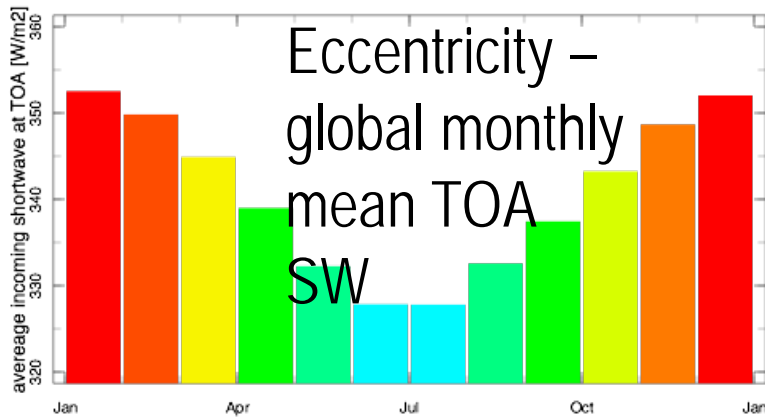
$$\cos x = \frac{A}{A'}$$

$$A' = \frac{A}{\cos x}$$

# Daily mean incoming solar radiation at top of the atmosphere ( $W/m^2$ )

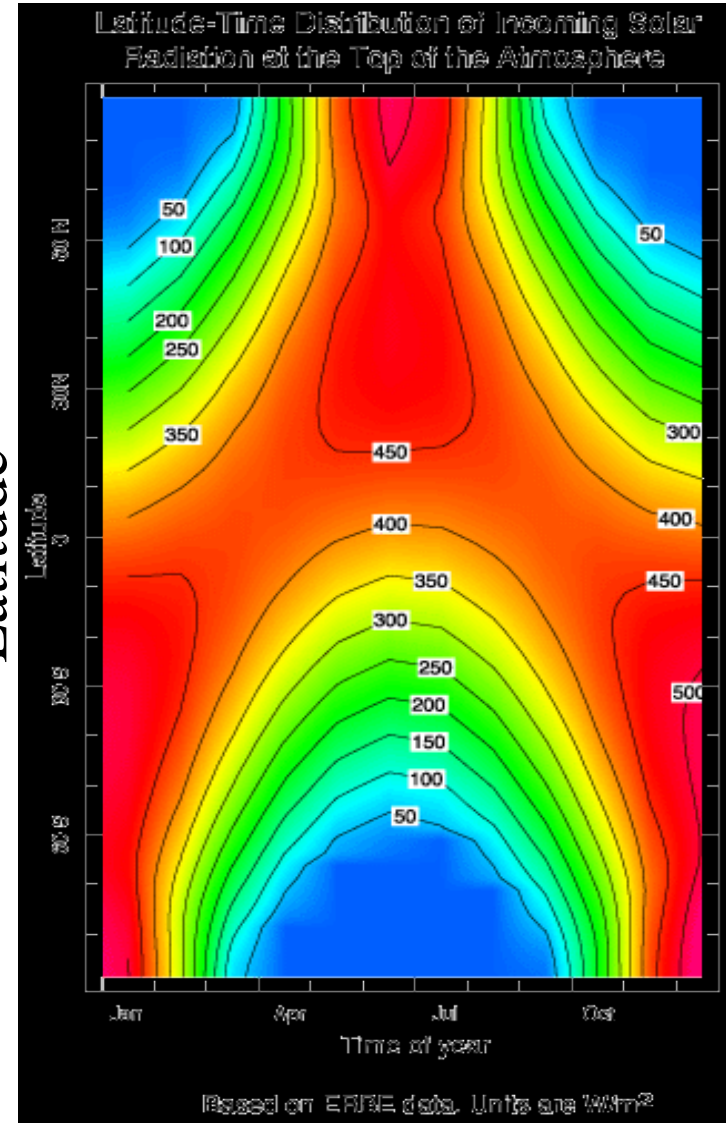


Length of day



Eccentricity – global monthly mean TOA SW

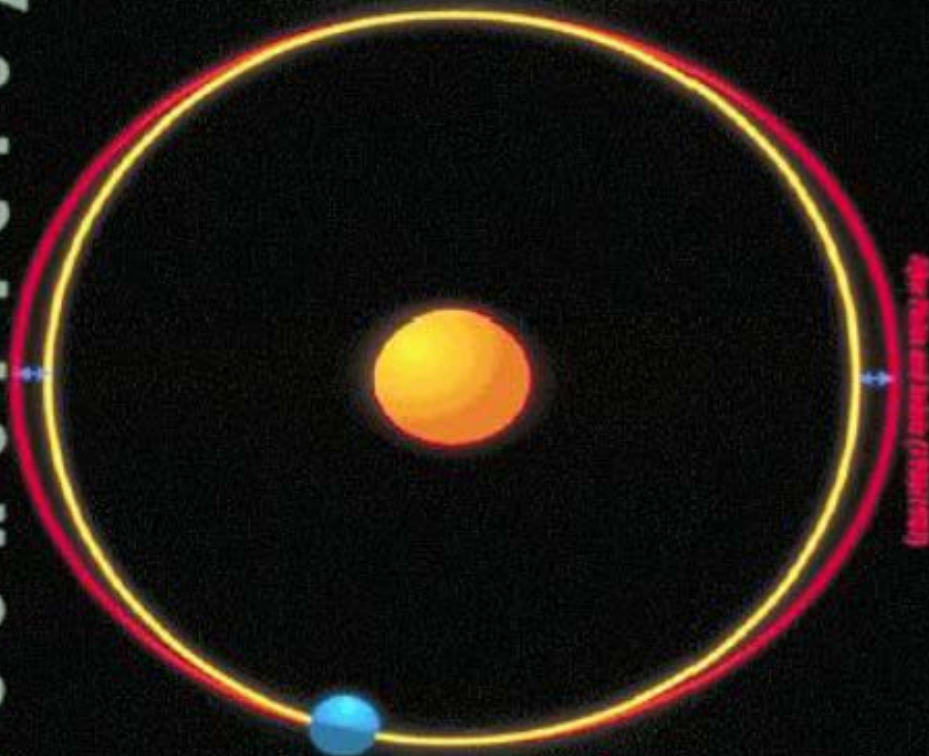
Latitude



Based on ERBE data Time of year

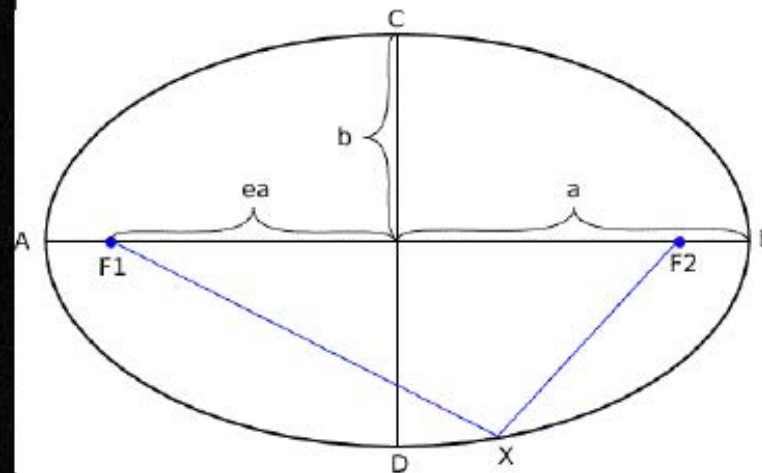
# eccentricity

The orbit of the Earth changes from nearly circular (eccentricity equal to 0.00) to more elliptical (eccentricity equal to 0.06). These changes occur in two broad frequency bands: one at periods of around 100,000 years and one at periods near 400,000 years.



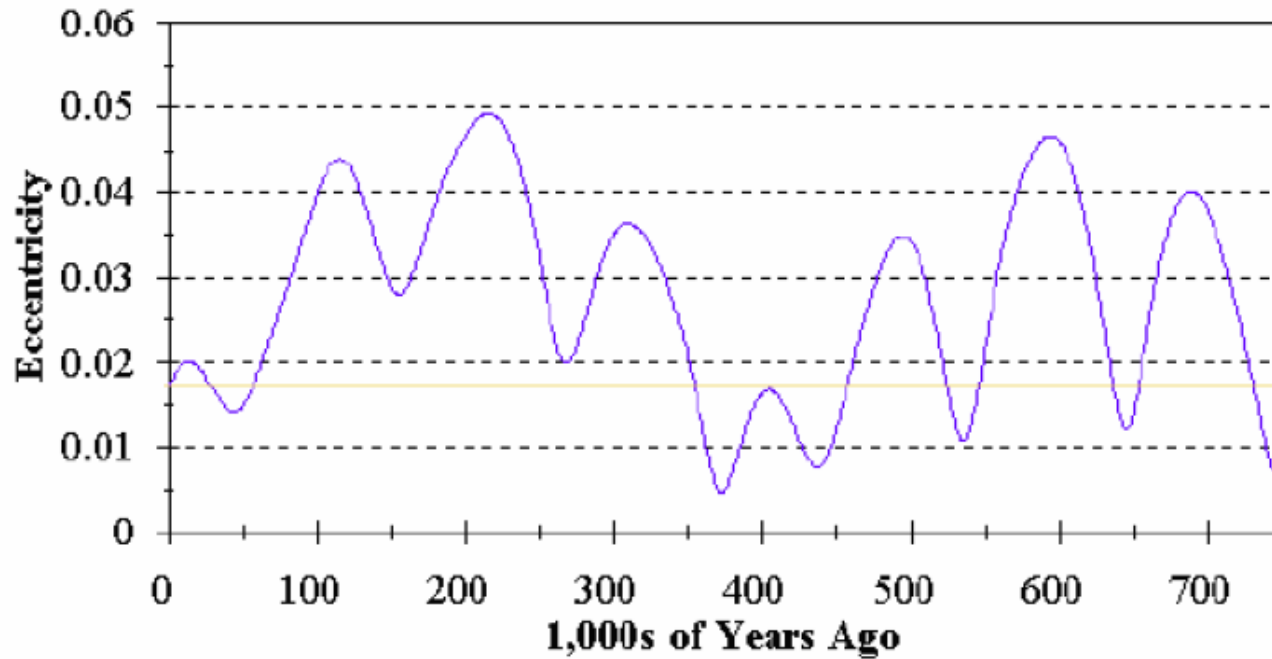
While variations in orbital eccentricity have a small impact on the total amount of radiation received at the top of Earth's atmosphere (ca. 0.1 percent), the primary importance of the eccentricity cycles is to modulate the amplitude of the precession cycle. When eccentricity is high (more elliptical), the effect of precession on the seasonal cycle is strong. When eccentricity is low (more circular), the position along the orbit at which the equinoxes occur is irrelevant since all points on the orbit become, in effect, perihelia.

## Eccentricity



$$e = \sqrt{1 - \frac{b^2}{a^2}}$$

## Orbital Eccentricity



100Kyr and 400Kyr periodicities. Small variations of 0.01% in insolation.

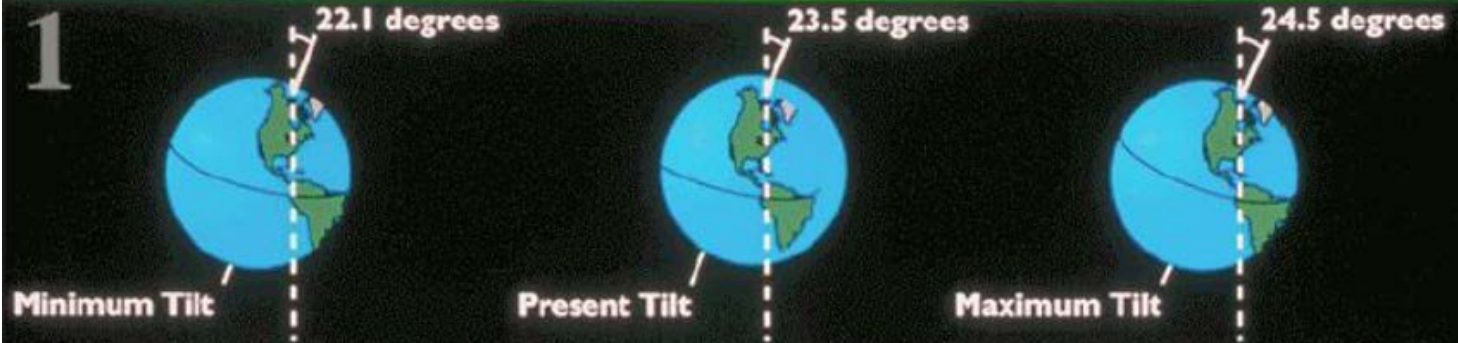
Eccentricity affects the global mean annual mean insolation.



# obliquity

*Earth's axial tilt varies from 24.5 degrees to 22.1 degrees at periods of close to 41,000 years.*

1



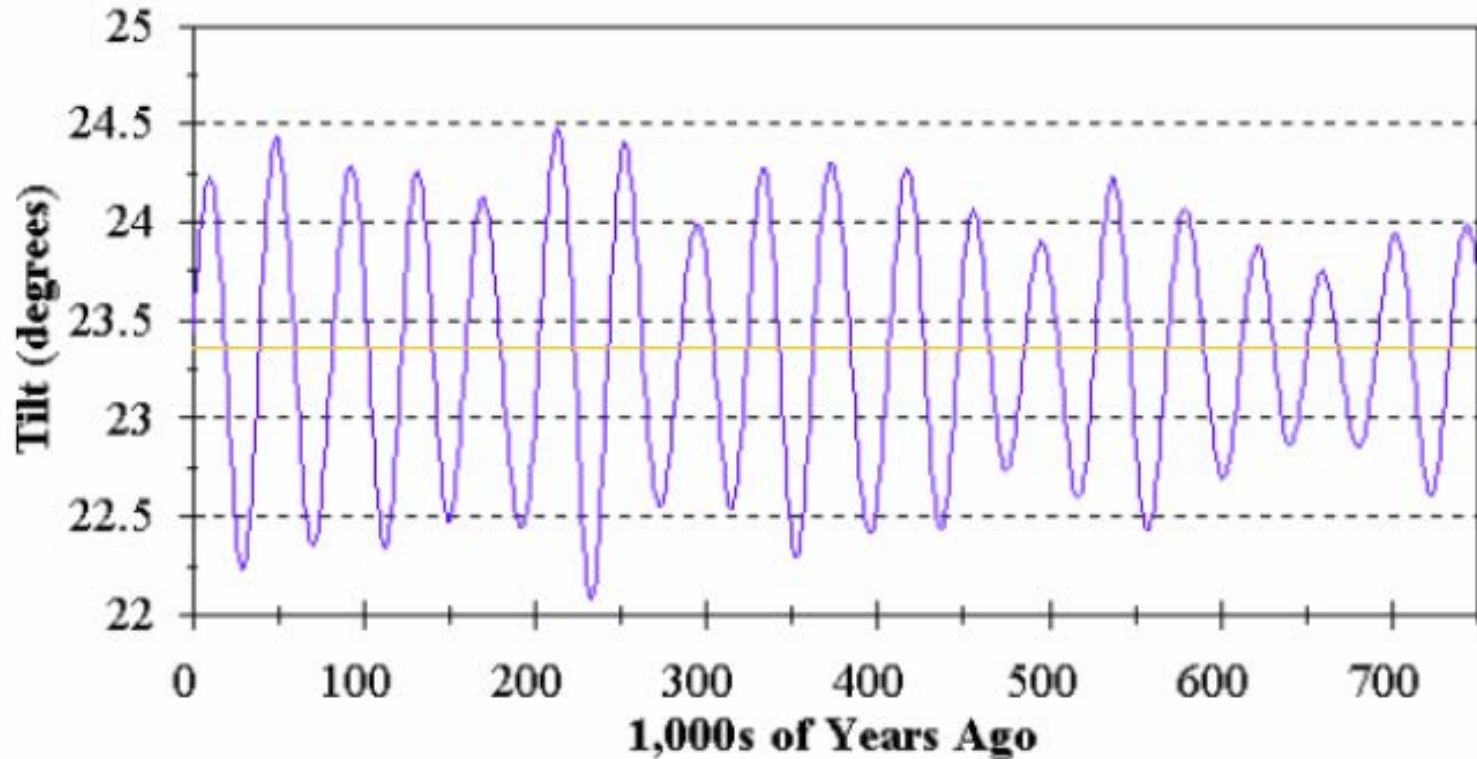
*Axial tilt affects the distribution of solar radiation on Earth's surface. When the tilt is decreased polar regions receive less sunlight; when it is increased, polar regions receive more sunlight.*

2



*After Peck & Imbrie (1982/1987)*

## Tilt



41Kyr periodicity dominant.

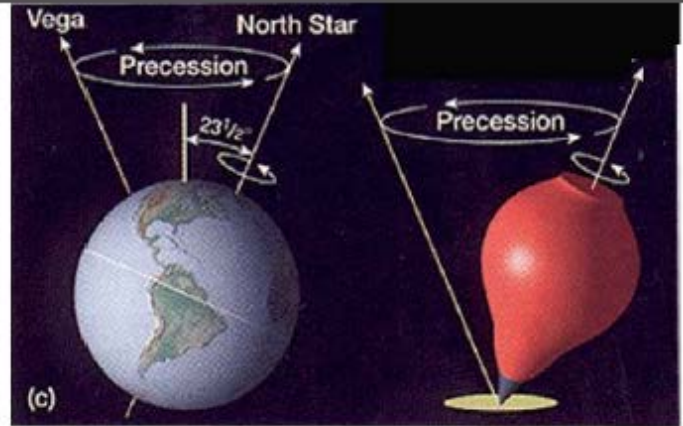
Latitude of peak insolation moves about 1km each century.

Global mean annual mean insolation unaffected, but high latitude insolation (important for ice sheet formation) varies strongly



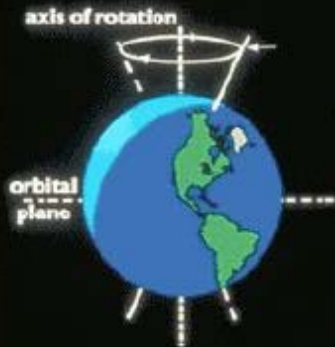


# Precession



Like a spinning top, Earth's axis of rotation "wobbles," so that the North Pole describes a circle in space

The 'wobble' of the Earth's axis causes the precession of the equinoxes. As shown in this figure, the positions of the equinoxes and solstices shift slowly around the Earth's elliptical orbit, completing one full cycle every 22,000 years. Precession changes the time at which the Earth reaches its perihelion (the point on the orbital path closest to the Sun), serving to amplify or soften climatic seasonality.



# P r e c e s s i o n

## Top view

## Side view

Today: Perihelion during northern winter.

(North Pole tilted away from Sun=northern winter)



5,500 years ago: Perihelion during northern spring.



11,000 years ago: Perihelion during northern summer.

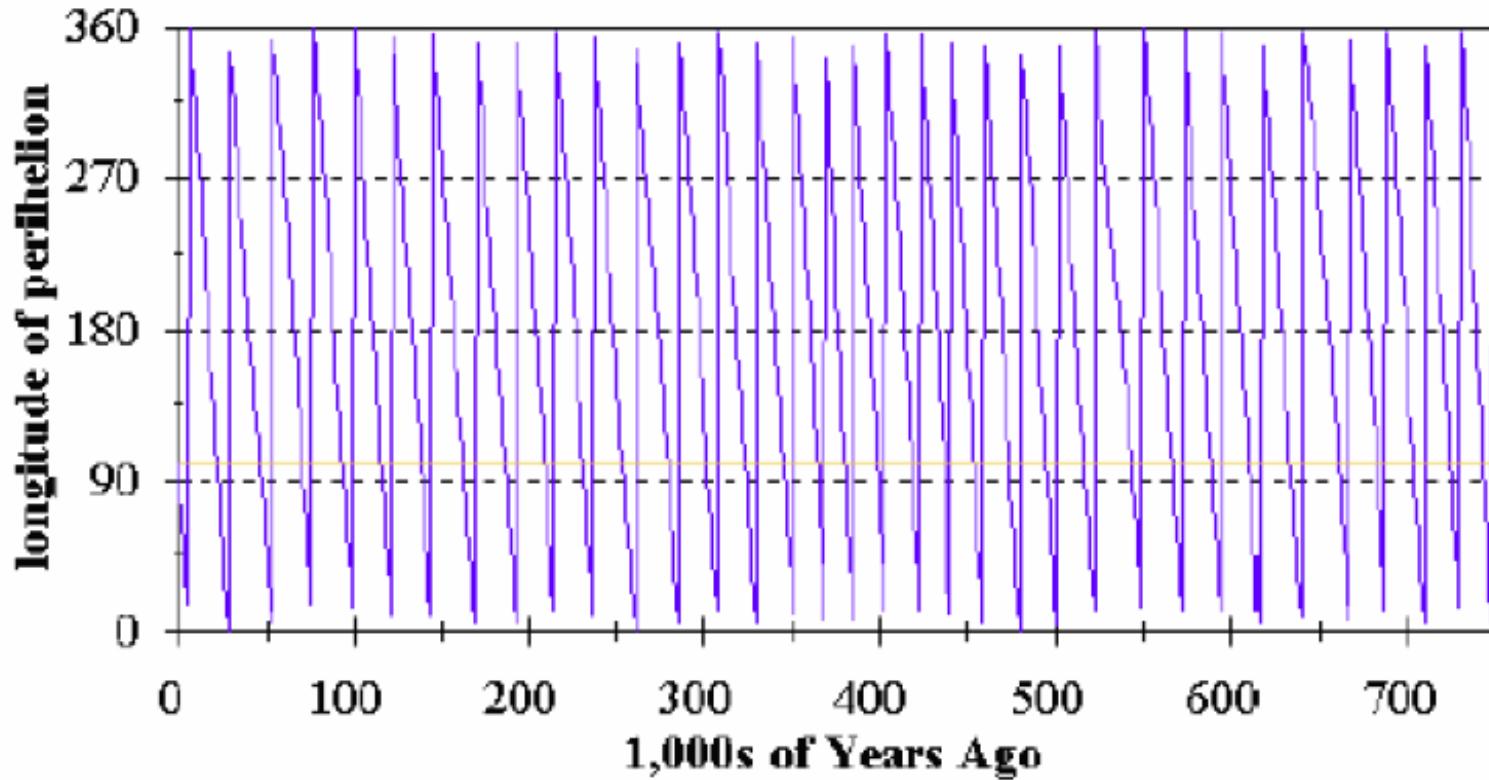
(North Pole tilted towards Sun=northern summer)



● Earth on December 21 ● Sun ● Perihelion

# Precession of Equinox

(Longitude of Earth at Equinox relative to Perihelion)

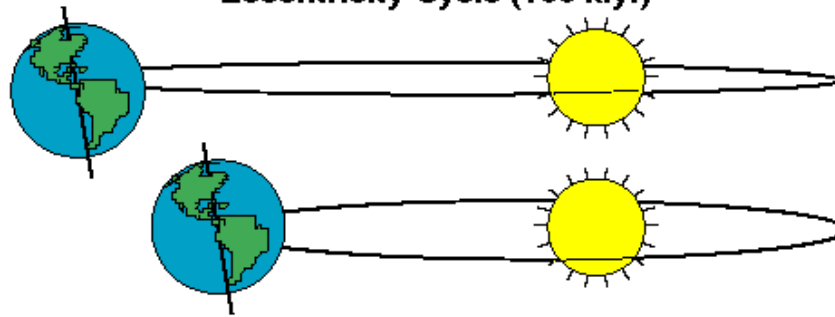


23Kyr and 19Kyr periodicities (low resolution peak of 22Kyr)

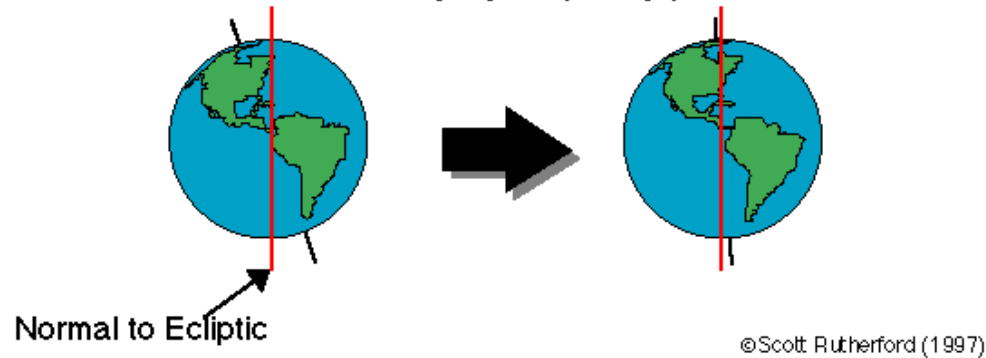
Global mean annual mean insolation unaffected.

Insolation depends on combination of eccentricity  $e$  and precession angle  $w$  :  $e \sin(w)$

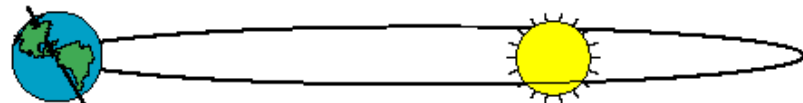
### Eccentricity Cycle (100 k.y.)



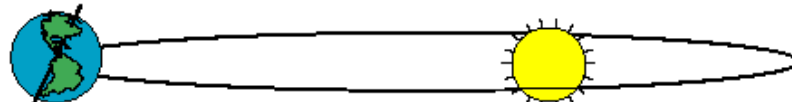
### Obliquity Cycle (41 k.y.)



### Precession of the Equinoxes (19 and 23 k.y.)

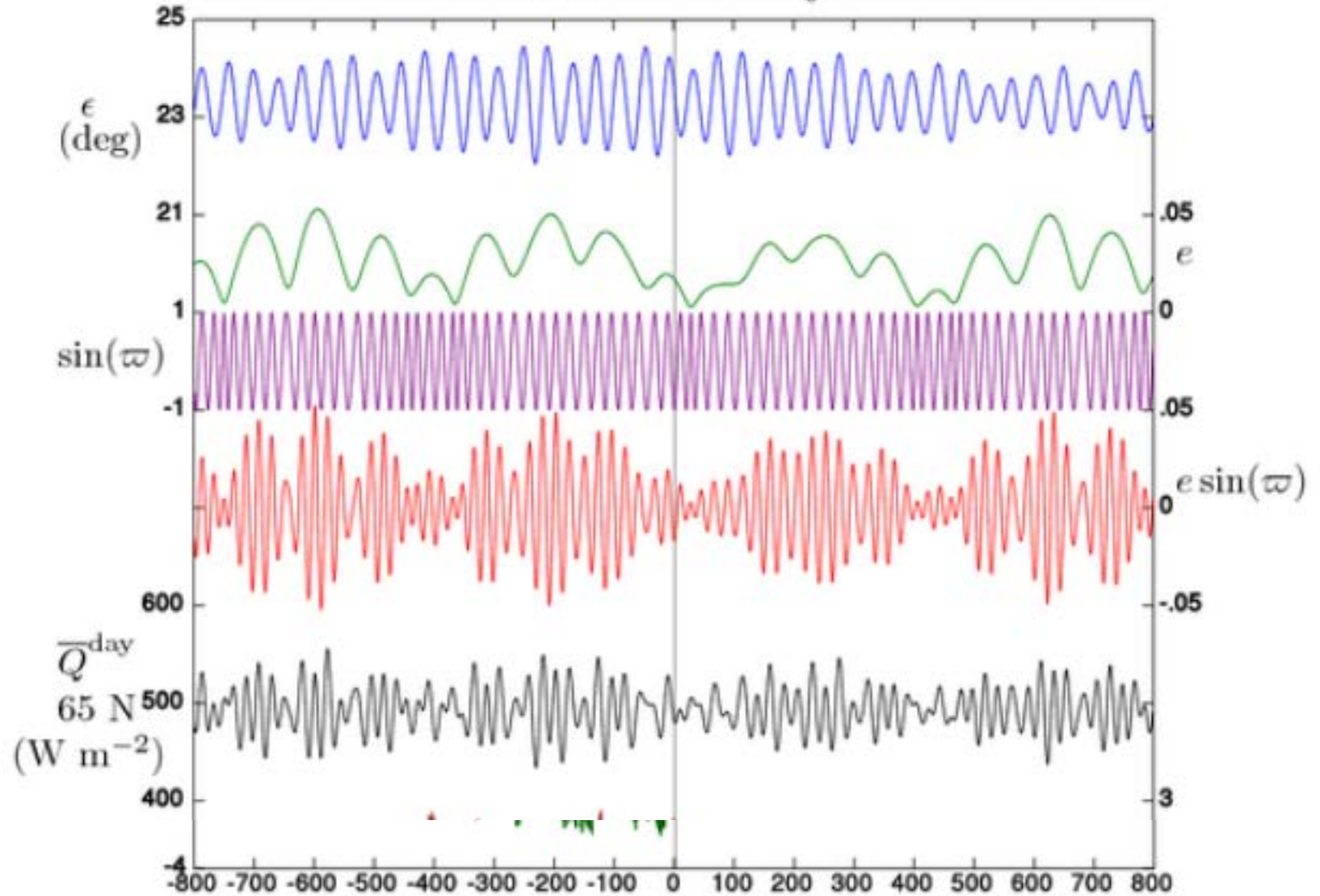


Northern Hemisphere tilted away from the sun at aphelion.



Northern hemisphere tilted toward the sun at aphelion.

# Milankovitch Cycles

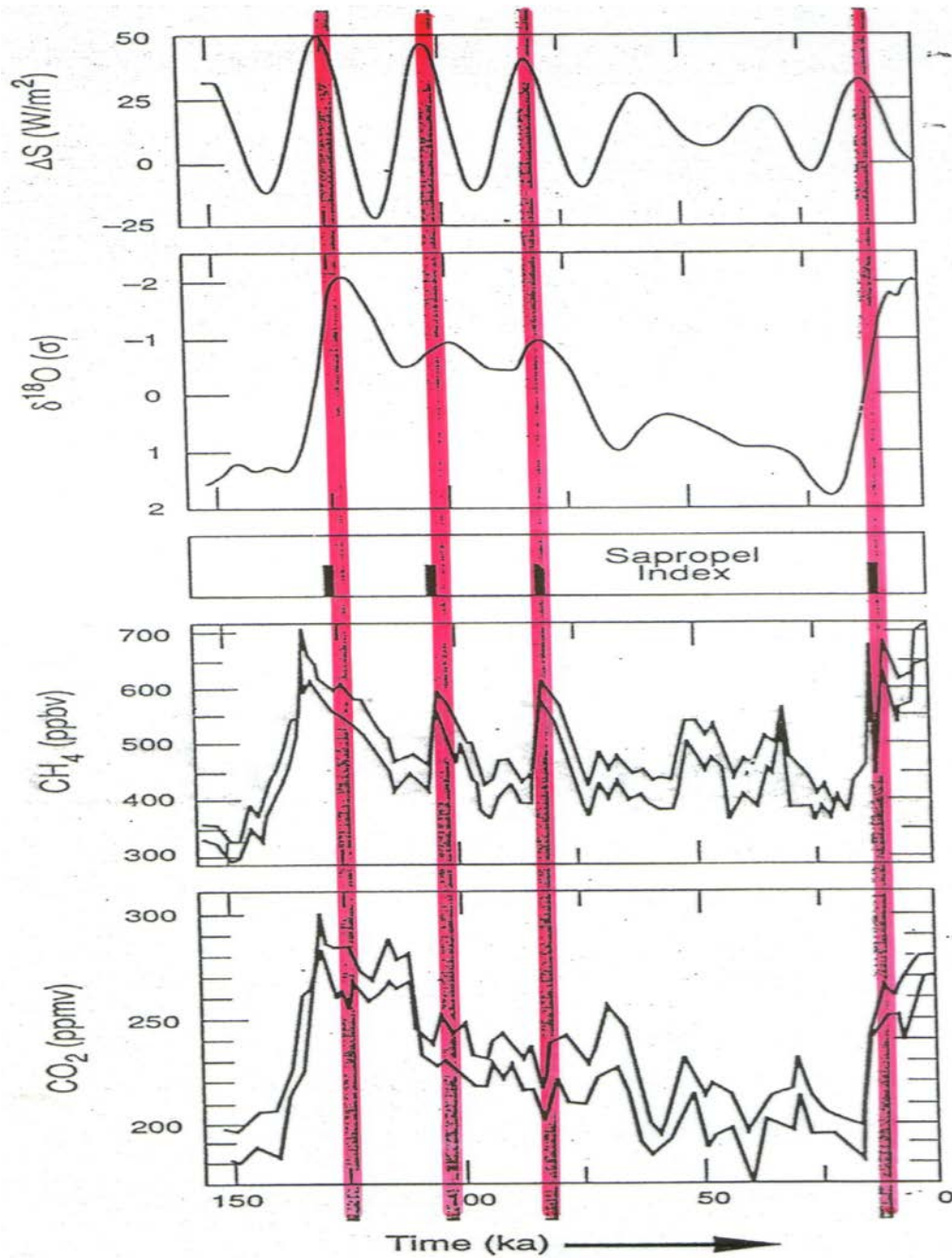


**Kyrs relative to present time**

We are at a period of minimum eccentricity.



# Is Milankovitch's theory of influence on climate correct?



**NH summer radiation**

**Volume of continental ice**

**Strength of African Monsoon**

**Methane concentrations from Vostok**

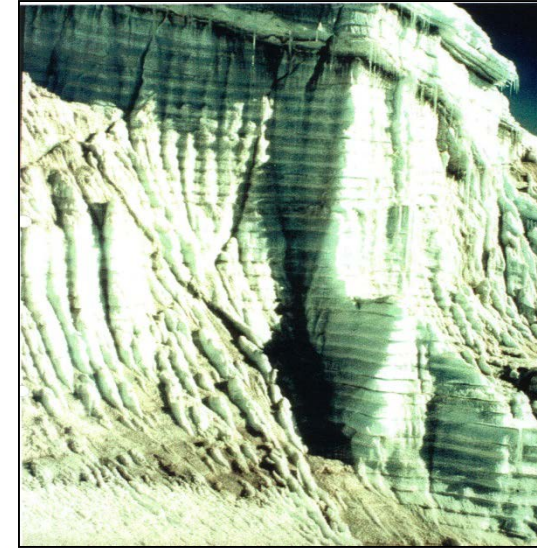
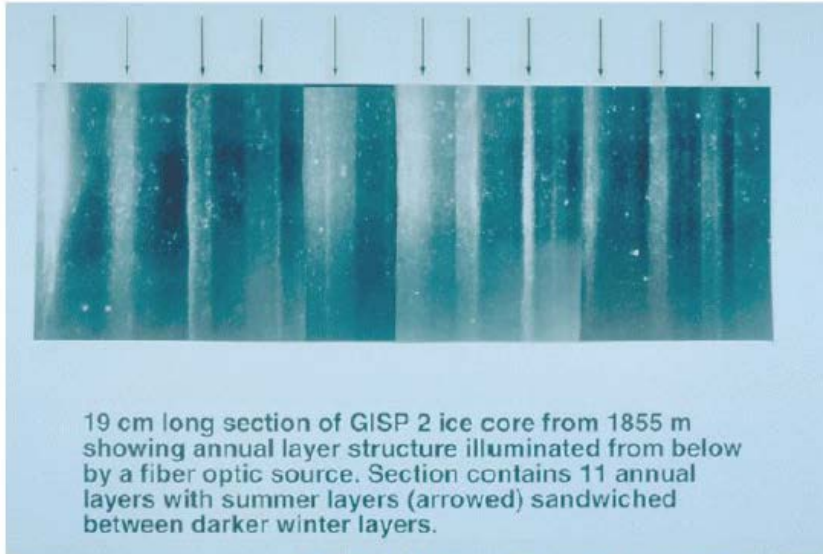
**CO<sub>2</sub> concentrations from Vostok**

## Ice Cores





## Example: 19cm representing 11 years



Can analyze  
both water and  
air bubbles



Analyze trace gases ( $\text{CO}_2$ ,  $\text{CH}_4$ ) in bubbles in ice.

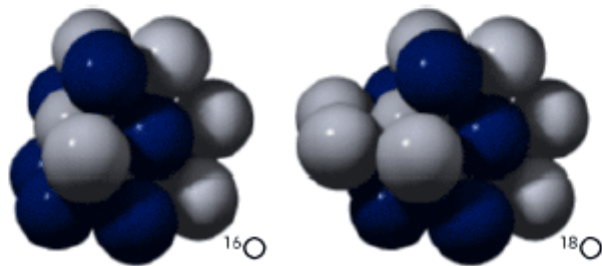
Determine temperature from oxygen isotopes ( $\text{O}^{18}/\text{O}^{16}$  ratios).

Useful in Greenland and Antarctica

Length of layered records: 10,000 up to 600,000 years or more for temperature, trace gases and accumulation.

# Isotopic Evidence

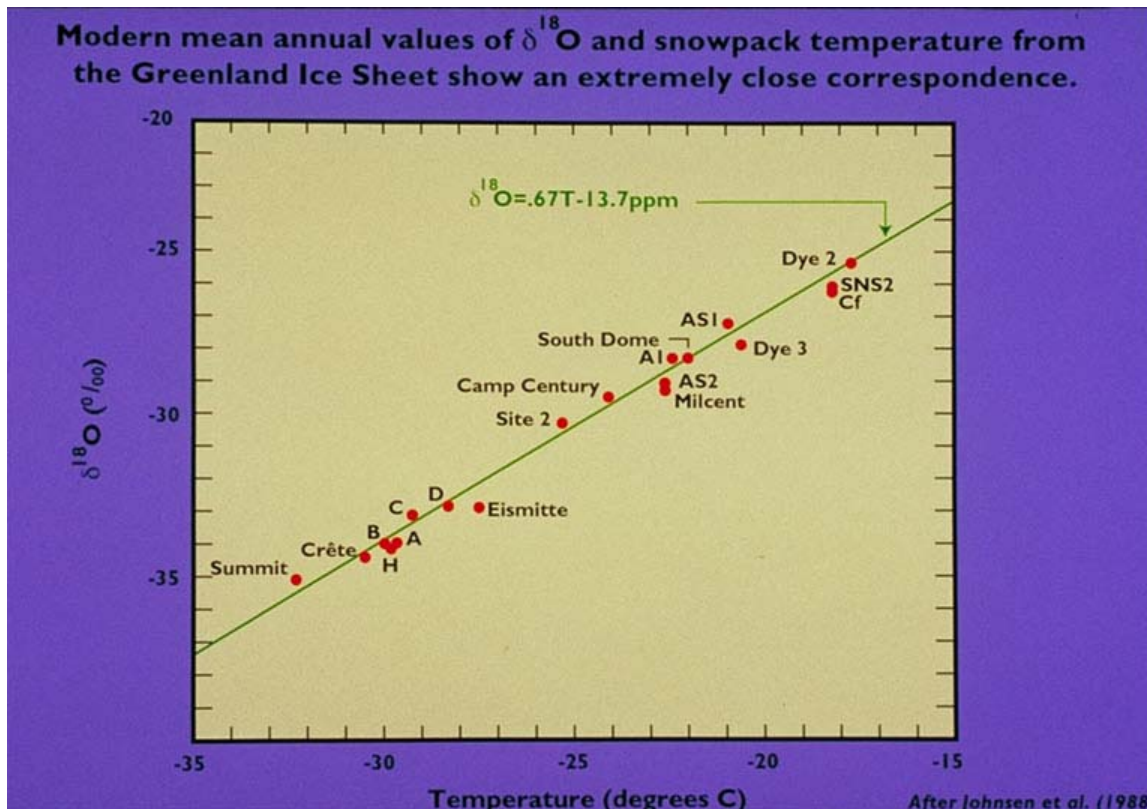
- Most elements have several stable isotopes, differing only by the number of neutrons and hence the weight of the atom and the molecule of which it is a part.
- For example: water,  $\text{H}_2\text{O}$ , can contain isotopes of hydrogen and oxygen.
- Hydrogen has  $\text{H}^1$  and  $\text{H}^2=\text{D}=\text{deuterium}$
- Oxygen has  $\text{O}^{16}$  (99.762%),  $\text{O}^{17}$  and  $\text{O}^{18}$ .
- Carbon has  $\text{C}^{12}$  (98.93%),  $\text{C}^{13}$  and  $\text{C}^{14}$  ( $\text{C}^{14}$  is radioactive with a half life of 5,730 years. It's made from Nitrogen.)



$$\delta^{18}\text{O} = \left( \frac{\left( \frac{^{18}\text{O}}{^{16}\text{O}} \right)_{\text{sample}}}{\left( \frac{^{18}\text{O}}{^{16}\text{O}} \right)_{\text{standard}}} - 1 \right) \times 1000 \text{‰}$$

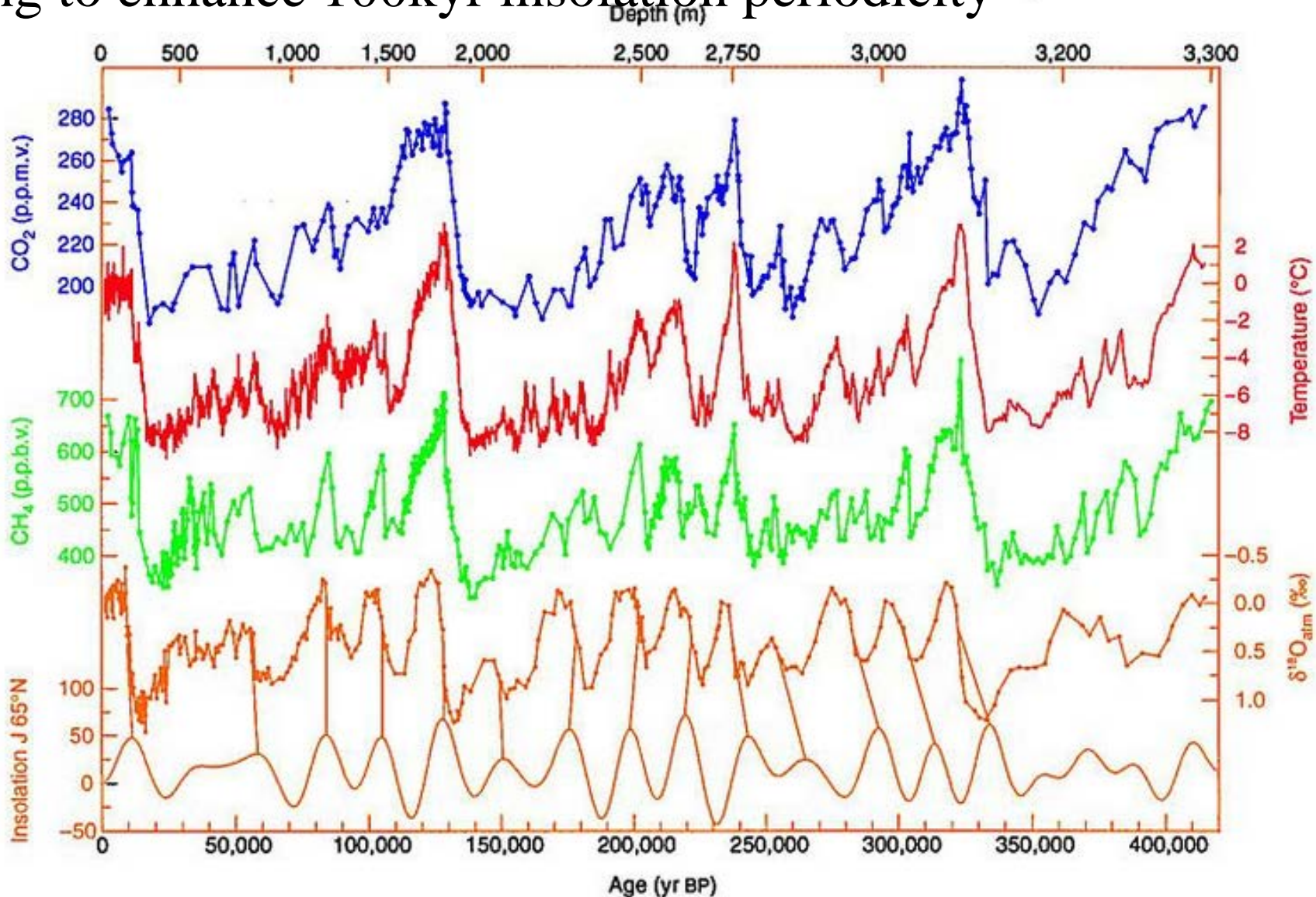
Measure anomaly from Standard Mean Ocean Water (SMOW)

This ratio is sensitive to sea temperature, and to global ice volume.  $^{16}\text{O}$  evaporates more easily



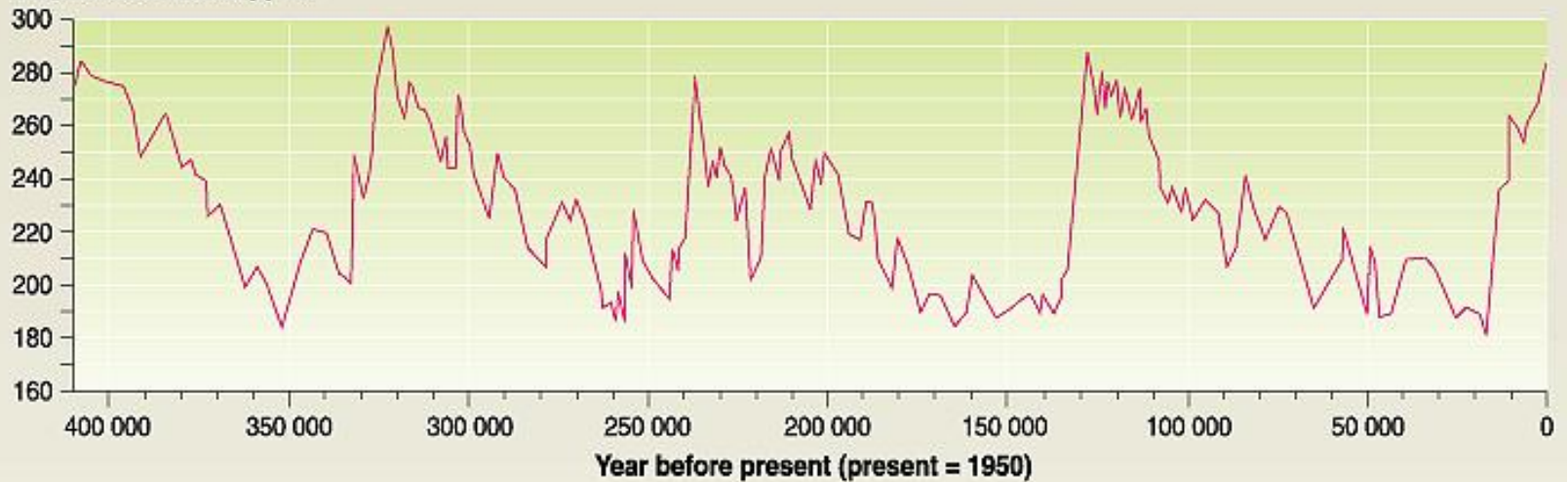


Vostok ice core: Temperature, CO<sub>2</sub>, methane and ice volume vary together, driven by insolation, with feedbacks acting to enhance 100kyr insolation periodicity

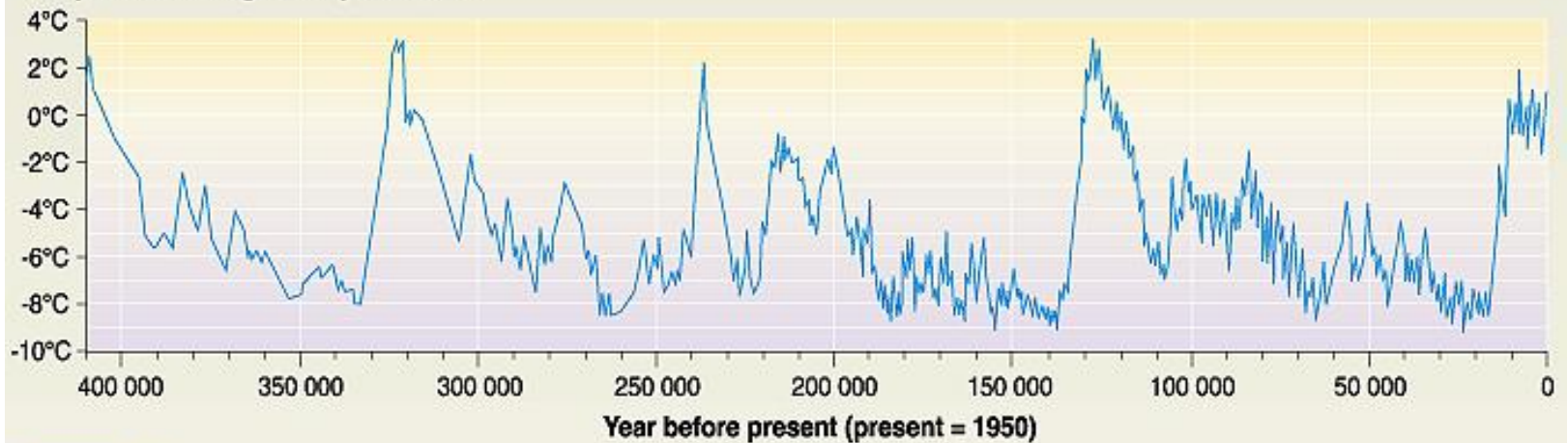


# Temperature and CO<sub>2</sub> concentration in the atmosphere over the past 400 000 years (from the Vostok ice core)

CO<sub>2</sub> concentration, ppmv



Temperature change from present, °C

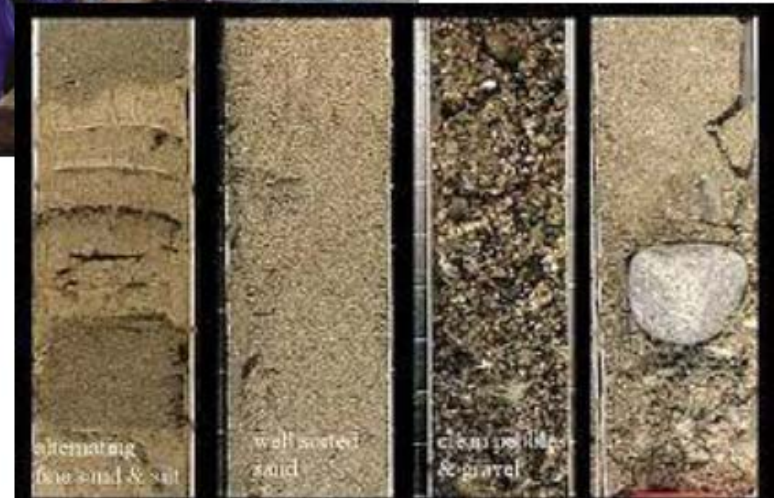




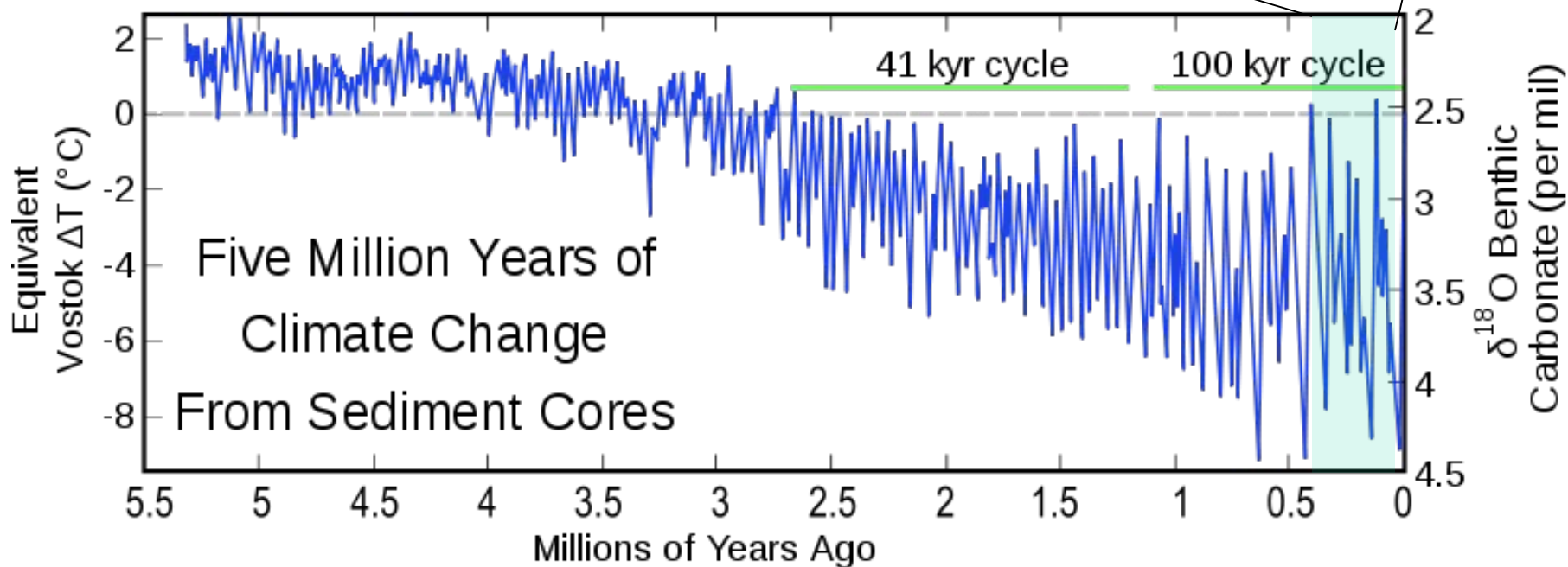
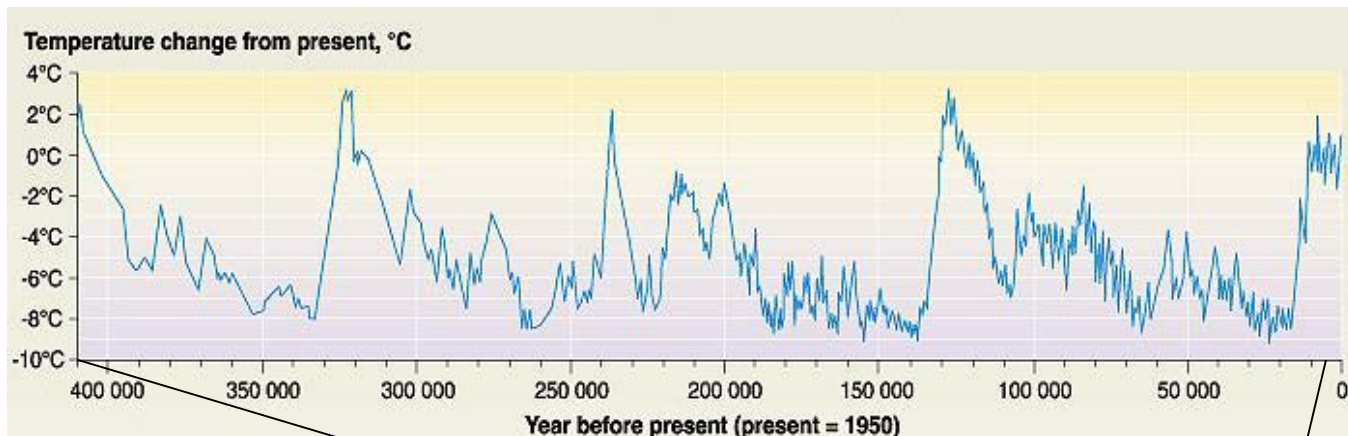
Lets put this in longer context – go back further in time.  
**Ocean Sediment Cores**



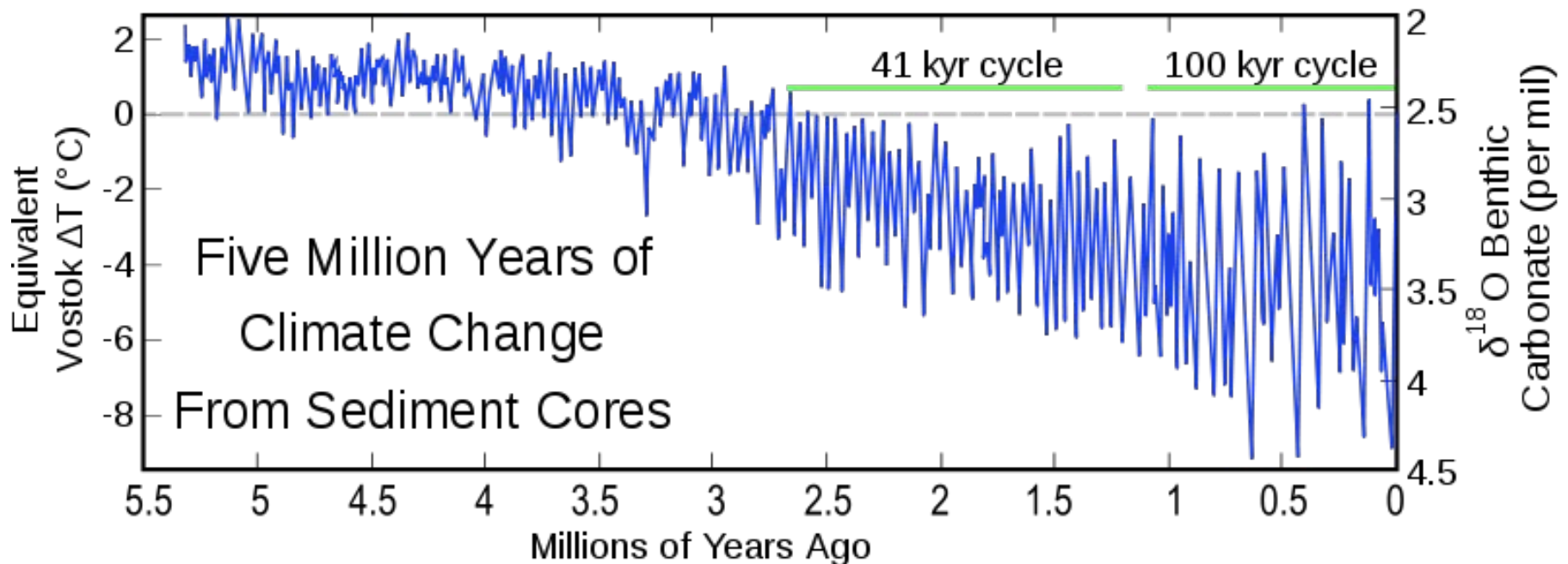
Find little dead creatures, and analyze their composition and deduce things from the relative abundance of different types

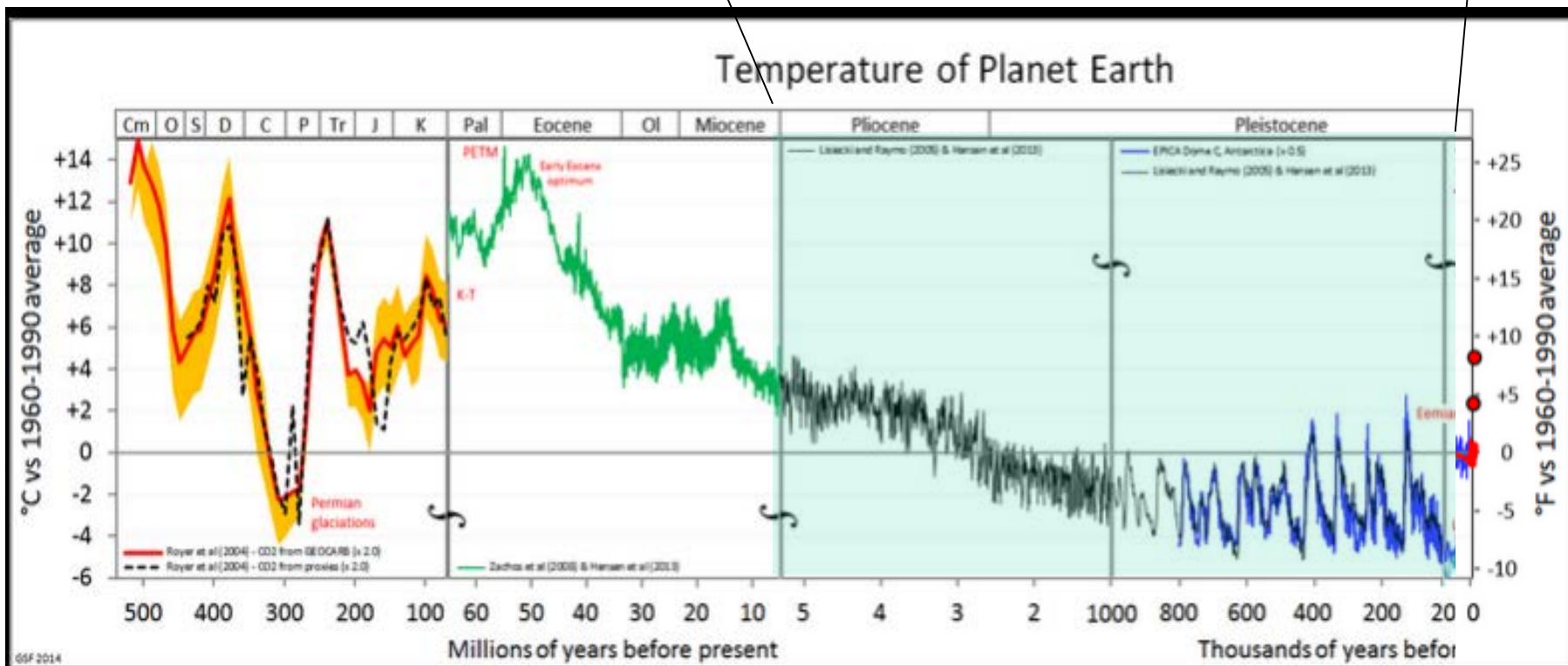
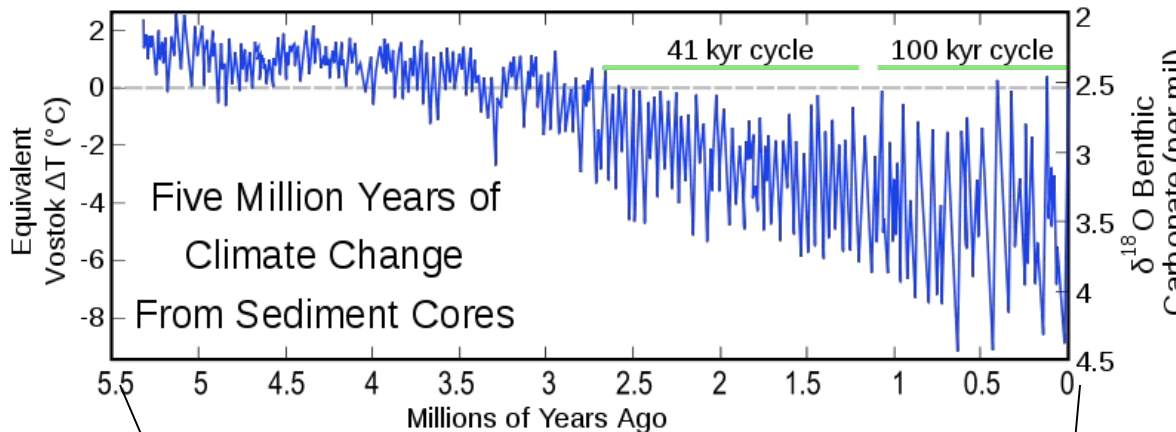




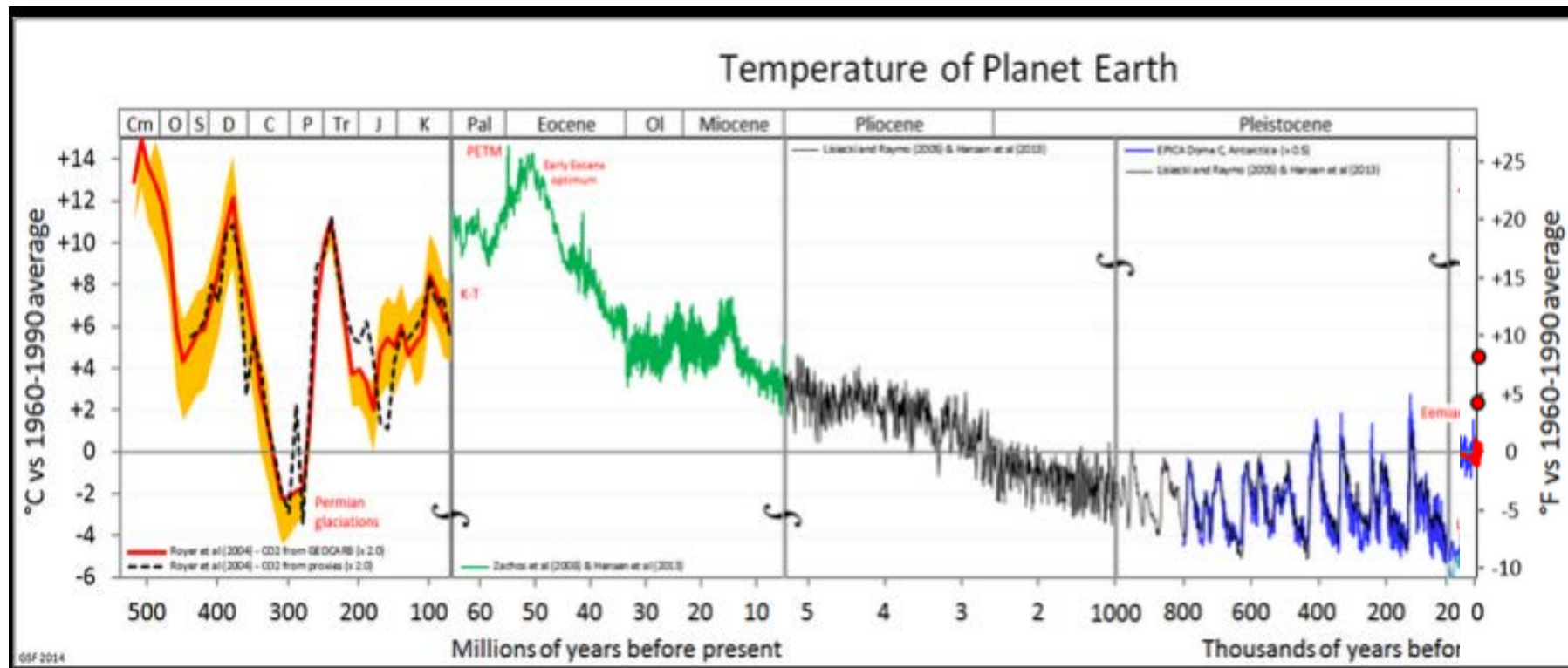
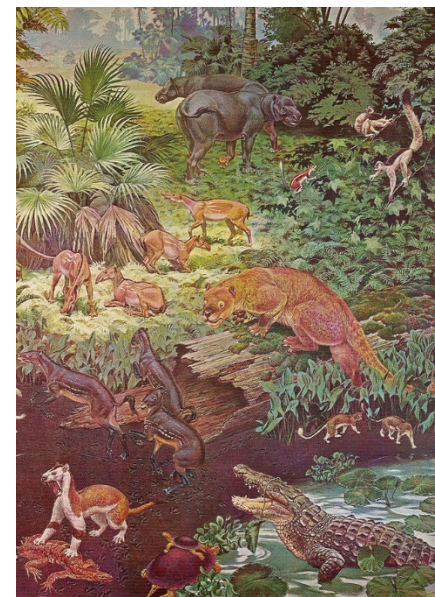


Earth has been cooling over the past 5.5 million years (My). Ice cycles started around 2.8 My ago, first with a periodicity of 41 Ky (obliquity), then changed to 100Kyr around 1 My ago.



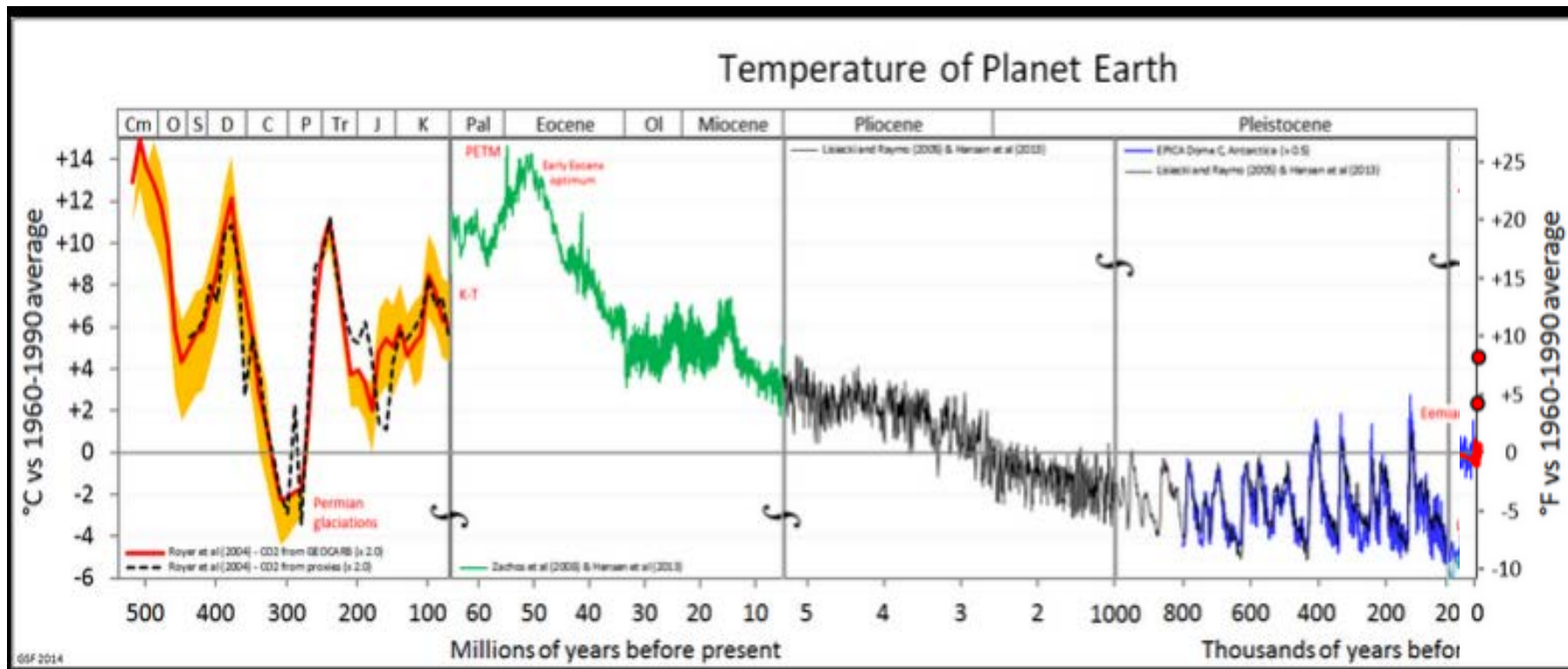


In fact Earth has been cooling over the past 50 My, and was warmest during the Eocene. CO<sub>2</sub> was higher than 500ppm  
Tropical plants and animals in North America

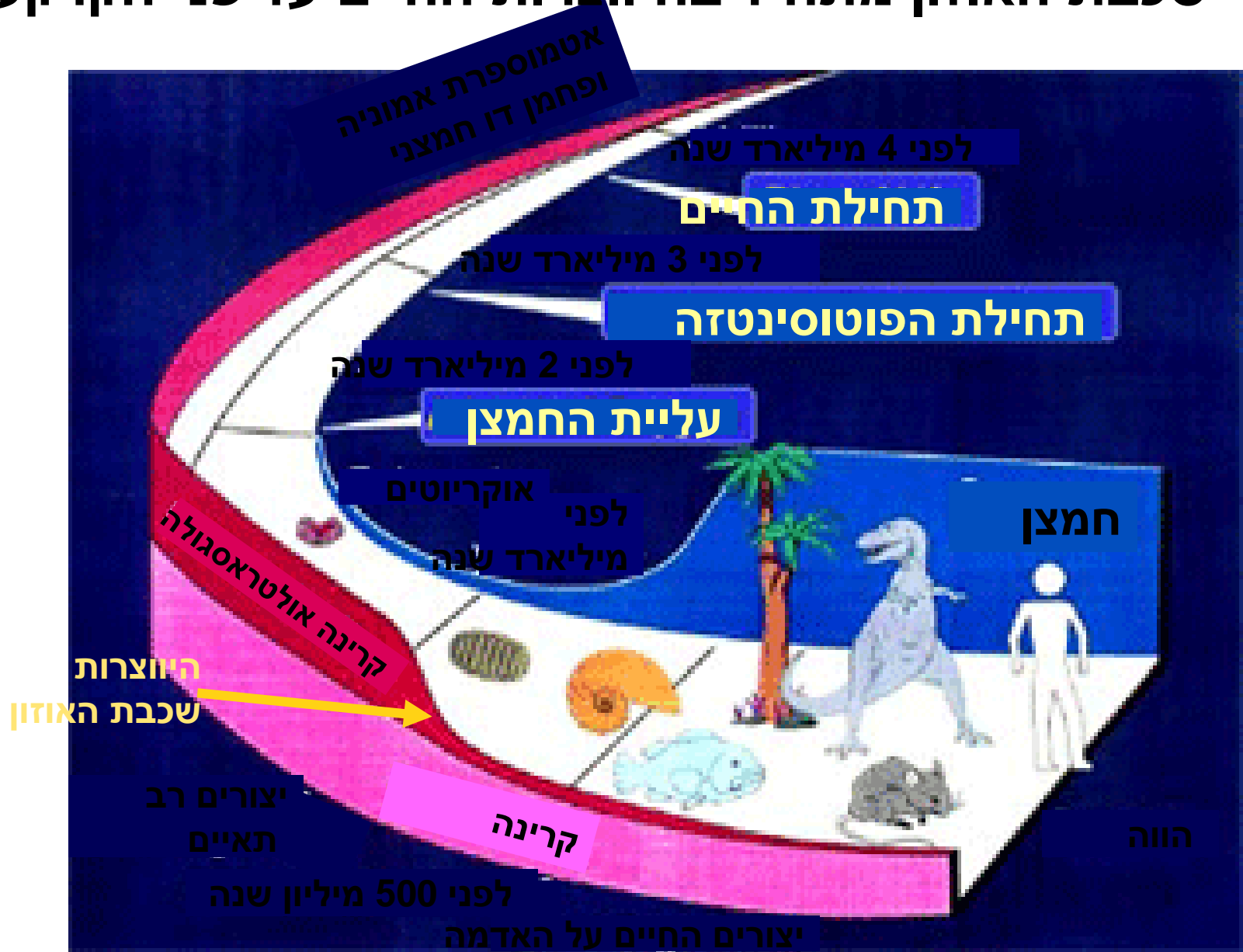




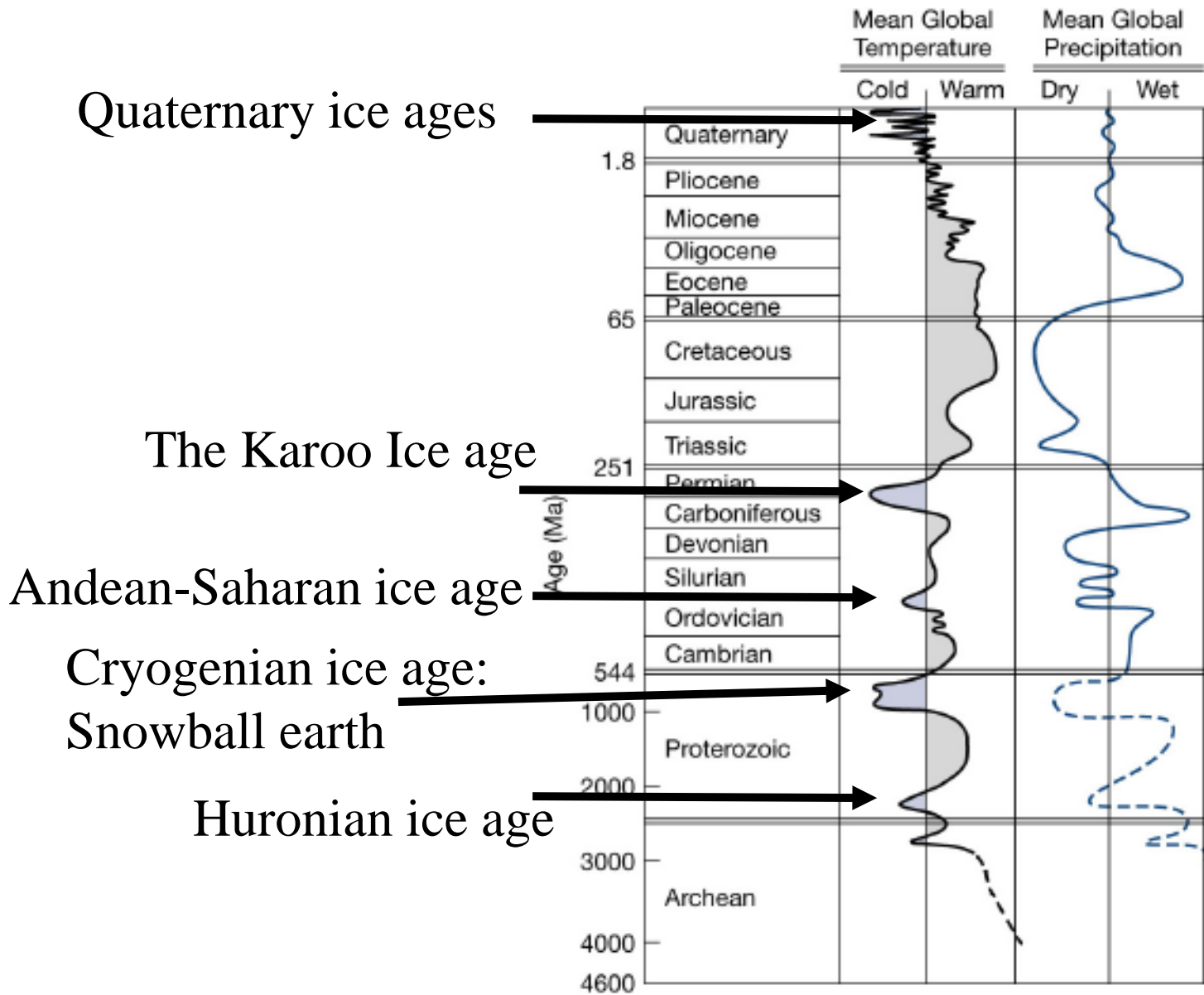
There was a prolonged ice age around 260-360My ago – the Karoo ice age. This is thought to be related to the reduction of CO<sub>2</sub> following the evolution of land plants and increase of oxygen in the atmosphere



# סיפור שכבת האוזון מתחיל בהיווצרות החיים על פני הקרקע

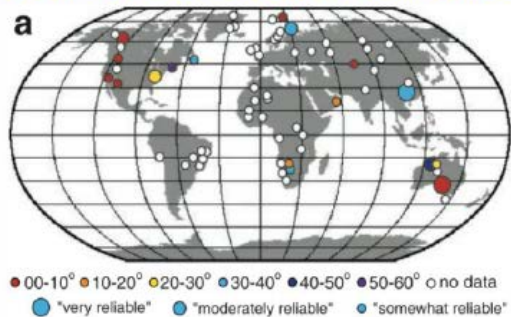




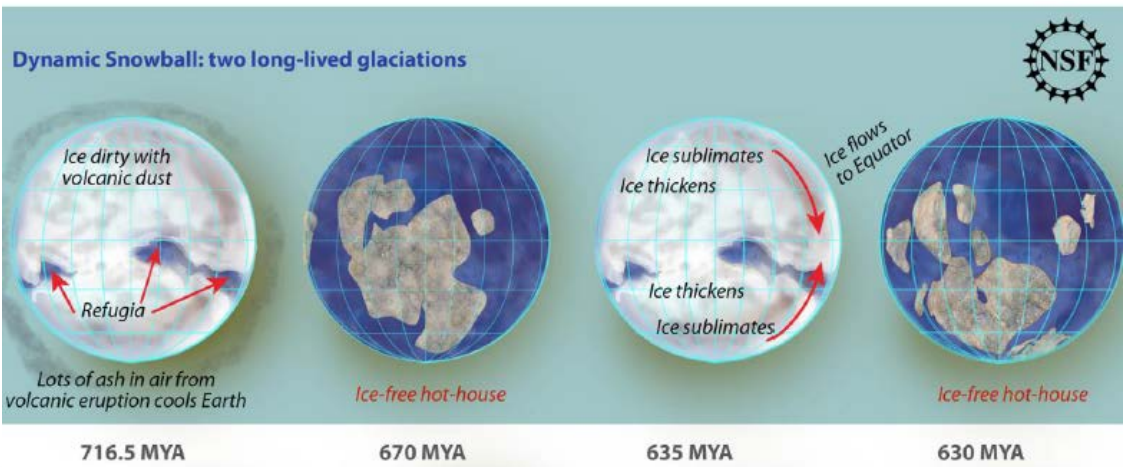


Snowball Earth: during the Cryogenian ice age – 630-850 My ago – there was at least one period, maybe 2-3 during which the Earth's surface was covered with ice. Still debated if fully ice covered or not so at equator.

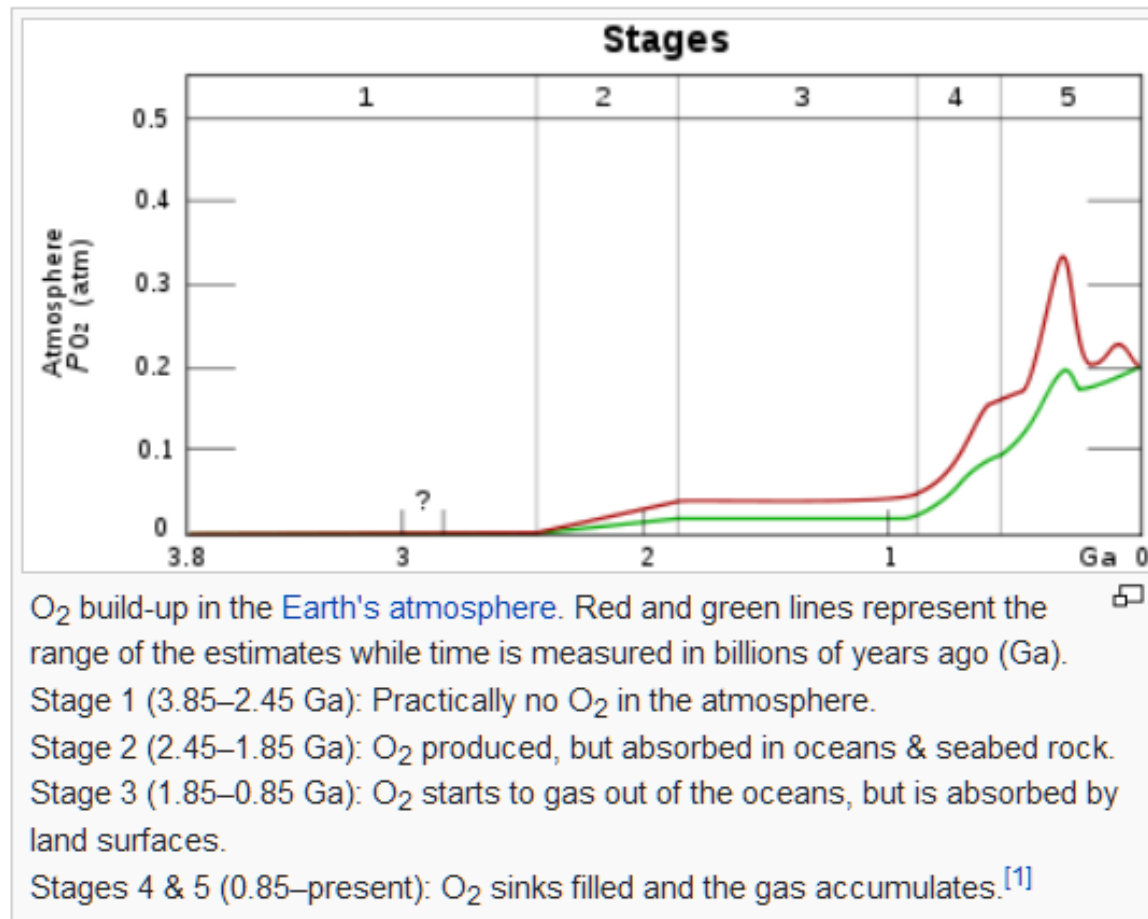
Glacial deposits at low paleo-latitude



Glacial deposits: Dropstone



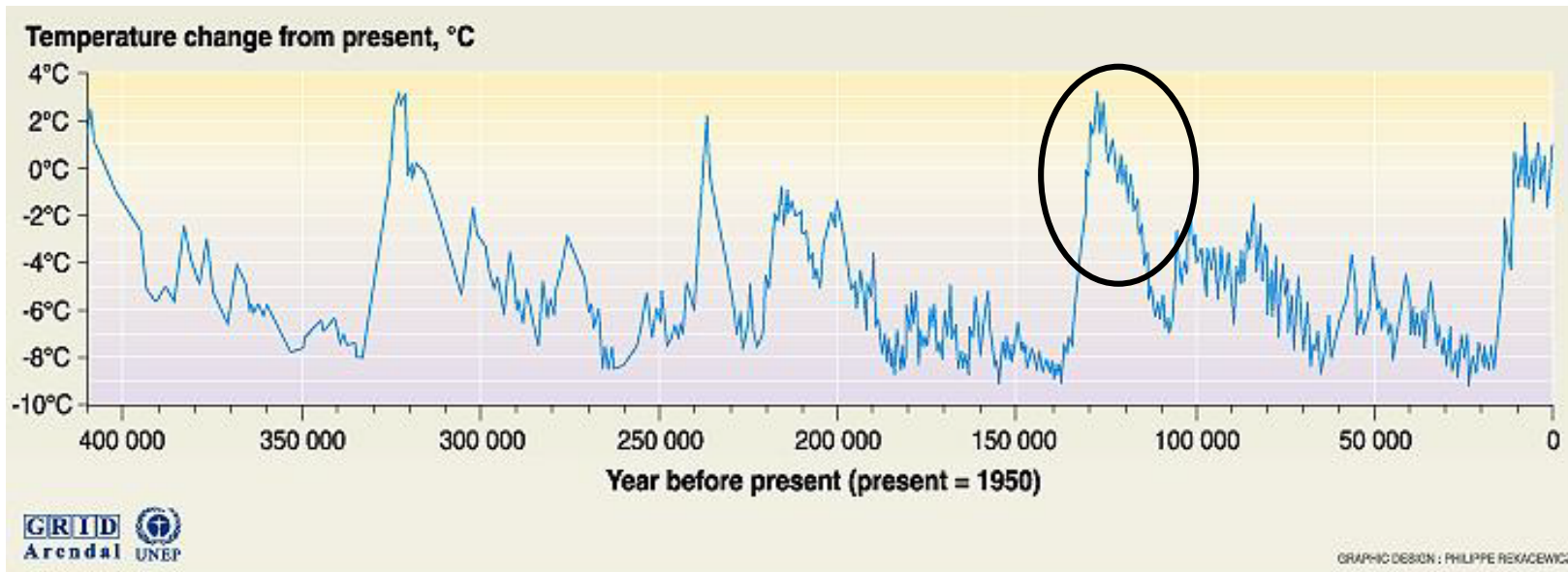
The earliest Huronian ice age- may have been onset by Methane reduction following the Great Oxygenation Event around 2.3 Billion years ago



Back to closer times the last interglacial, around 125Kyr ago, (the Emian) was slightly warmer than present, and lasted around 10Kyr

Want to connect to present and to expected future natural variations

**Emian  
interglacial**





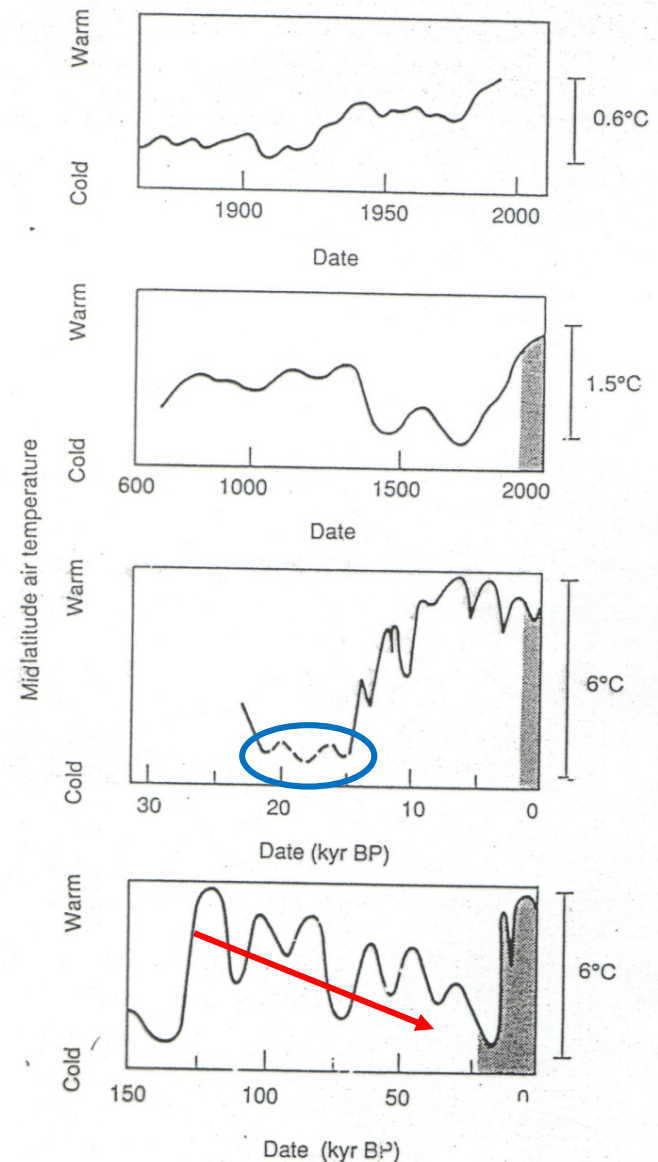
# Recent Climate Changes (last 150 kyr)

*115 kyr – 18 kyr: Generally declining temperatures superimposed with large fluctuations in the climate.*

*22-14 kyr: Last Ice Age, maximum ~18kyr ago.*

*The “Laurentide Ice Sheet” over N. America and the Scandenvian Ice Sheet” over parts of Europe, reached a maximum extent 22-13kyr ago.*

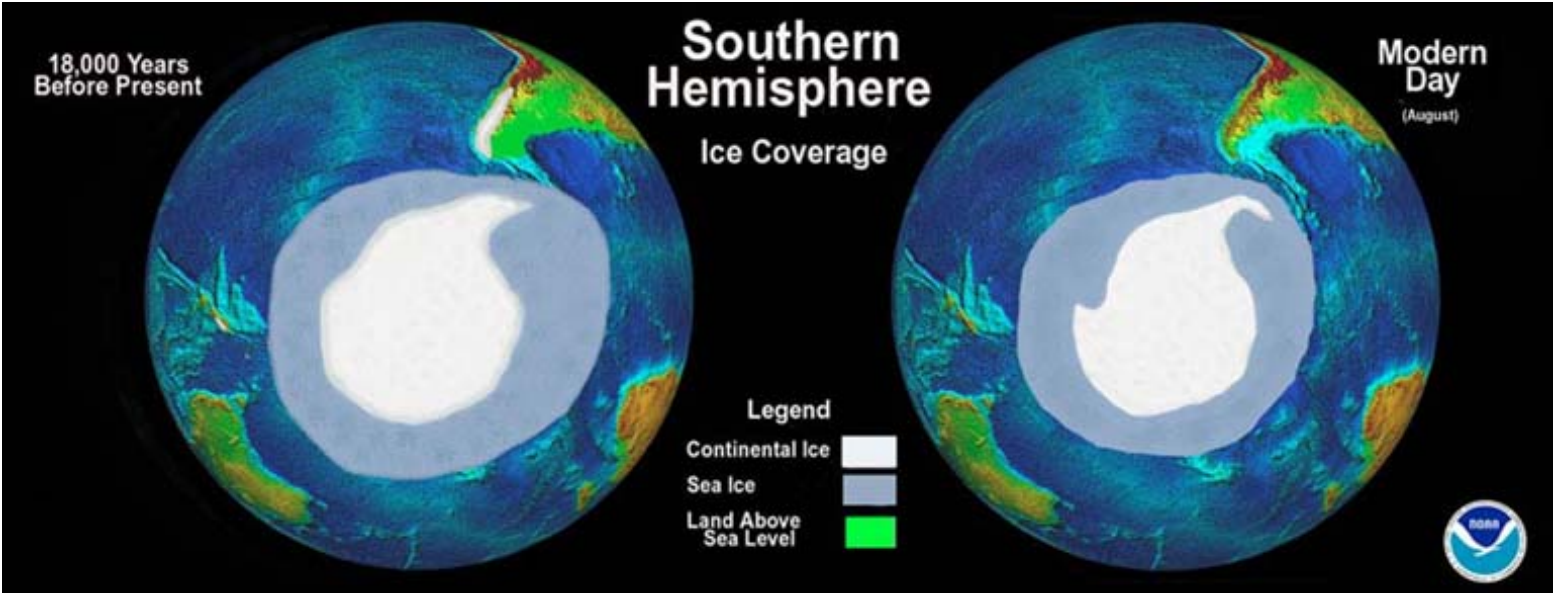
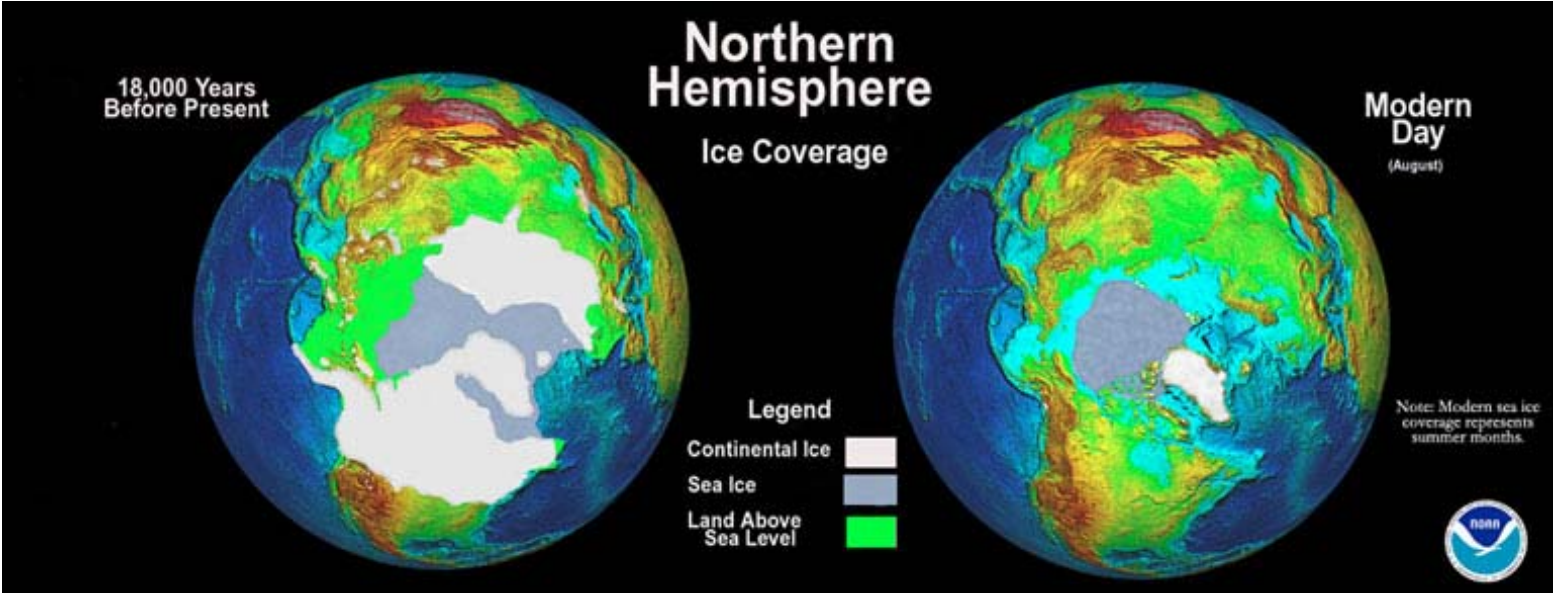
*“Cordilleran Ice Sheet” over western N. America reached maximum 14kyr ago.*



During the last ice age maximum:

New York buried under 1 km of ice.  
Maximum area of NH ice sheets was  
~90% of maximum ice cover during  
the last 1 MY. Sea level dropped ~85  
m; sea surface temperatures (SST) fell  
as much as 10C in mid latitudes of the  
N. Atlantic, 3C in the Caribbean.



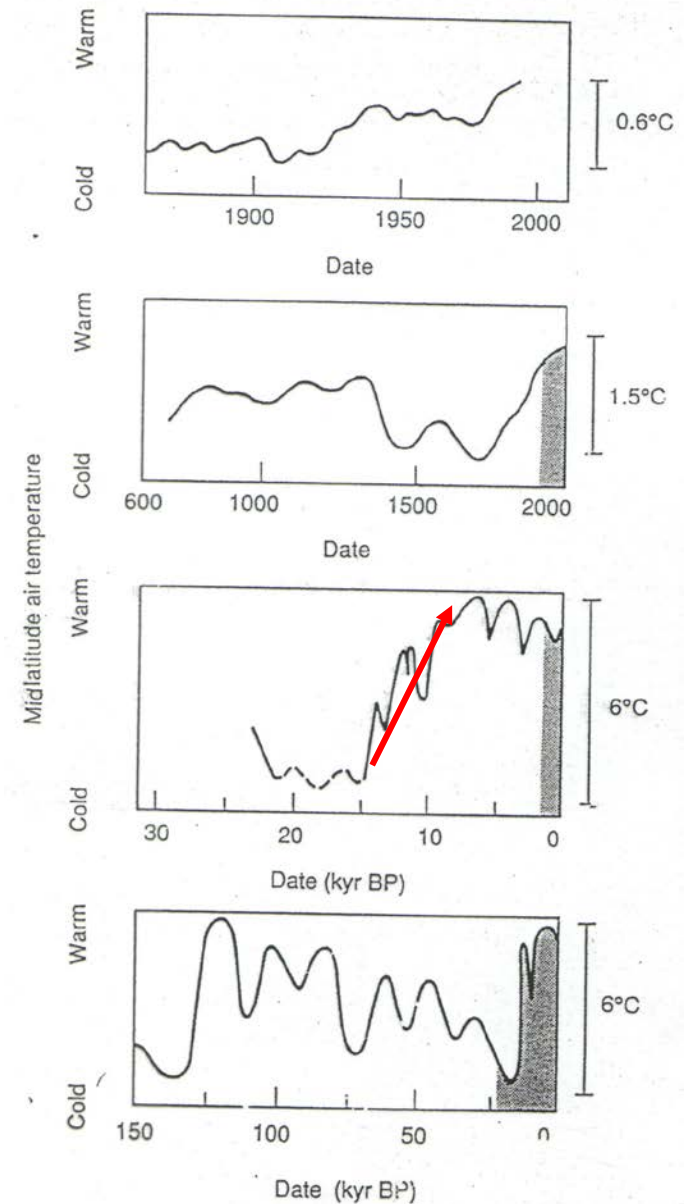


## 14-7kyr: Deglaciation.

Widespread deglaciation began rather abruptly ~14,000 years ago.

Cordilleran Ice Sheet melted rapidly and was gone by ~10,000 years ago.

Scandinavian Ice Sheet lasted only slightly longer. By 8,500 years ago the conditions in Europe had reached their present state; in N. America by ~7,000 years ago.

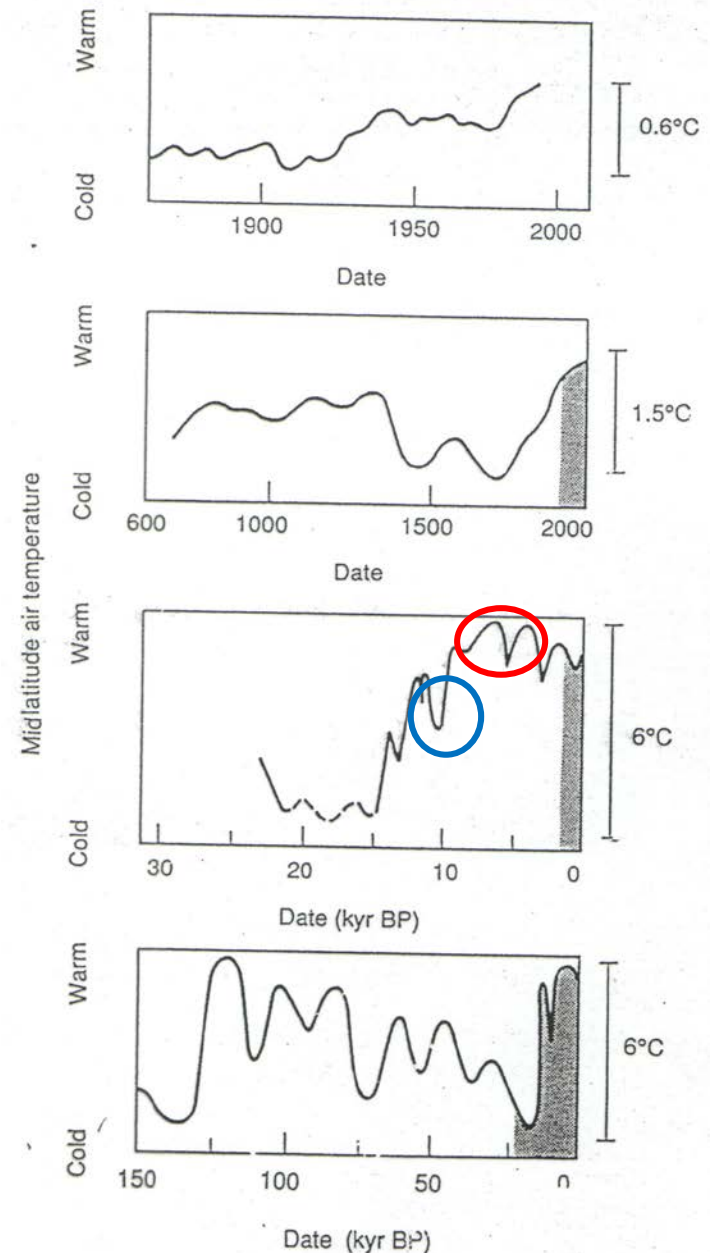




**10.8-10.1kyr ago: Younger Dryas cooling** Interspersed in the general warming were periods of widespread cooling and glacial advance spaced about 2,500 years apart. The Younger Dryas Event established itself within 100 years and lasted 700 years.

**7-4kyr ago: Post Glacial Optimum**

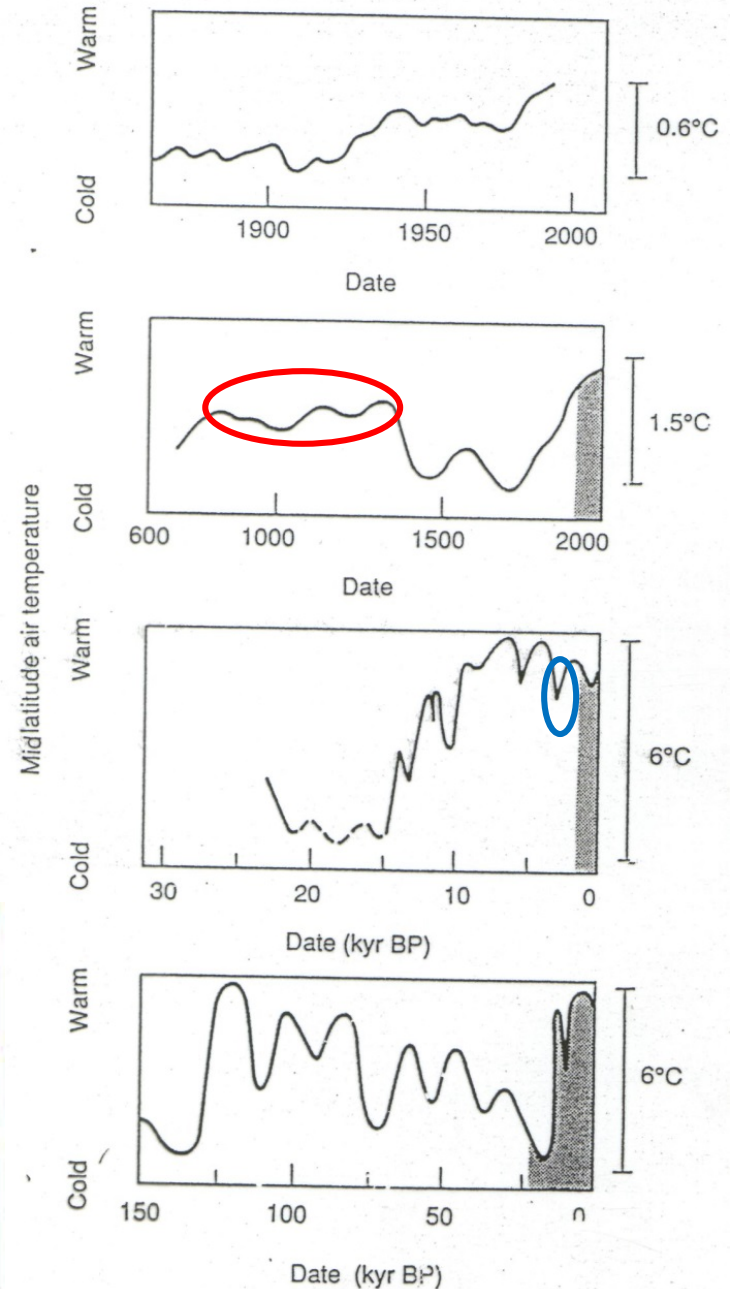
Warmest period culminated between 7-5kyr ago. Sea level rose rapidly maybe reaching its peak ~4kyr ago— higher than today by ~3 meters. Mainly due to reduction of ice on land, but sea ice in Arctic Ocean also reduced. Both Antarctica and Europe (in summer) 2-3C warmer than they are today.



## 2.9-2.3kyr ago: Iron Age Cold Epoch (900-300 BC)

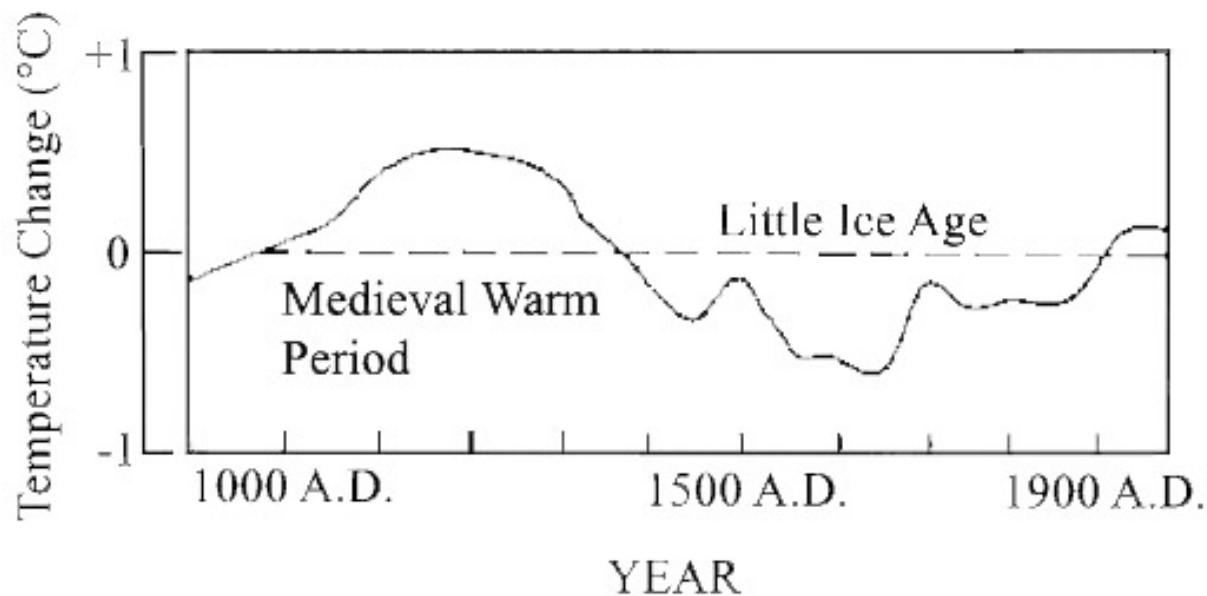
## 900-1200 AD: The medieval Warm Period (the Medieval Optimum)

Weaker and shorter warming than post-glacial optimum. Melting of the Arctic ice – founding of the Norse Colonies across the northern N. Atlantic to America (Iceland and Greenland). Summer temperatures probably 1C higher than today in western and central Europe. Lasted until 1300 AD.



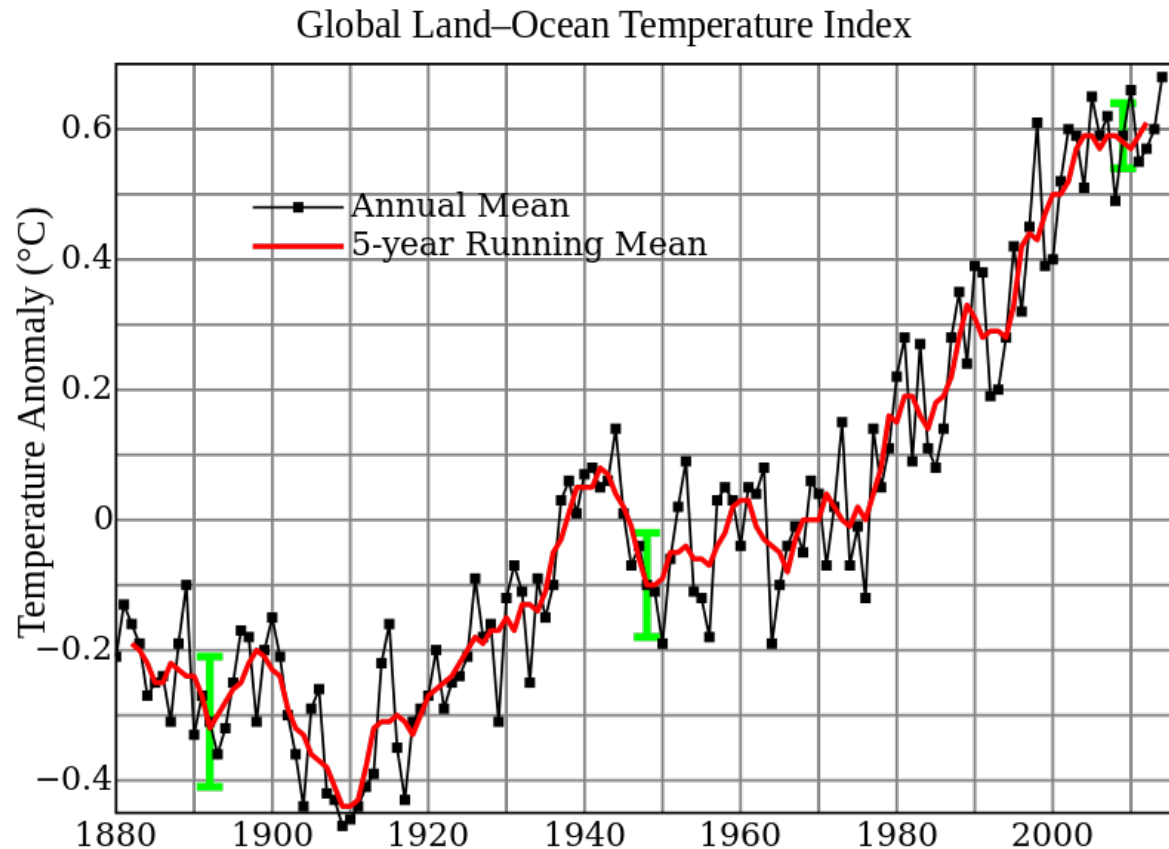
## 1430-1850: The Little Ice Age

Two peaks (1400 and late 1600s). Most harsh in Britain during the second half of 17<sup>th</sup> Century. Thames river often covered with ice. Arctic pack ice expanded south with important effects for Greenland and Iceland. By 1780-1820 the temperature across the N. Atlantic north of 50N was 1-3C colder than today. Glaciers advanced in Europe, Asia and N. America. However, the Southern Hemisphere seems to have partly escaped this cold period, i.e., not a global phenomenon. In fact, Antarctic may have been warmer than today during the period 1760-1830.

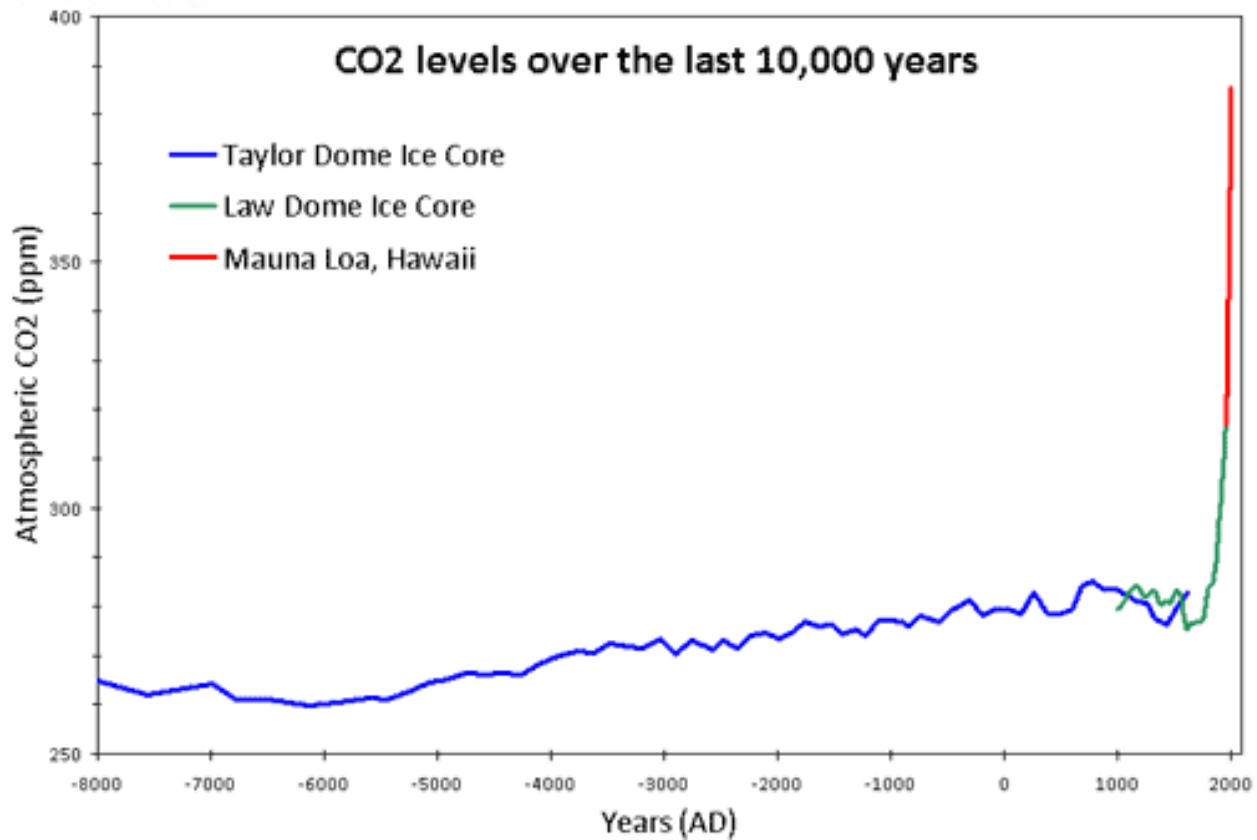
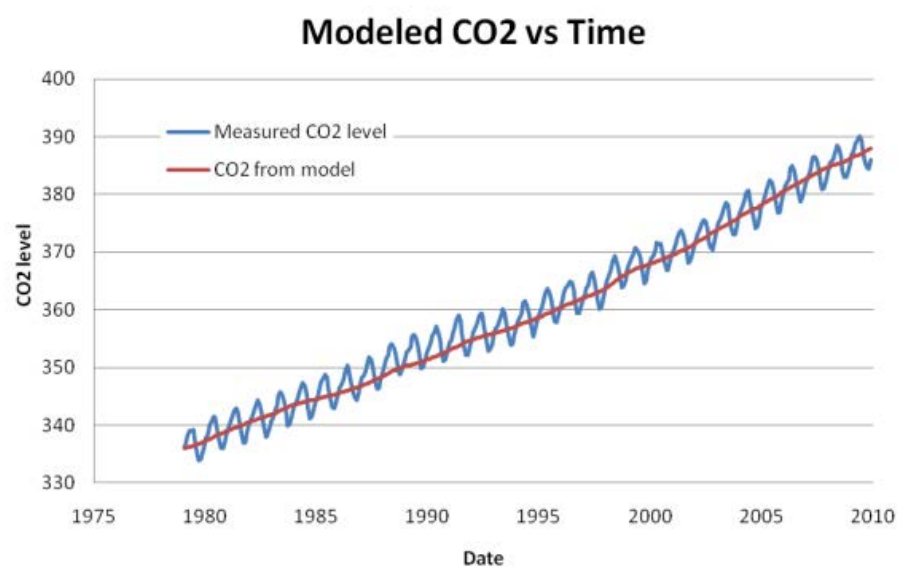
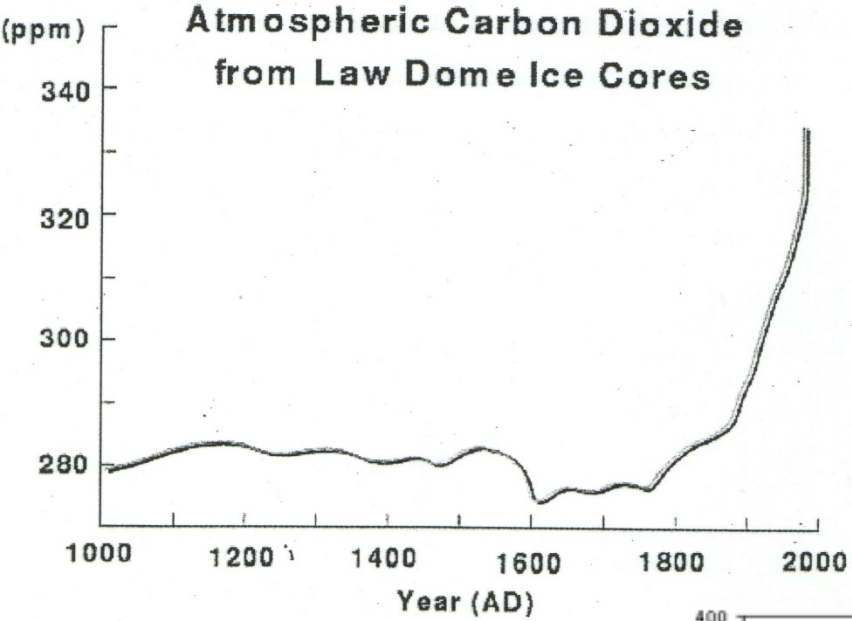


## 1880-present: Global Warming Trend

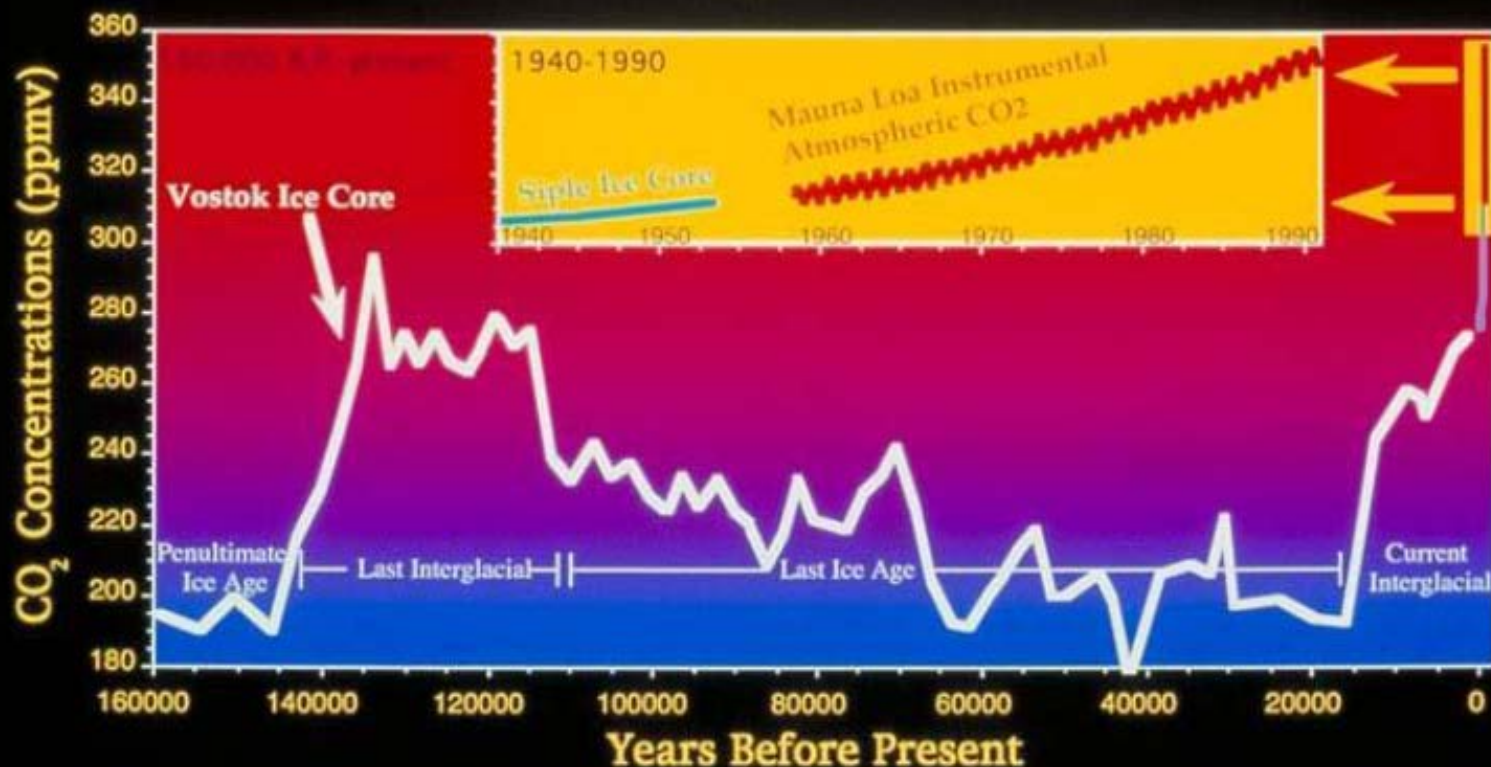
- ~0.8C global warming since 1880
- 1940-1960 - a slight cooling in the global temperature record.
- Since 1970 the rate of warming has increased.
- Ten of the hottest years since the beginning of instrumental record (~1860) have occurred since 2000: 2014 in first place. 14 of hottest 15 years in 21<sup>st</sup> century.







**Measurements of the Greenhouse Gas Carbon Dioxide (CO<sub>2</sub>) Have Been Taken from Several Ice Cores. Data from the Vostok and Siple Cores Show That CO<sub>2</sub> Concentrations Are Currently at Their Highest Level in the 160,000 Year Record.**



**Sources:** Vostok: Barnola *et al.* (1987); Siple: Friedli *et al.* (1986); Mauna Loa: Keeling and Whorf (1991).

# To relate very recent warming to more distant past we need to discuss the observational methods

## 1. *Instrumental Observations:*

- barometer/thermometer (invented in 17th Century)
- temporal resolution: hours/days
- length of record: ~100-200 years.

## 2. *Historical Records:*

- books, manuscripts, paintings, logs, ....
- temporal resolution: years
- length of record: up to 1000 years.

## 3. *Tree Rings:*

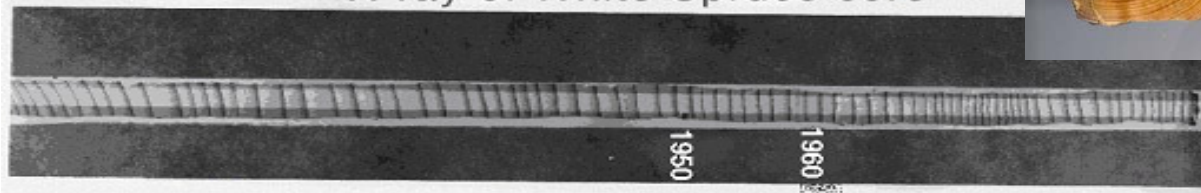
- respond to temperature, moisture and sunlight
- annual variations: ring widths.
- useful in mid and high latitudes
- length of record: up to 1000 yrs (rare cases 8000 yr)
- at low latitudes not useful due to small annual T and P variability



# Tree Rings

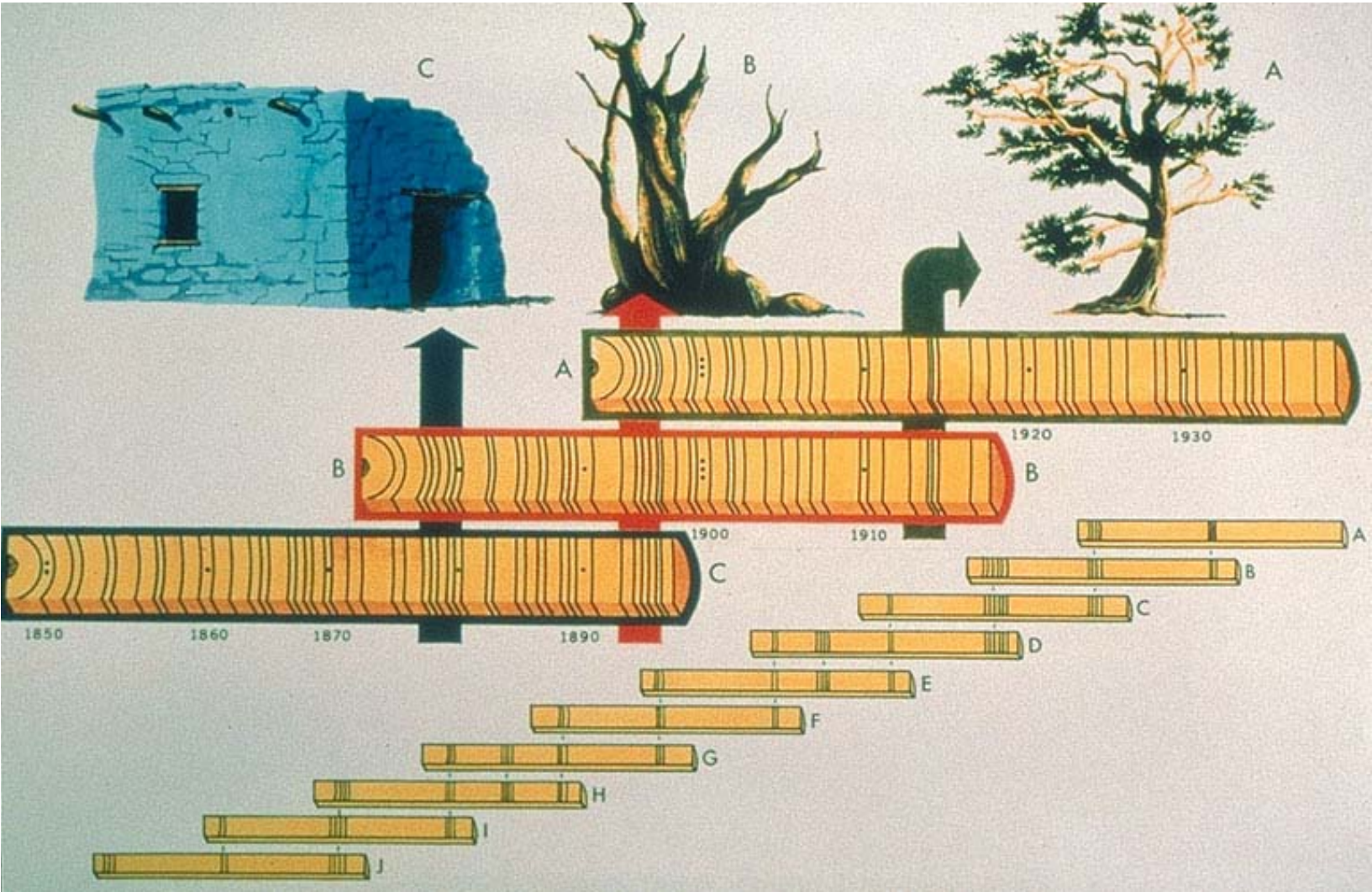


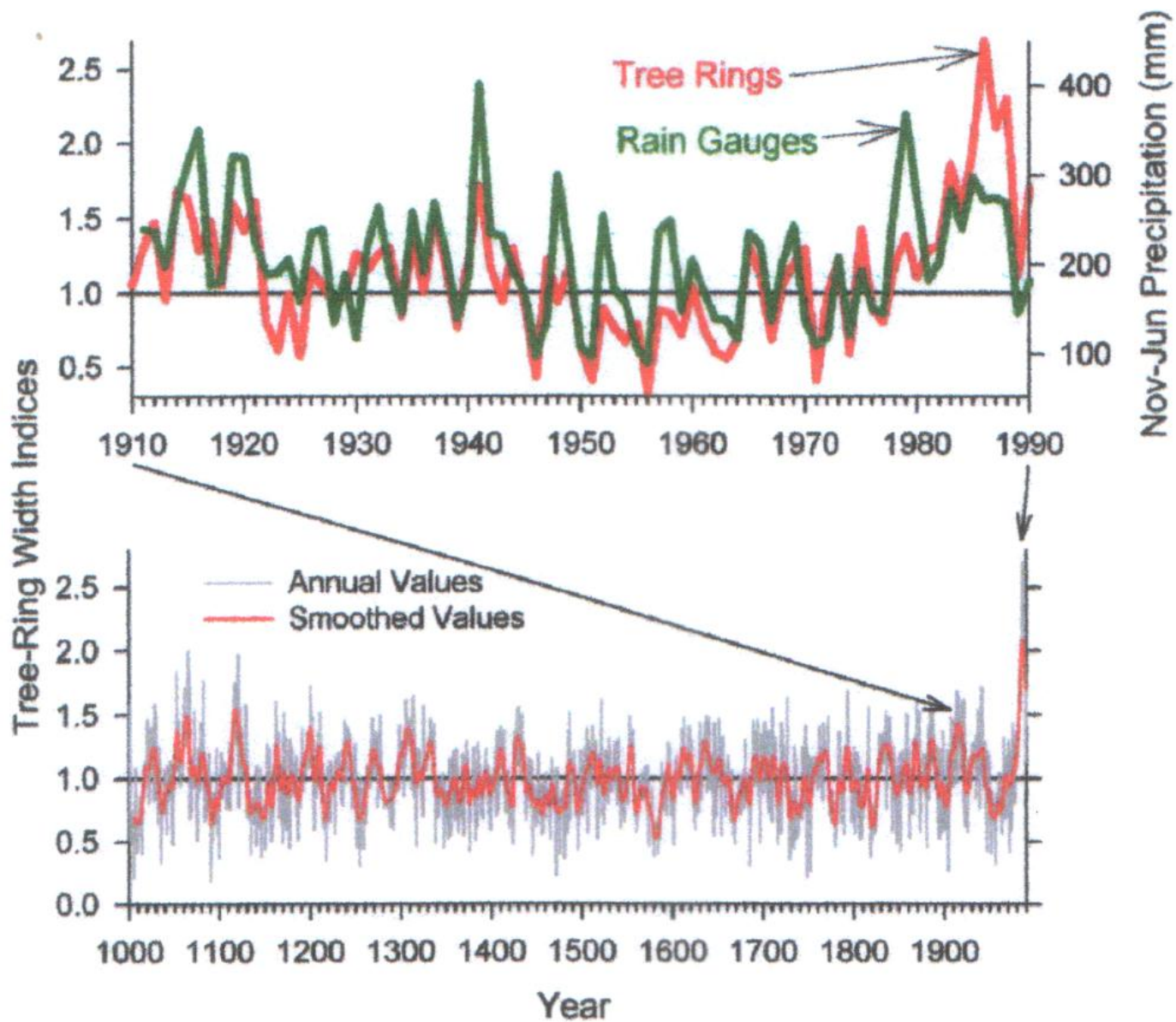
X-ray of White Spruce core



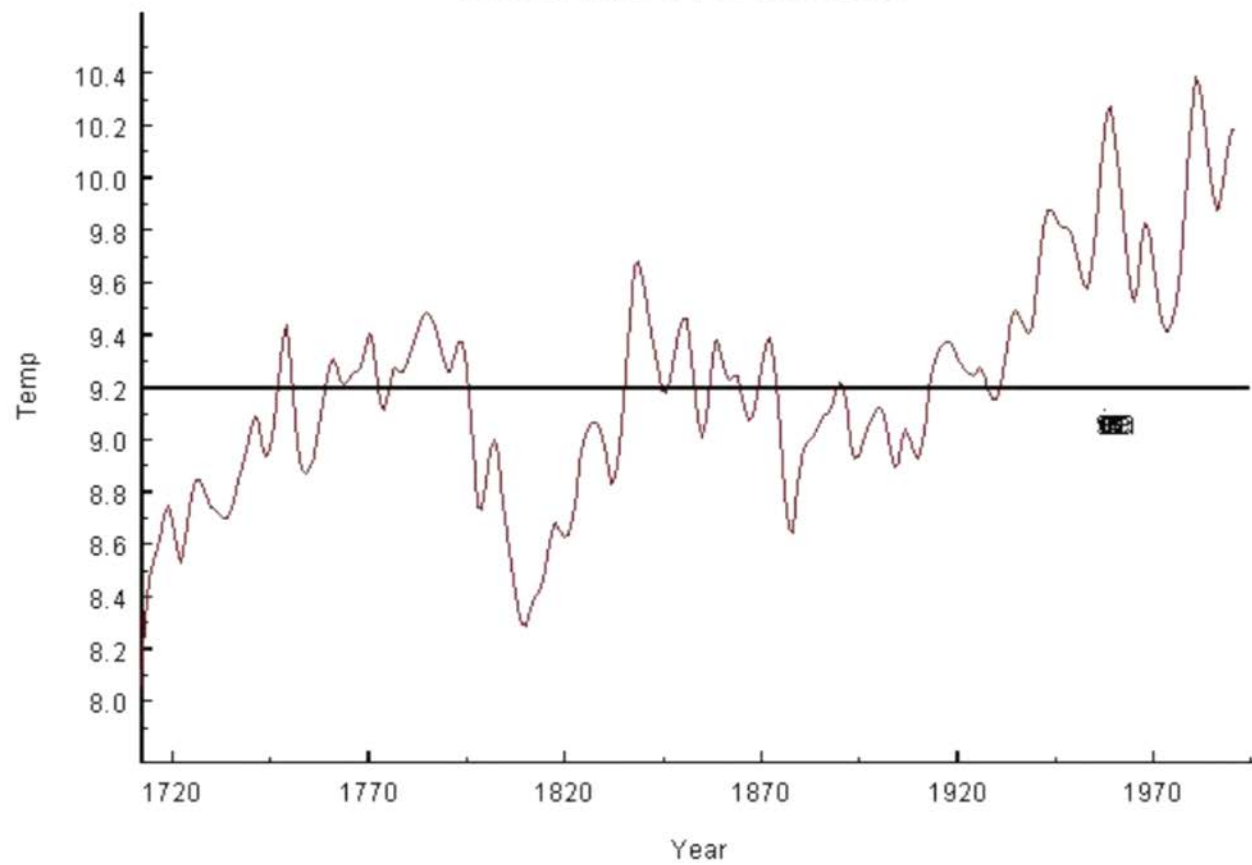


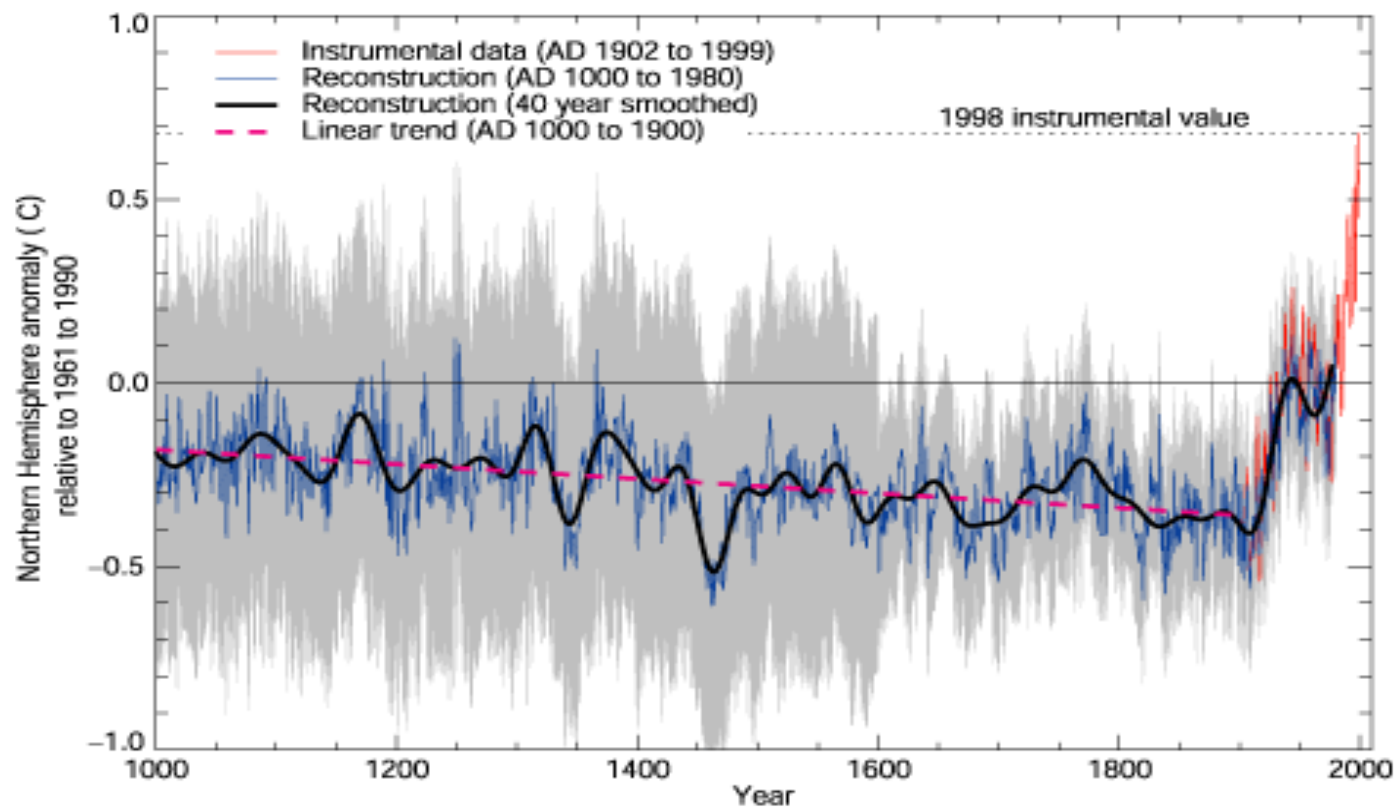
# Cross dating: Extending the time series back in time





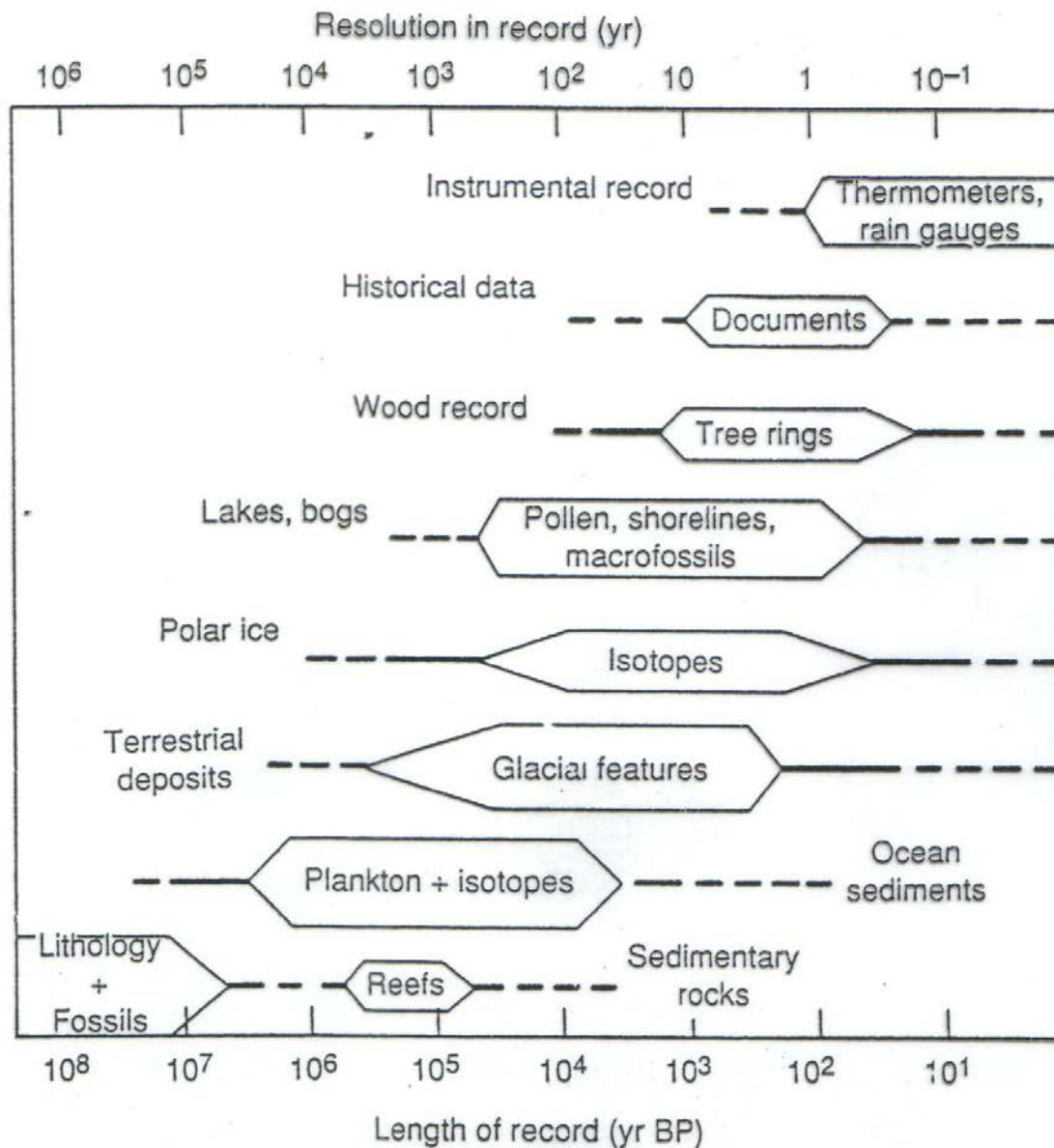
Reconstructed mean May-July temperature, Kenai Peninsula, Alaska  
smoothed with a 13-year lowpass filter





Millennial Northern Hemisphere temperature reconstruction (blue) and instrumental data (red) from AD 1000 to 1999, adapted from Mann et al. (1999). Smoother version of Northern Hemisphere series (black), linear trend from AD 1000 to 1850 (purple dashed) and two standard error limits (grey shaded) are shown.





**Figure 10.1** The time span of various climatic records from decades to hundreds of millions of years. The resolution in each record is also indicated, given the rule of thumb that it is about 1% of the length of the individual climatic record. The thickness of each polygon is an attempt to reflect the relative contribution for each contributing record at each time in the past. (Adapted with permission from T. Webb III, J. Kuczbach, and F. A. Street-Perrott, in *Global Change*, T. F. Malone and J. D. Roederer, eds., pp. 182-218. Copyright 1985 by Cambridge University Press, UK.)

# Temperature of Planet Earth

